

## Further contributions to the early Miocene forest vegetation of the Galatian Volcanic Province, Turkey

N. Neslihan Acarca Bayam, Ünal Akkemik, Imogen Poole, and Ferdi Akarsu

### ABSTRACT

Ten new petrified forest sites have been discovered on the Galatian Volcanic Province (GVP). The silicified woods, preserved in the lower Miocene Hançili and Pazar formations, represent a diverse plant assemblage that grew across western and central GVP during the Burdigalian, ca. 20–17 Ma. The assemblage comprises a total of 233 petrified woods collected from the 10 sites in and around Ankara (Beypazarı, Çamlıdere, Kızılcahamam) and Bolu (Kırıbrışık). Taxonomic identifications of the fossil wood floras suggest four distinct vegetational associations can be recognized on the GVP: (1) swamp associations of *Glyptostrobus* Endlicher, *Sequoia* Endlicher and *Taxodium* Richard; (2) riparian forest comprising palms, *Salix* L., *Liquidambar* L. and *Quercus* L. section *Ilex*; (3) mesic-xeric woodland comprising *Acer* L., *Quercus* L., *Pinus* L. and *Juniperus* L.; and (4) mixed coniferous-angiosperm forest comprising *Picea* A. Dietrich, *Cedrus* Trew., *Pinus*, *Juniperus* L. and *Podocarpus* L'Herit (ex Pers.) alongside. These associations support the idea that the landscape was varied and the climate warm and frost free.

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**Keywords:** Miocene; silicified wood; petrified forest; wood identification; Galatian Volcanic Province; Köroğlu Volcanics, Turkey

Submitted: 4 September 2017 Accepted: 2 November 2018

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Acarca Bayam, N. Neslihan, Akkemik, Ünal, Poole, Imogen, and Akarsu, Ferdi 2018. Further contributions to the early Miocene forest vegetation of the Galatian Volcanic Province, Turkey. *Palaeontologia Electronica* 21.3.40A 1-42. <https://doi.org/10.26879/816>  
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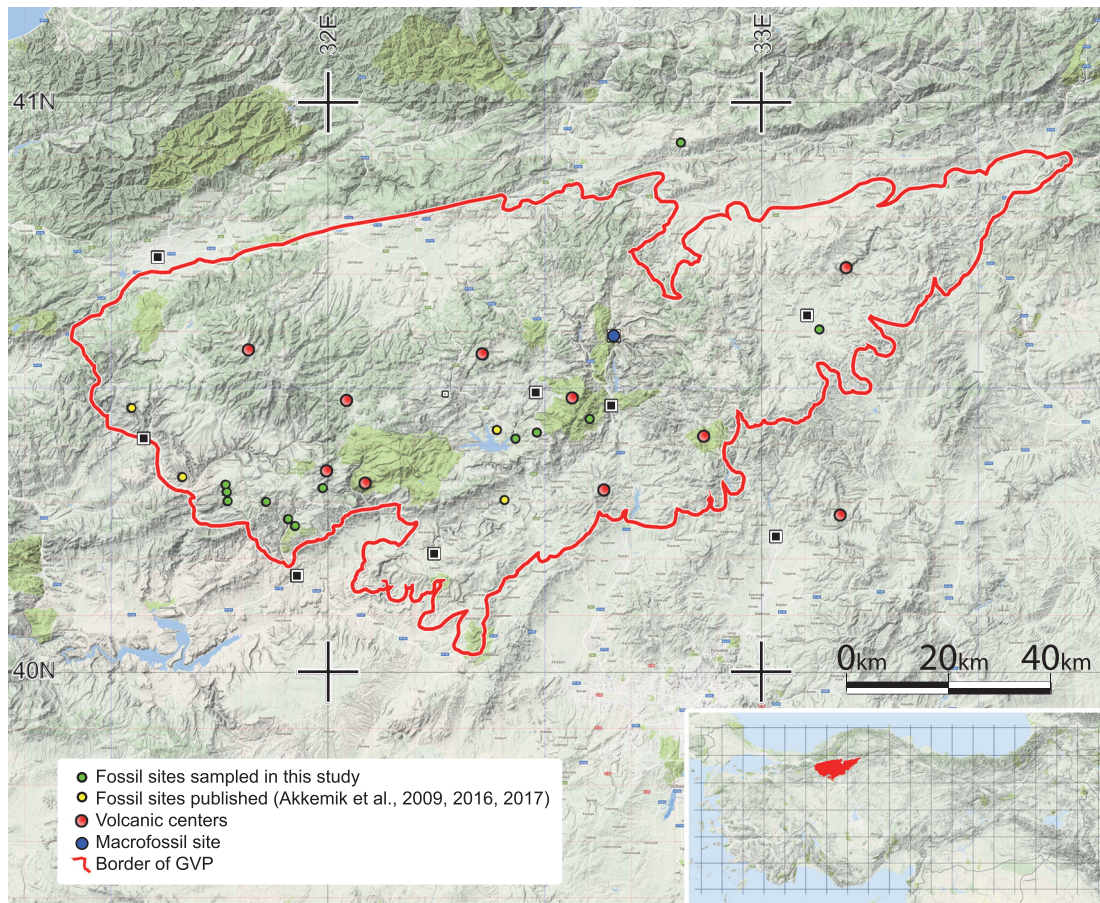
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## INTRODUCTION

The Galatean Volcanic Province (GVP) (or Köroglu Volcanics) is a major Neogene volcanic belt covering an area of approximately 12,800 km<sup>2</sup> (Koçyiğit et al., 2003) and located in central Anatolia in the northwestern part of Turkey (Toprak et al., 1996). The GVP and three other volcanic provinces in western, central and eastern Anatolia together form the four main Neogene and Quaternary volcanic fields in Turkey (Kazancı, 2012) that evolved as a consequence of the tectonic collision between the African Plate and the Eurasian Plate (Wilson et al., 1997; Tankut et al., 1998; Rögl, 1999; Koçyiğit et al., 2003). Most of the volcanic products of the GVP belong to the eruptions from

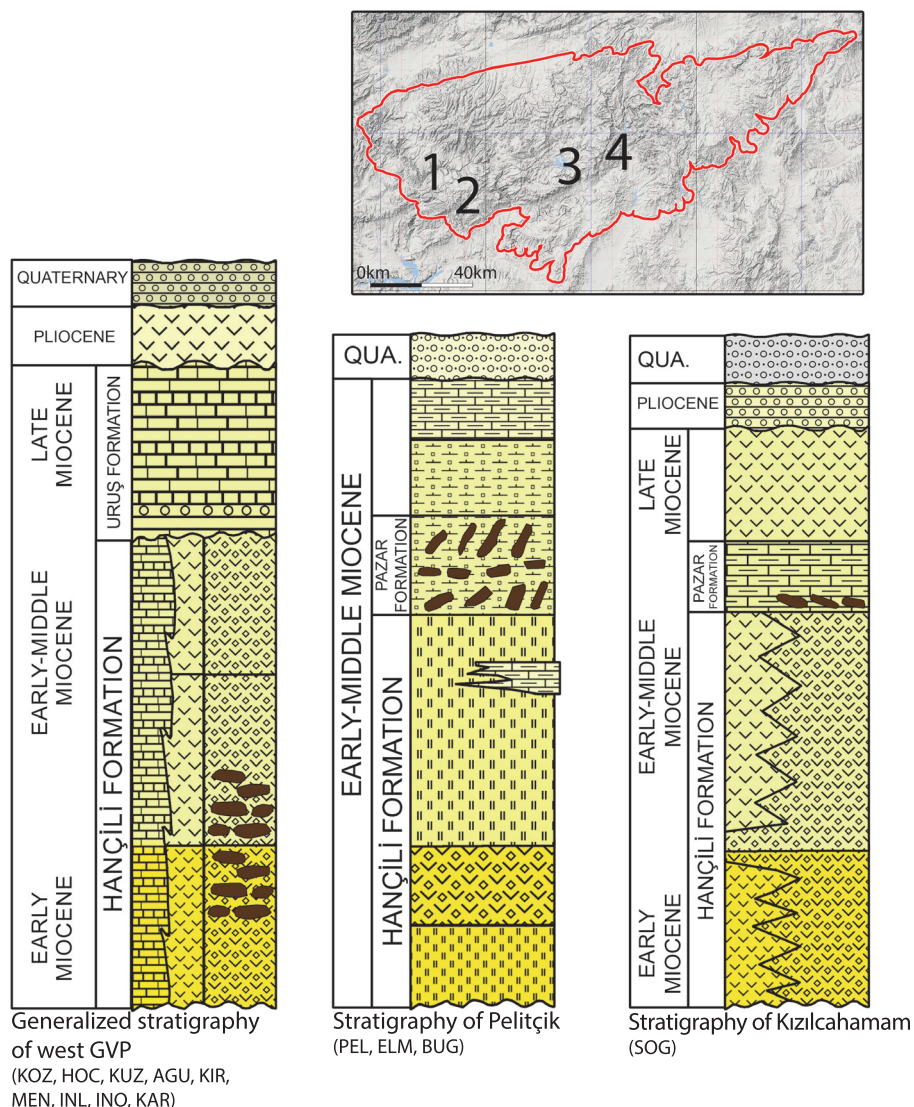
nine eruption centers (Figure 1) so far identified, which grade from basalts to andesite and rhyolites (Kazancı, 2012). Rocks were deposited during three major volcanic phases: (1) late Campanian-Lutetian subductional phase; (2) post-Lutetian collisional phase; and (3) Oligocene-Miocene post-collisional phase (Koçyiğit et al., 2003).

Keller et al. (1992), Wilson et al. (1997) and Adıyaman et al. (2001) suggest that during the Oligocene-Miocene volcanic phase three episodes of volcanism took place across the GVP: the first around 22–19 Ma, followed by a second episode between 18–17 Ma and a third episode of volcanism during the late Miocene (10–9 Ma). Potassium-argon dating of volcanic rocks determined that the major phase of volcanic activity took place



**FIGURE 1.** Map showing the area covered by the Galatian Volcanic Province (GVP) (or Köroglu Volcanics). Sites discussed in this work are marked by green circles: ELM, Elmali Village; SOG, Soguksu National Park; BUG, Bugralar Village; INO, Inozu Valley South Side; INL, Inozu Valley North Side; KAR, Karasar Village; MEN, Menciler Plateau; KIR, Kiraluc Sitene near Nuhhoca Village; AGU, Asagiguney Village; KUZ, Kuzca Village. Yellow circles (PEL, Pelitcik Village; GUD, Gudul; HOC, Hoçaş Village; KOZ, Kozyaka Village) indicate those sites discussed in Akkemik et al. (2009, 2016, 2017). The sites located in the western part (INO, INL, KAR, MEN, KIR, AGU, KUZ, HOC and KUZ) are from early-middle Burdigalian and Hancili Formation (Altun et al., 2002; Akbaş et al., 2002); the sites in the central part (GUD, BUG, ELM, PEL and SOG) are from middle-late Burdigalian and Pazar Formation (Kazancı, 2012; Sen et al., 2017).





**FIGURE 2.** Stratigraphic columns of the GVP showing the fossils (dark symbols) originating from the Hançili Formation of the western part (1-2) (Akbaş et al., 2002; Altun et al., 2002) and from within the pyroclastic layers of the Pazar Formation in the central part of the GVP at Çamlıdere-Pelitcik (3) and Kızılcahamam (4). Columns redrawn from figure 3 for Kızılcahamam and figure 6 for Pelitcik of Kazancı (2012) and figure 2 of Şen et al. (2017).

during the early Miocene (20–17 Ma; Wilson et al., 1997), which resulted in the accumulation of thick volcanic and volcanoclastic rocks intercalated with continental sedimentary rock (Toprak et al., 1996).

The fossil wood material from the western part of the GVP originated from the lower part of the early to early-middle Miocene Hançili Formation (Akbaş et al., 2002; Altun et al., 2002) (Figure 2), a sequence of conglomerates, coals and organic-rich shales and claystones reflecting deposition in lacustrine and coastal environments under warm, humid climatic conditions (Vosoughi Moradi et al.,

2016). Radiometric dating of the early Miocene volcanic basalts across the western part of the GVP have been undertaken by several workers: Toprak et al. (1996), summarizing Türkecan et al. (1991) and Keller et al. (1992), give the age of the exposed volcanoclastic rocks found in the southwestern part of the Kartalkaya-Köroğlu Complex near Seben (Figure 1) in the western part of the GVP as ranging from  $24.83 \pm 2.01$  to  $21.79 \pm 4.93$  Ma. Subsequently Akkemik et al. (2016) undertook radiometric dating of volcanic rocks taken from around the silicified wood samples that originated

from the Seben-Hoçaş Fossil Forest (Bolu Province, Figure 1), just west of the sites sampled in this study. These returned an age of  $18.2 \pm 0.8$  Ma (early-middle Burdigalian). These data suggest an age range of ca. 24.8–18 Ma (Aquitania–Burdigalian) for the fossil woods found in the western part of the GVP.

The fossils found in the central part of the GVP were recovered from the Pazar Formation (Figure 2), a sequence of argillaceous limestones, marl and claystones with locally cherty limestone, sandstone and tuff, and occasional thin coal seams (Süzen and Türkmenoglu, 2000; Yavuz-Işık and Demirci, 2009; Sen et al., 2017). Tatlı (1975) stated that Pazar Formation was formed in lake and swamp environments. Sen et al. (2017) suggested that the Pazar Formation was deposited during the early and early middle Miocene. The fossils originated from the basal layers of the Pazar Formation at Kizilcahamam, but were more evenly dispersed throughout the formation at the Pelitcik sites (Figure 2). Representative basalt samples from the volcanic layers covering the central part of the GVP, near the towns of Özmus and Gündül (Figure 1), were analyzed by Wilson et al. (1997) and returned ages of  $18.2 \pm 0.5$  Ma and  $18.8 \pm 0.6$  Ma (both middle Burdigalian), respectively. Further basalt samples from Güvem and Çamlıdere (Figure 1) returned an age of middle to late Burdigalian ( $18.2 [\pm 0.5] - 16.9 [\pm 0.5]$  Ma, respectively) (Table 1) (Tankut et al., 1995; Wilson et al., 1997). Based on this evidence we consider the age of the fossils preserved in the same volcanic rocks covering the central part of the GVP to be middle to late Burdigalian and thus coeval or slightly younger than those fossils found in the west. Given that the two regions are interpreted to have formed from the same volcanic processes, and that there exists a degree of overlap in the ages of these two regions, until further dating is undertaken we consider the material found in the western and the central region of the GVP to be relatively contemporaneous.

The GVP is one of the most important areas for fossil forest preservation in Turkey (Akkemik et al., 2016 and references cited therein) at a time when particularly warm conditions favoured the growth of rich forest vegetation across the Northern Hemisphere (Mai, 1995; Zachos et al., 2008; Denk et al., 2011, 2017a, 2017b, 2017c). Fossilized animal remains alongside petrified wood, leaf, fruit and cones are commonly found in the lower Miocene volcano-sedimentary units in the central part of the GVP (Kasaplıgil 1977; Dernbach et al.,

1996; Akkemik et al., 2009, 2016, 2017; Denk et al., 2017b). Here evidence supports a mixed conifer-deciduous and evergreen angiosperm early Miocene vegetation growing in the Kizilcahamam-Çamlıdere-Güvem area (Figure 1), with *Sequoia*, *Taxodium* Rich., *Quercus* section *Ilex*, *Acer*, *Salix*, *Populus*, *Ulmus*, *Zelkova*, *Alnus* Mill. *Betula* L. *Fabaceae*, *Engelhardioideae* (Akkemik et al., 2009; Dernbach et al., 1996; Kasaplıgil, 1977; Denk et al., 2017b). To the south, near Gündül (Figure 1), Kasaplıgil (1977) and Akkemik et al. (2017) identified leaves and woods belonging to *Glyptostrobus* and *Glyptostroboxylon* Conwentz and concluded this taxon formed a main constituent of a prevailing swamp forest vegetation. Further north, near Pelitcik (Figure 1), a *Sequoia* type wood along with wood of *Taxodium* were identified as representative forest trees (Akkemik et al., 2009). The ‘*Sequoia* type’ wood is distinct in that it includes anatomical features of *Sequoia*, *Sequoiadendron* J. Buchholz and *Metasequoia* Miki and thus cannot be identified as belonging to either one of these modern genera; with better preserved samples a closer identification may be possible.

In the western part of the GVP, Akkemik et al. (2016) also described wood samples recovered from two early Miocene sites: one near the village of Hoçaş and the second near the village of Kozyaka (Figure 1). Both sites lie at an altitude of ~900 m close to the town of Seben (Figure 1) in the province of Bolu, and together are referred to as the Seben Fossil Forest. The Hoçaş site had preserved both in situ and in situ samples that include *Palmae*, *Salix* L., *Liquidambar*, *Quercus* section *Ilex* and *Juniperus*, *Acer*, *Cedrus* and *Picea*. The Kozyaka site also yielded samples of *Acer*, *Quercus* L. (*Ilex* type), *Cedrus*, along with *Salix/Populus* L., *Ulmus* L., and *Pinus* (Akkemik et al., 2016).

Although these studies have provided an indication of the local vegetation that grew across the western and central parts of the GVP during the early Miocene, ideas pertaining to the flora that grew here are still limited. The discovery of new petrified wood sites and the subsequent analyses of the wood samples discussed in context of the vegetational hypotheses presented in previous publications, forms the basis of this work and helps further understanding of the flora and ecological dynamics that prevailed across Anatolia during the Miocene.

## MATERIAL AND METHODS

Ten new petrified forest sites within GVP have been identified (Figure 1, Table 1). From these



**TABLE 1.** Summary of the range of dimensions of the samples and the total number of samples collected from the 10 new fossiliferous sites along with site details and observations on the autochthonous/allochthonous nature of the material collected.

Locality	Code	Longitude	Latitude	Altitude (m)	Total number collected	Length (cm)	Radial (cm)	Tangential (cm)	Observation
<b>Western GVP (Hancili Formation)</b>									
Bolu-Kıbrısçık-Kuzca-Alanhümmet Villages	KUZ	E31°45'45.7554"	N40°19'49.6914"	1272	22	2.5–25	1–9	2–11	Sample KUZ20 allochthonous; all others ex-situ autochthonous
Beypazarı-Nuhhoca Village (Kıraluç)	KIR	E31°45'53.4234"	N40°18'3.528"	1268	73	4–22	1.5–10	1–16	ex-situ autochthonous
Beypazarı-Aşağıgüney Village	AGU	E31°45'54.7554"	N40°19' 6.6714"	1141/1250	28	5–24	3–9.5	2–10	13 samples allochthonous; all others ex-situ autochthonous
Beypazarı-Karaşar Village	KAR	E31° 58' 53.2194"	N40° 19' 20.6034"	1370	1	11	6.5	6.5	ex-situ autochthonous
Beypazarı-Mençeler Valley	MEN	E31°51'1.9794"/E31°51'1.9794"	N40°17'33.1794"/N40°17'31.776"	1340/ 1374	34	5–20.5	6–12	1–11	ex-situ autochthonous
Beypazarı-Inozu Valley (north side)	INL	E31°54'55.26"	N40°15'24.8754"	1117	10	7–28	2–8.5	1.5–25	ex-situ autochthonous
Beypazarı-Inozu Valley	INO	E31°54'34.596"	N40°15'43.308"	998	11	3–24	2.5–8	1.5–14	ex-situ autochthonous
<b>Central GVP (Pazar Formation)</b>									
Çamlıdere-	BUG	E32°26'3.2994"	N40°24'34.484"	1005	30	4–21	2–12.5	1–9	ex-situ honous
Çamlıdere-Elmalı Village	ELM	E32°28'31.692"	N40°25'0.804"	1233	18	3–19.5	4–13.5	1.5–8.5	ex-situ autochthonous
Kızılcahamam-Soğuksu	SOG	E32°37'49.9404"	N40°28'7.2948"	1052	6	1–11	2–8.5	1–6	ex-situ autochthonous

sites a total of 233 silicified wood samples were collected ranging from 3 to 30 cm in length/height and 0.5 to 17 cm in diameter. The woods comprise both in situ stumps (Figure 3) decapitated by a volcanic eruption, but still in their life position (autochthonous), or relatively small pieces of in situ (allochthonous) material including those specimens from trees felled by volcanic activity and subsequently moved from the original growth position (Figure 4). Preservation is variable with some specimens showing good preservation, whilst others are poorly preserved. The in situ fossil wood specimens collected from the Çamlıdere-Elmalı Village site (ELM) were generally too badly preserved for any further analysis.

Each petrified wood sample is coded using the first three letters of the site name (Table 1) and

numbered sequentially. Those samples found *in situ* are identified by 'D' following the site name identifier (and prefixing the sample number). Samples are deposited in the Wood Anatomy Laboratory of the Forest Botany Department, Faculty of Forestry at Istanbul University.

Sixty-two of the samples were thin sectioned along the three planes of section (transverse section, TS, radial longitudinal section, RLS, and tangential longitudinal section, TLS), and identified using transmitted light microscopy. Descriptions follow the terminology of the IAWA Committee (1989) for hardwood identification and IAWA Committee (2004) for softwood identification wherever possible. Measurements were taken and documented using guidelines recommended by the IAWA Committee (1989, 2004) (i.e., the mean fol-



**FIGURE 3.** Examples of petrified wood specimens. **1**, Example of small wood samples found at the Bugarlar Fossil Site (BUG), from the central part of GVP, with an age of middle-late Burdigalian. **2**, Petrified woods found in their life position in Asagiguney Site (AGU). **3**, Material found in Kuzca Village (KUZ) in the west part of the GVP of early-middle Burdigalian age.

lowed by minimum and maximum measurements in parentheses). Identifications were made following reference to well-described extant wood, such as vouchered specimens housed in Forest Botany Department of Faculty of Forestry, University of Istanbul, comparisons with published literature (Jacquot, 1955; Greguss, 1955, 1967; Barefoot and Hankins, 1982; Fahn et al., 1986; Schweingruber, 1990, 1993; Tidwell, 1998; Eliçin, 1977; Vassio



**FIGURE 4.** Example of allochthonous petrified woods from Kuzca Village (KUZ) in the Galatian Volcanic Province.

et al., 2008; Wheeler, 2011; Akkemik and Yaman, 2012; Uhl et al., 2014) and online databases such as InsideWood for angiosperm wood (<http://inside-wood.lib.nsc.edu>).

#### SYSTEMATIC PALAEOBOTANY

A total of 14 different genera were identified from 10 fossil sites across the GVP. Seven of them were conifers, and the rest angiosperms. A summary of the wood characters of each taxon outlined below is based on all samples listed. Measurements for each individual sample are given in Appendices 1 and 2.

Family CUPRESSACEAE Rich. ex Bartling, 1830

Genus JUNIPERUS Linnaeus, 1753

*Juniperus* sp.

**Material.** AGU04, AGU05, AGU09, AGU11 (in situ); AGUD3, AGUD7, AGUD8, AGUD10 (in situ).

**Locality.** Near to the village of Beypazarı-Aşağıgüney (AGU).

**Description.** Ex-situ and in-situ material with dimensions ranging from 4 to 16 cm in height/length, and 2 to 8 cm in diameter.

Transverse section (TS). Growth ring widths 0.2–1.5 mm with distinct boundaries. Transition from earlywood to latewood gradual with latewood tracheids forming a layer 1–2 rows deep; false rings common. Resin canals absent. Intercellular spaces obvious and occur throughout the wood. Tracheid tangential and radial diameters measure 15 (5–25)  $\mu\text{m}$  and 16 (3–28)  $\mu\text{m}$ , respectively, in earlywood, and 9 (2–17)  $\mu\text{m}$  and 5 (1–13)  $\mu\text{m}$ , respectively, in the latewood. Tracheid wall thickness measures 4 (1–9)  $\mu\text{m}$  (thin walled) in earlywood and 6 (2–12)  $\mu\text{m}$  (thick walled) in latewood. Axial parenchyma is



common, and both diffuse (in the transition from earlywood to latewood and within the latewood itself) and tangentially zonate (Figure 5).

Tangential longitudinal section (TLS). Rays exclusively uniseriate. Ray height is 2–5 (max. 12) cells or 45 (9–226)  $\mu\text{m}$ ; ray width is 8 (4–18)  $\mu\text{m}$  in width. Transverse end walls of axial parenchyma cells are clearly, or slightly, nodular (Figure 5). Tracheid pitting in tangential walls is common.

Radial longitudinal section (RLS). Rays homogeneous. Ray tracheids absent. End walls of ray parenchyma cells nodular and indentures are present. Nodular end walls present. Cross-field pitting is cupressoid with 3–5 pits per cross-field. Tracheid pitting in radial walls of the earlywood is predominantly uniseriate (Figure 5).

**Comments.** The combination of wood features described here is consistent with that of *Juniperus* and are similar to those of the eight juniper species growing today in Turkey. All eight species grow across the whole of Anatolia with three of them, *Juniperus excelsa* M. Bieb., *J. drupacea* Labill. and *J. oxycedrus* L., being arboreal in habit (Akkemik and Yaman, 2012). Akkemik et al. (2016) also identified *Juniperus* wood from the Hoçaş Fossil Forest site and given the close anatomical similarity and geographical proximity, we conclude the material described here was probably derived for the same taxon as that found in the Hoçaş Fossil Forest site. Therefore, we assign the material described here to *Juniperus* species. Recently, Denk et al. (2017b) reported twigs of *Juniperus* from early Burdigalian deposits of Beskonak (Güvem), which supports evidence for juniper growing in this region during the early Miocene.

Genus SEQUOIA (D. Don) Endlicher, 1847  
*Sequoia* sp.

**Material.** INO04, BUG20 and SOG04.

**Locality.** Within the Beypazarı Inozu Valley on the south side (INO), and near the villages of Çamlıdere-Buğrular (BUG) and Kızılcahamam-Soğuksu (SOG).

**Description.** The following description was based on *ex-situ* material collected from three sites with dimensions ranging from 4 to 6.5 cm in diameter by 6 to 10.5 cm in height.

Transverse section (TS). Growth ring widths 0.5–3 mm with distinct boundaries. Transition from earlywood to latewood abrupt with a wide zone (3–22 cells) of latewood tracheids with reduced radial diameters. No resin canals observed. Tracheid tangential and radial diameters measure 26 (11–42)  $\mu\text{m}$  and 29 (14–44)  $\mu\text{m}$ , respectively, in the earlywood and 17 (5–38)  $\mu\text{m}$  and 10 (1–27)  $\mu\text{m}$  respec-

tively in the latewood. Tracheid wall thickness measures 3 (1–9)  $\mu\text{m}$  (thin walled) in earlywood and 9 (1–18)  $\mu\text{m}$  (thick walled) in latewood. Axial parenchyma common and diffuse to tangentially zonate in the transition zone and latewood. Vascular traces absent (Figure 6).

Tangential longitudinal section (TLS). Rays uniseriate and (low) medium to high (2–27 cells), rarely partly-biseriate. Ray height is 174 (28–437)  $\mu\text{m}$ ; ray width is 13 (6–27)  $\mu\text{m}$ . Helical thickenings absent. Axial parenchyma present with either slightly nodular or smooth transverse end walls; brown contents (possible preservational feature) present in each parenchyma cell (Figure 6).

Radial longitudinal section (RLS). Tracheid pitting generally uniseriate in the latewood, and uni- to biseriate and opposite in the earlywood. Crassulae commonly associated with both uniseriate and biseriate pitting. Two to six taxodioid pits per cross-field and horizontally arranged in one or two rows. No ray tracheids observed. Ray parenchyma walls smooth. Axial parenchyma cells with (brown) contents (Figure 6).

**Comments.** The combination of wood characters described here is consistent with that of the extant *Sequoia/Sequoiadendron/Metasequoia* complex and very similar to those fossil woods also identified as *Sequoia* from the Pelitçik Fossil Forest site located between Beypazarı Inozu Valley (INO) and Kızılcahamam-Soğuksu (SOG) (Akkemik et al., 2009). Further evidence for the presence of *Sequoia/Sequoiadendron/Metasequoia* comes from leaf imprints and pollen identified from the Güvem area (Kasaplıgil, 1977; Denk et al., 2017b). *Sequoia* and the *Sequoia/Sequoiadendron/Metasequoia* complex are the most common tree taxa found growing in Anatolia during the Oligocene (Akkemik et al., 2005; Akkemik and Sakıncı, 2013) and into the early to mid-Miocene (Kasaplıgil, 1977; Akkemik et al., 2009; Denk et al., 2017b).

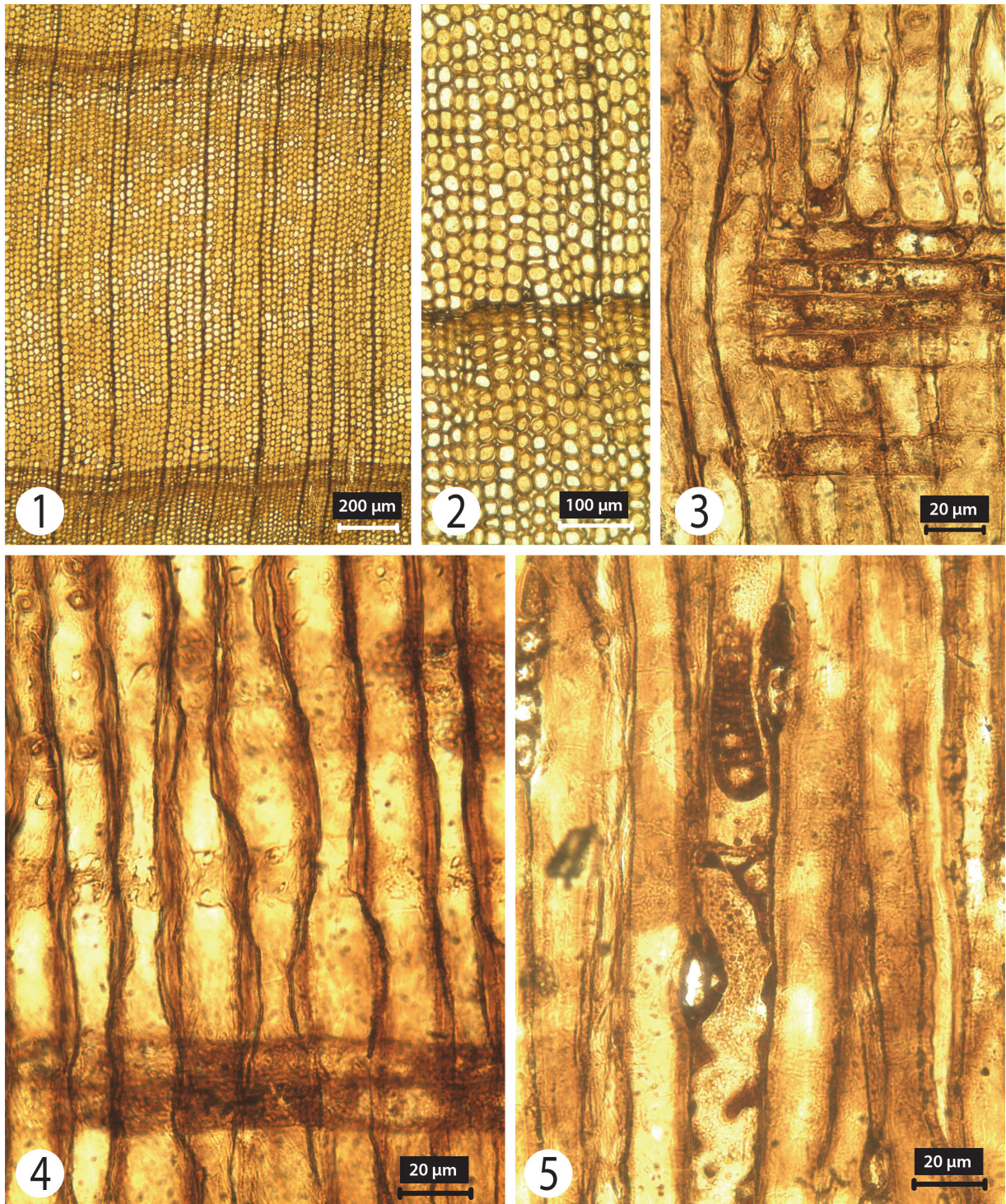
Genus TAXODIUM Richard, 1810  
*Taxodium* sp.

**Material.** BUG01.

**Locality.** Near to the village of Çamlıdere-Buğrular

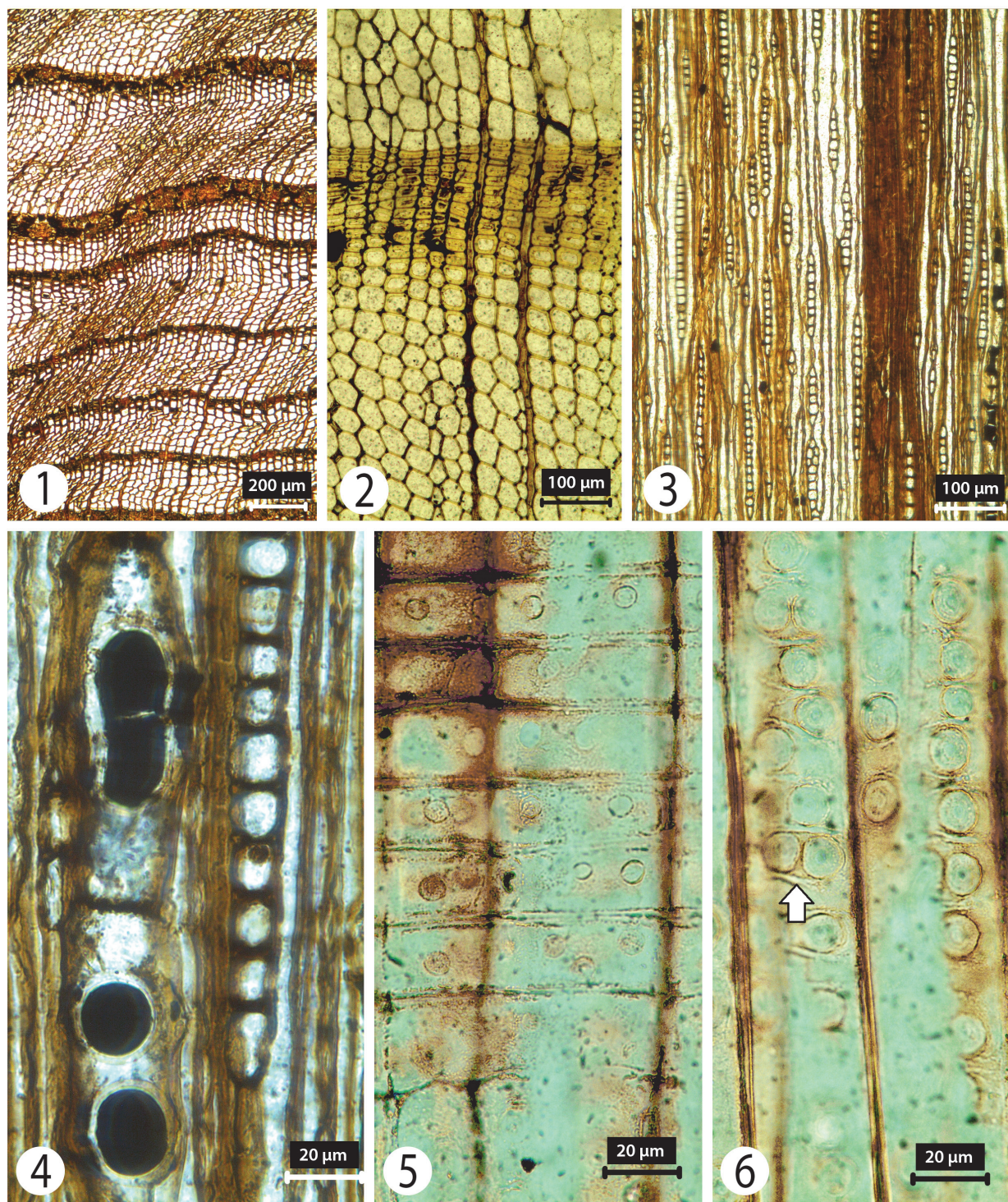
**Description.** The following description is based on one poorly preserved *ex-situ* sample measuring 12 cm in diameter by 14 cm in length.

Transverse section (TS). Growth ring boundaries indistinct with ring widths ranging from 0.7 to 2 mm. Transition from earlywood to latewood gradual with a narrow zone (1–3, rarely more) of thin walled, latewood tracheids. Tracheid tangential and radial diameters measure 39 (26–53)  $\mu\text{m}$  and 50 (33–63)  $\mu\text{m}$ , respectively, in the earlywood and 31 (19–50)



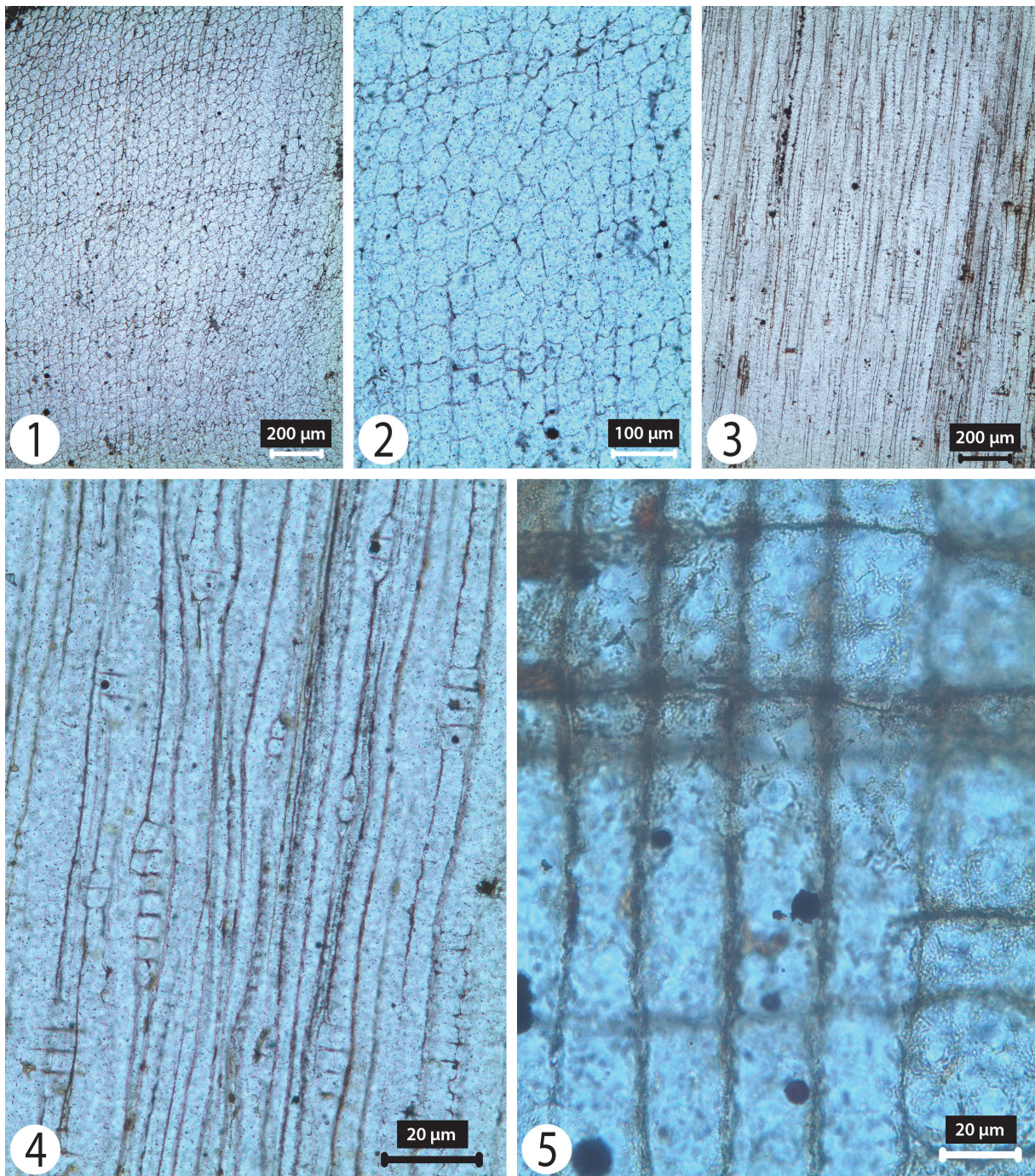
**FIGURE 5.** *Juniperus* sp. 1, TS showing growth rings with false rings in summer wood. 2, Close-up of tracheidal arrangement within one growth ring (not inter-tracheidal spaces) and uniseriate rays (TS). 3, RLS showing procumbent cells comprising the ray. 4, RLS showing cupressoid type cross-field pits (RLS). 5, TLS showing axial parenchyma cells with nodular end walls.





**FIGURE 6.** *Sequoia* sp. 1, TS showing distorted (preservational feature) growth rings with clear latewood. 2, TS of one growth ring showing a late wood zone of up to 11 rows of tracheids. 3, TLS showing uniseriate, homogenous rays. 4, TLS showing axial parenchyma cells with smooth end walls and dark contents (preservational feature). 5, Rays with taxodioid type cross-field pits (RLS). 6, Bordered pits in one and two rows with crassulae visible (arrow) (TLS).





**FIGURE 7.** *Taxodium*. **1**, TS showing very narrow, indistinct growing rings. **2**, TS showing latewood zones of only 2 rows of tracheids. **3-4**, TLS showing uniseriate rays. **5**, Rays with taxodioid type cross-field pits (RLS).

µm and 27 (20–37) µm, respectively, in the latewood. Tracheid wall thickness measures 3 (1–5) µm (thin walled) in earlywood and 3 (1–4) µm (thin walled) in latewood. Axial parenchyma diffuses and rare. Resin canals absent (Figure 7).

Tangential longitudinal section (TLS). Rays uniseriate, rarely partly biseriate and homogenous. Rays measure 422 (303–569) µm in height and 18 (6–27) µm in width. Tracheidal pitting rare, uniseriate and non-contiguous. Helical thickening absent (Figure 7).



Radial longitudinal section (RLS). Tracheid pitting uniseriate or biseriate (rarely triseriate) and opposite. Rays low. Three to seven seemingly taxodioid pits per cross field both horizontally and randomly arranged. Ray tracheids absent. Ray parenchyma walls smooth (Figure 7).

**Comments.** The combination of wood characters described here is consistent with that of *Taxodium*. Akkemik et al. (2009) identified this genus from Pelitçik, a site proximal to Çamlıdere-Buğralar Village (Figure 1). *Taxodium* is one of the more frequently encountered tree genera from the early Miocene with micro- and macrofossils having been identified from Anatolia (Kasaplıgil, 1977; Akkemik et al., 2009), Italy (Biondi and Brugiapaglia, 1991) and Greece (Velitzelos et al., 2014).

Family PINACEAE Sprengel ex F. Rudolphi, 1830  
Genus CEDRUS Trew, 1757  
*Cedrus* sp.

**Material.** INL03, INO05 and INO06.

**Localities.** Within the Beypazarı-Inozu Valley on the south side (INO) and north side (INL).

**Description.** All samples are *ex-situ* with dimensions ranging from 7 to 10 cm in length and 2 to 4.5 cm in diameter.

Transverse section (TS). Growth ring width ranging from 0.4 to 1.5 mm. Growth ring boundaries distinct with transition from earlywood to latewood gradual. Latewood zone is generally wide (5–30 cells in radial direction). Tangential and radial tracheid diameters measure 21 (6–38)  $\mu\text{m}$  and 30 (15–52)  $\mu\text{m}$ , respectively, in the earlywood and 13 (4–26)  $\mu\text{m}$  and 7 (1–15)  $\mu\text{m}$ , respectively, in latewood. Tracheid wall thickness measures 5 (1–9)  $\mu\text{m}$  (thin walled) in earlywood and 12 (5–21)  $\mu\text{m}$  (thick walled) in latewood. Axial parenchyma marginal, diffuse or in 2–3 rows. Resin canals absent although traumatic resin canals in both the latewood and transition zones of some rings are common (Figure 8).

Tangential longitudinal section (TLS). Rays entirely uniseriate, 13–22 cells high or 215 (27–1004)  $\mu\text{m}$  in height and 12 (4–31)  $\mu\text{m}$  in width. Many traumatic vertical resin canals observed, rays with horizontal resin canals absent (Figure 8).

Radial longitudinal section (RLS). Rays are heterogeneous. Smooth walled ray tracheids are common and form 1–2 rows with enlarged cells when situated towards the outside of the rays. End- and horizontal walls of ray parenchyma cells are distinctly pitted. Tracheid pitting in radial walls (in early wood) is predominantly uniseriate with scalloped tori. Cross-field pitting is cupressoid with 1–4 pits per cross-field (Figure 8).

**Comments.** Woods with resin canals also occur in wood of *Pinus* and *Picea* but differ in that they are arranged in one row within each growth ring. The combination of wood characters described here is consistent with that of *Cedrus*. *Cedrus* fossils have been described from the central part of the GVP. Denk et al. (2017b) reported a seed of *Cedrus* at Güvem and pollen, which is considered to be relatively common in the Güvem area, by Yavuz-Isik (2008). At the Hoçaş Fossil Forest site in the west Akkemik et al. (2016) identified the wood of *Cedrus*. The published wood characteristics of extant *Cedrus libani* A. Rich. (Akkemik and Yaman, 2012) are virtually identical to those of the petrified woods but although identical at species level, when the age of the material is considered we cannot ignore the fact that this material might be a representative of an ancestral lineage (Akkemik et al., 2016) and for this reason these specimens are referred to *Cedrus*. as *Cedrus* cf. *libani*. Today one species, *Cedrus libani* (Taurus cedar), grows in the Mediterranean Region of Turkey. However, two more species (*C. brevifolia* A. Henry ex Elwes et A. Henry and *C. atlantica* [Endl.] Manetti ex Carrière) are native to land surrounding the Mediterranean Sea.

Genus PICEA A. Dietrich, 1824  
*Picea* sp.

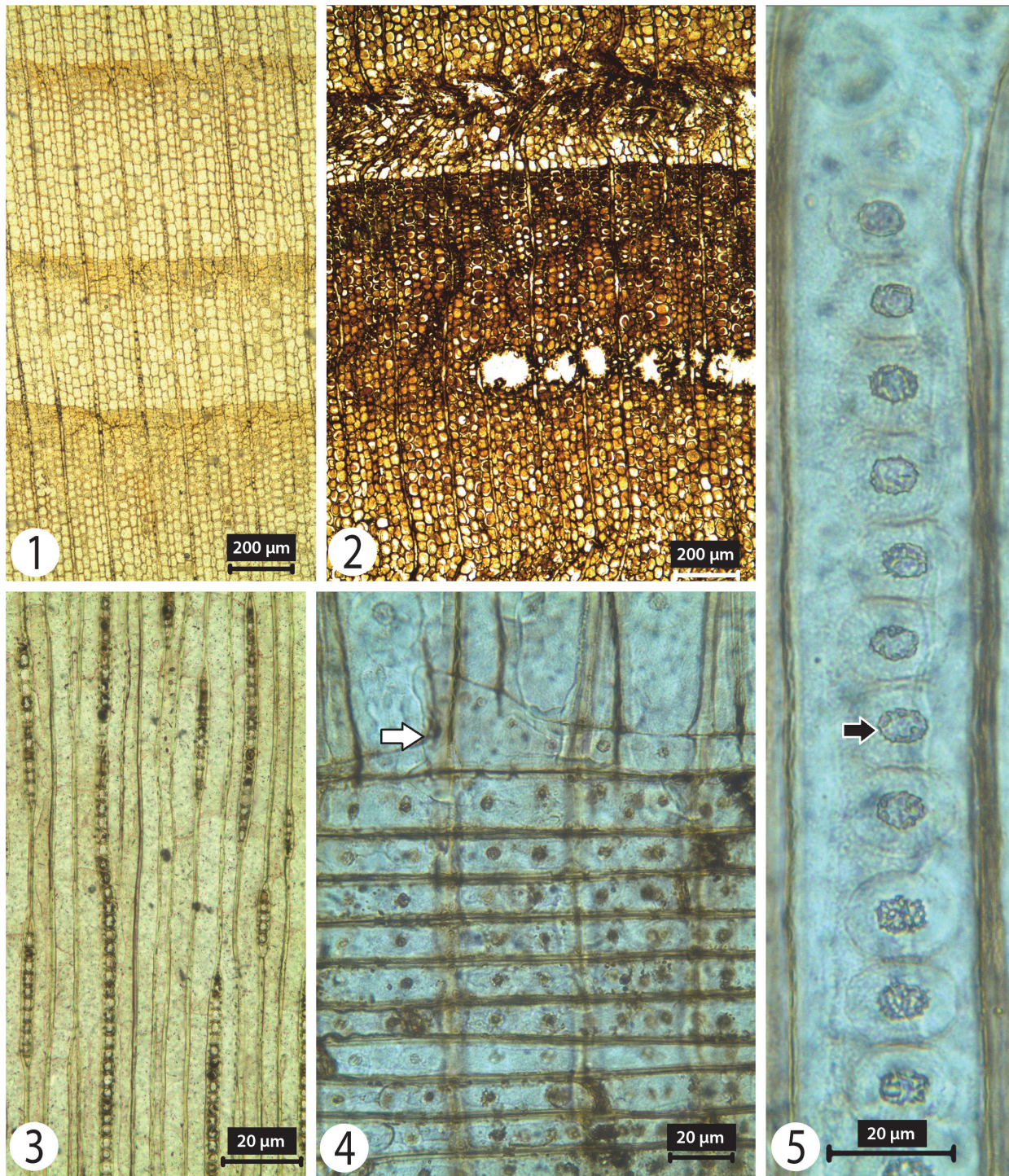
**Material.** AGUD2 and KUZ20.

**Localities.** Near the villages of Beypazarı-Aşağıgüney and Bolu-Kıbrısçık-Kuzca- Alanhümmet.

**Description.** The following description is based on two specimens, one in situ AGUD2 in life position whereas the second, KUZ20, is in situ. Dimensions of the samples range from 11–17 cm in height/length and 5.5–7.5 cm in diameter.

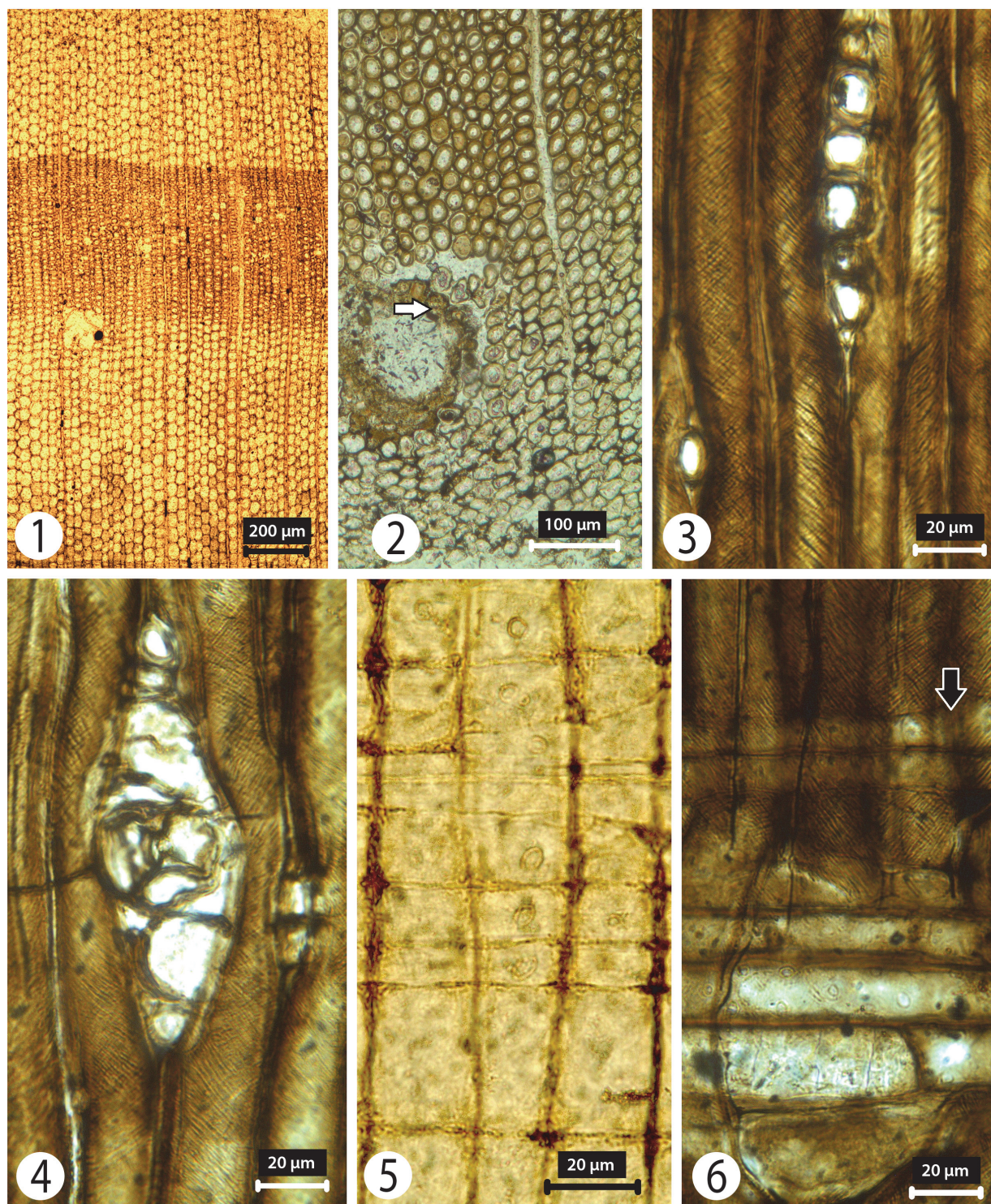
Transverse section (TS). Growth ring boundaries distinct with transition from earlywood to latewood gradual. Growth ring width is very wide, 3.5–9 mm. Tracheid tangential and radial diameters measure 17 (8–26)  $\mu\text{m}$  and 24 (9–54)  $\mu\text{m}$ , respectively, in earlywood, and 8 (4–16)  $\mu\text{m}$  and 7 (1–18)  $\mu\text{m}$ , respectively, in latewood. Tracheid wall thickness measures 3 (1–6)  $\mu\text{m}$  (thin walled) in earlywood and 5 (2–9)  $\mu\text{m}$  (thick walled) in latewood. Resin canals common and surrounded by epithelial cells of varying dimensions (Figure 9), sometimes paired and generally restricted to transition zone; tangential and radial diameters measure 129 (75–192)  $\mu\text{m}$  and 141 (112–193)  $\mu\text{m}$ , respectively.

Tangential longitudinal section (TLS). Rays generally uniseriate unless with horizontal resin canals. Rays with resin canal uniseriate to partly biseriate. Rays measure 298 (128–405)  $\mu\text{m}$  in height when



**FIGURE 8.** *Cedrus*. 1, TS showing clear growth rings and no (traumatic) resin canals. 2, TS showing growth rings and associated traumatic resin canals. 3, TLS showing uniseriate rays. 4, RLS showing rays with ray tracheids (white arrow) and cupressoid type cross-field pits. 5, Bordered pits with scalloped tori (RLS) (black arrow).





**FIGURE 9.** *Picea*. **1**, TS showing wide growth rings and resin canal. **2**, TS resin canal with no visible epithelial cells (white arrow). **3**, TLS showing a ray without any horizontal resin canals and tracheids with helical thickenings. **4**, TLS showing a ray with a horizontal resin canal and helical thickenings in the tracheids. **5**, Cupressoid type cross-field pits. **6**, Ray tracheid (black arrow) with smooth wall and cupressoid type pits in the cross-fields (RLS).



resin canals are present, and 144 (28–353)  $\mu\text{m}$  in those rays without resin canal; ray widths measure 44 (38–48)  $\mu\text{m}$  in those rays with resin canals, and 12 (6–20)  $\mu\text{m}$  in rays without resin canals. When present in rays, resin canals are surrounded by a layer of epithelial cells. Helical thickenings are very common in AGUD2 but absent in KUZ20 (Figure 9).

Radial longitudinal section (RLS). Rays heterogeneous. Ray tracheids common in 1–2 rows with smooth, thin walls, rarely slightly wavy. Both horizontal and end walls of ray parenchyma cells are distinctly pitted. Cross-field pitting is cupressoid with 1–3 pits per cross-field. Tracheid pitting in radial walls of the early wood is predominantly uniseriate (Figure 9).

**Comments.** The combination of wood characters described here is consistent with that of *Picea*. Today *Picea* is represented by three species around the Mediterranean and Europe (i.e., *Picea abies* (L.) Karsten, *P. omorika* (Panëiæ) Purkyne, and *P. orientalis* (L.) Link.), but identification to species level is not always possible using wood characters alone. Within these three species, Schweingruber (1990) notes that cross-field pitting is generally cupressoid in *Picea abies* (and *Picea omorika*, whereas Akkemik and Yaman (2012) have observed piceoid cross-field pitting in *Picea orientalis*. Our samples exhibit predominantly cupressoid cross-field pitting suggesting that the material is probably more similar to European *Picea* species, but the fossil wood differs from all extant European and Mediterranean species by the presence of helical thickening. Therefore, the fossil material described here is assigned to *Picea* sp.

Genus PINUS Linnaeus, 1753

*Pinus* sp.

**Material.** AGU01, AGU22, AGU23, INO12, KIR21, KUZ01, KUZ07, KUZ09, KUZ17, MEN32 and MEN33 (in situ); AGUD12 and KUZ01 (in situ).

**Localities.** Near the villages of Beypazarı-Aşağıgüney (AGU), Beypazarı-Nuhhoca (Kıraluç Site) (KIR), Bolu-Kıbrısçık-Kuzca- Alanhümmet (KUZ) and Beypazarı-Mençeler (MEN), and in the Beypazarı Inozu Valley (south side) (INO).

**Description.** In situ and in situ material with dimensions ranging from 9 to 16 cm in height/length by 2 to 8.5 cm in diameter.

Transverse section (TS). Growth ring boundaries distinct with transition from earlywood to latewood abrupt in AGU22 and KIR21, and gradual in the others. Width of growth rings range from 1–3.5 to 7 mm with a mode of 4–5 mm; i.e., generally wide

except in specimen AGU22. Tracheid tangential and radial diameters measure 25 (6–59)  $\mu\text{m}$  and 29 (9–55)  $\mu\text{m}$ , respectively, in the earlywood and 10 (1–38)  $\mu\text{m}$  and 7 (1–22)  $\mu\text{m}$ , respectively, in the latewood; tracheid wall thickness measure 4 (1–15)  $\mu\text{m}$  (thin walled) in earlywood and 6 (1–19)  $\mu\text{m}$  (thick walled), respectively, in latewood. Resin canals present, frequent in the earlywood-latewood transition zone; thin-walled epithelial cells observed surrounding some resin canals (Figure 10). Tangential and radial diameters of resin canals are 138 (41–354)  $\mu\text{m}$  and 139 (61–295)  $\mu\text{m}$ , respectively.

Tangential longitudinal section (TLS). Rays with and without horizontal resin canals. Rays without horizontal canals are uniseriate, 144 (29–425)  $\mu\text{m}$  in height and 14 (4–30)  $\mu\text{m}$  in width. Rays with resin canals present measure 301 (83–501)  $\mu\text{m}$  in height and 41 (13–67)  $\mu\text{m}$  in width (Figure 10).

Radial longitudinal section (RLS). Rays heterogeneous. Ray tracheids commonly in 1–4 rows with smooth walls, rarely slightly wavy. Both transverse walls and end walls of ray parenchyma cells are unpitted. Cross-field pitting is clearly pinoid with 1–2 pits per cross-field. Tracheid pitting uniseriate (Figure 10).

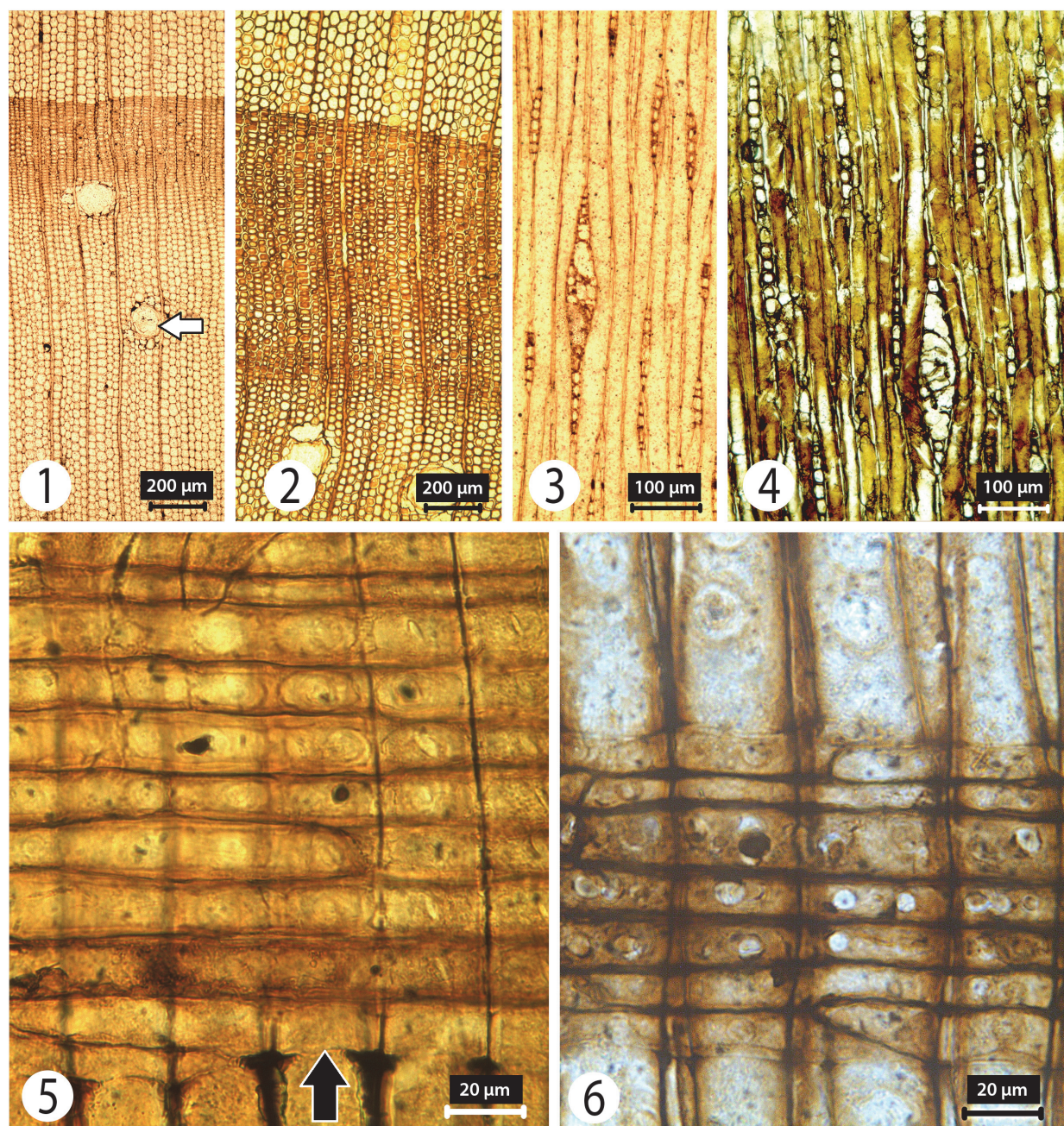
**Comments.** The combination of wood characters such as resin canals, heterogeneous rays and pinoid type cross-field pitting place this morpho-taxon within the genus *Pinus*. Today Turkey is represented by five species of *Pinus*, namely *P. nigra* Arnold, *P. sylvestris* L., *P. brutia* Ten., *P. pinea* L. and *P. halepensis* Miller. The fossils show greatest anatomical similarity to *Pinus pinea* with its smooth ray tracheids and 1–2 pinoid type of pits per cross-field (Akkemik and Yaman, 2012). *Pinus* wood has also been identified from the Kozyaka Fossil Forest site (Figure 1) in the west part of GVP (Akkemik et al., 2016), and other macrofossil (leaf and cone) evidence (Kasaplıgil, 1977; Denk et al., 2017b) found at both the Güvem and Hoçaş sites (Figure 1) that revealed *Pinus* comprised both the three needled and five needled varieties. With this in mind, it cannot be assumed that all *Pinus* fossil evidence originated from the same taxon. Moreover, since identification to species based on wood anatomy alone is often inconclusive, we refer the specimens described herein to *Pinus*.

Family PODOCARPACEAE Endlicher, 1847

Genus PODOCARPUS L'Heritier ex Persoon,  
1807

*Podocarpus* sp.

**Material.** KIR40, KIR70, KIR73 and MEN19.



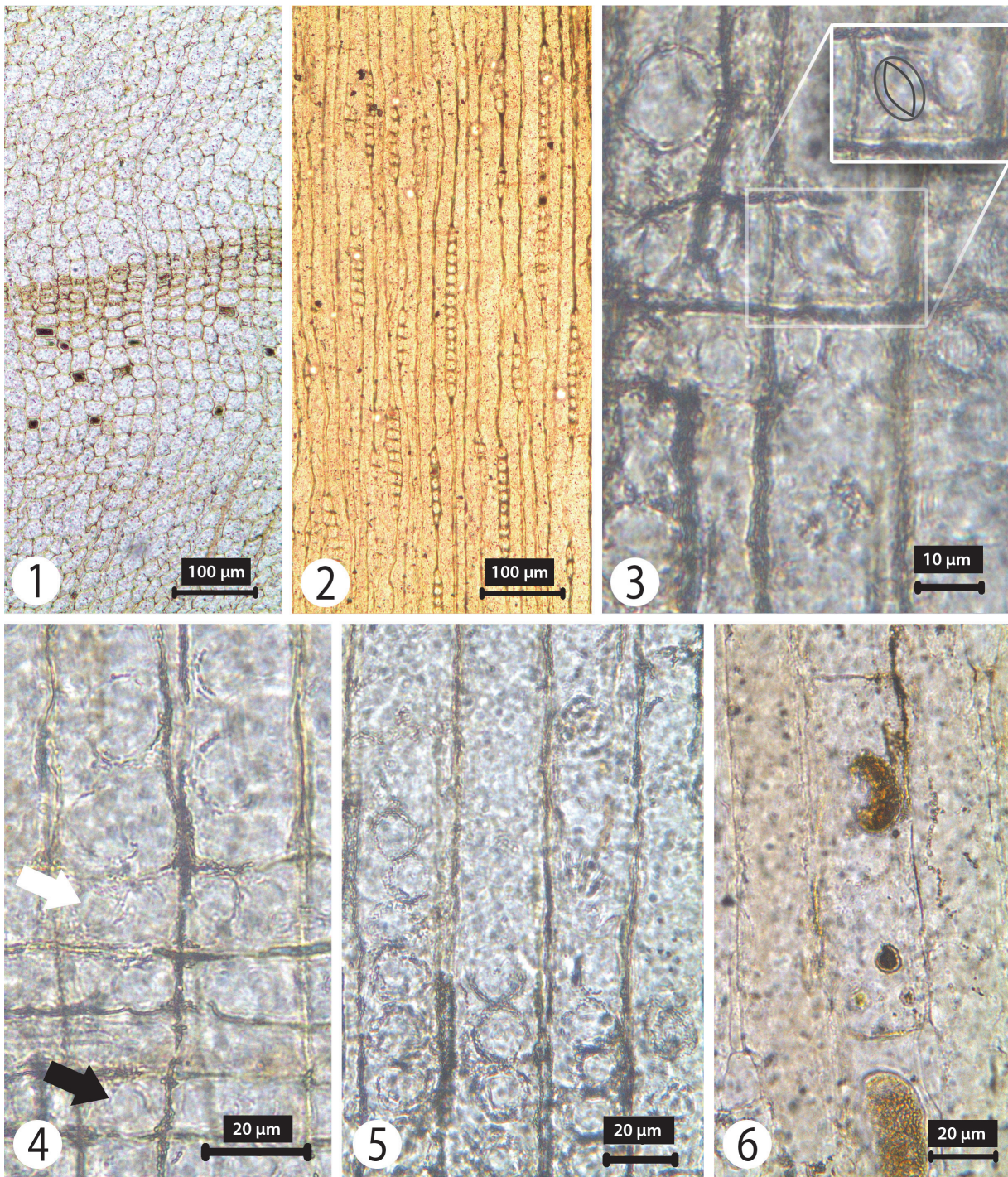
**FIGURE 10.** *Pinus*. **1**, TS showing wide growth rings and resin canal. **2**, TS resin canal with no visible epithelial cells (white arrow). **3-4**, Rays with and without horizontal resin canal and tracheids (TLS). **5**, RLS showing procumbent cells forming the ray. **6**, Pinoid type cross-field pits, and ray tracheids (black arrow) with smooth walls.

**Localities.** Near the village of Beypazarı-Nuhhoca (Kıralıç Site) (KIR) and in the Beypazarı-Mençeler Valley (MEN).

**Description.** All samples are ex-situ and dimensions range from 6–7cm in height and 1–4 cm in diameter. One sample, KIR40, is very poorly preserved, and some features were not observed.

Transverse section (TS). Growth rings boundaries distinct with ring widths ranging from 0.4 to 1.2 mm; transition from earlywood to latewood almost abrupt, with 4–7 rows of latewood tracheids in KIR73 but gradual in KIR70. Resin canals absent (Figure 11). Tracheid tangential and radial diameters measure 22 (5–40)  $\mu\text{m}$  and 23 (10–45)  $\mu\text{m}$ , respectively, in the earlywood, and 14 (7–22)  $\mu\text{m}$





**FIGURE 11.** *Podocarpus*. **1**, TS across one growth ring with diffuse axial parenchyma cells. **2**, Homogenous uniseriate rays (TLS). **3**, RLS showing large cupressoid and taxodioid type cross-field pits with two cross-field pits shown in the upper box. Left indicates the borders of pit drawn, right shows the oblique appearance of the pit. **4**, RLS showing procumbent nature of the cells forming the ray, and cross-field pits; white arrow indicates the oblique position of the pit; black arrow shows vertical pit. **5**, Radial pitting with one row in large (earlywood) tracheids. **6**, Axial parenchyma cells with thin end walls (vertical arrow).

and 7 (1–15)  $\mu\text{m}$ , respectively, in the latewood. Tracheid wall thickness is 3 (1–5)  $\mu\text{m}$  (thin walled) in the earlywood and 5 (1–12)  $\mu\text{m}$  (thick walled) in the latewood. Axial parenchyma common, diffuse and in short tangential bands both through the transition zone and in late latewood (Figure 11).

Tangential longitudinal section (TLS). Rays uniseriate, homogenous, very rarely biseriate, mostly short, 5–15 (2–21) cells in height or 95 (22–347)  $\mu\text{m}$  high and 10 (4–18)  $\mu\text{m}$  in width (Figure 11).

Radial longitudinal section (RLS). Tracheid pitting on radial walls, uniseriate with abietinean radial pitting (i.e., spaced and uniseriate; sensu Philippe and Bamford, 2008)—round, not contiguous, relatively large with diameters of 16–19  $\mu\text{m}$ . Poor preservation has mostly rendered only the outer borders of tracheid pits recognizable. Generally, two (1–3) pits per cross-field, large cupressoid and taxodioid. Vertical and oblique apertures observed in some pits (Figure 11). Horizontal cross-field pits were not observed.

**Comments.** The combination of wood characters including the presence of distinct growth rings, uniseriate bordered pits on radial walls and 1–2 (up to 3) large cupressoid and taxodioid pits per cross field place this material within the genus *Podocarpus* (Brezinova et al., 1994). Since identification to species based on wood anatomy alone is often inconclusive, we refer these specimens to *Podocarpus*. Fossil woods of putative affinity to the Podocarpaceae have been described by many authors and assigned to *Podocarpoxylon* Gothan, *Protopodocarpoxylon* Eckhold, *Phyllocladoxylon* Gothan, *Mesembrioxylon* Seward, *Metapodocarpoxylon* Duperon-Laudoueneix and Pons, *Protophyllocladoxylon* Kräusel, *Circoporoxylon* Kräusel and even *Cupressinoxylon* Göppert (Pujana and Ruiz, 2017). Pujana and Ruiz (2017) have recently reviewed fossil *Podocarpoxylon* woods and conclude that although many supposed woods of the podocarpaceous type have been found in the Northern Hemisphere, the deficit of other podocarpaceous fossil remains substantiates the fact that this the fossil genus may include wood from taxa other than those of Podocarpaceae. Possible podocarpaceous woods have previously been identified: Süss and Velitzelos (2000) identified two *Podocarpoxylon* wood species from the lower Miocene of Lesbos, and from Thrace (Sakıncı et al., 2007) and Akkemik (work in progress) have both identified wood samples as *Podocarpoxylon*. This suggests a possible podocarpaceous tree grew across this area during the early Miocene. The

sample described here is the first record of *Podocarpus* in the wood flora of GVP.

Family FAGACEAE Dumortier, 1829

Genus QUERCUS Linnaeus, 1753

*Quercus* section *Ilex* sp.

**Material.** INL09, KIR07, KAR01, MEN02, MEN22 and MEN28.

**Localities.** Within the Beypazarı-Inozu valley (north side) (INL) and the Beypazarı-Mençeler valley (MEN) and near to the villages of Beypazarı-Nuhhoca (Kıraluç Site) (KIR) and Beypazarı-Karaşar (KAR).

**Description.** All samples are in situ and range from 4 to 12 cm in length and 4 to 8.5 cm in diameter.

Transverse section (TS). Growth ring boundaries indistinct, wood diffuse porous. Vessel tangential diameter 82 (22–182)  $\mu\text{m}$ , and radial diameter 109 (33–206)  $\mu\text{m}$ ; average vessel density is 22 (13–29) vessels per  $\text{mm}^2$ . Tyloses common in wider vessels. Large multiseriate rays can clearly be seen with the naked eye. In well-preserved woods axial parenchyma is clearly visible arranged in short tangential bands in single rows (Figure 12).

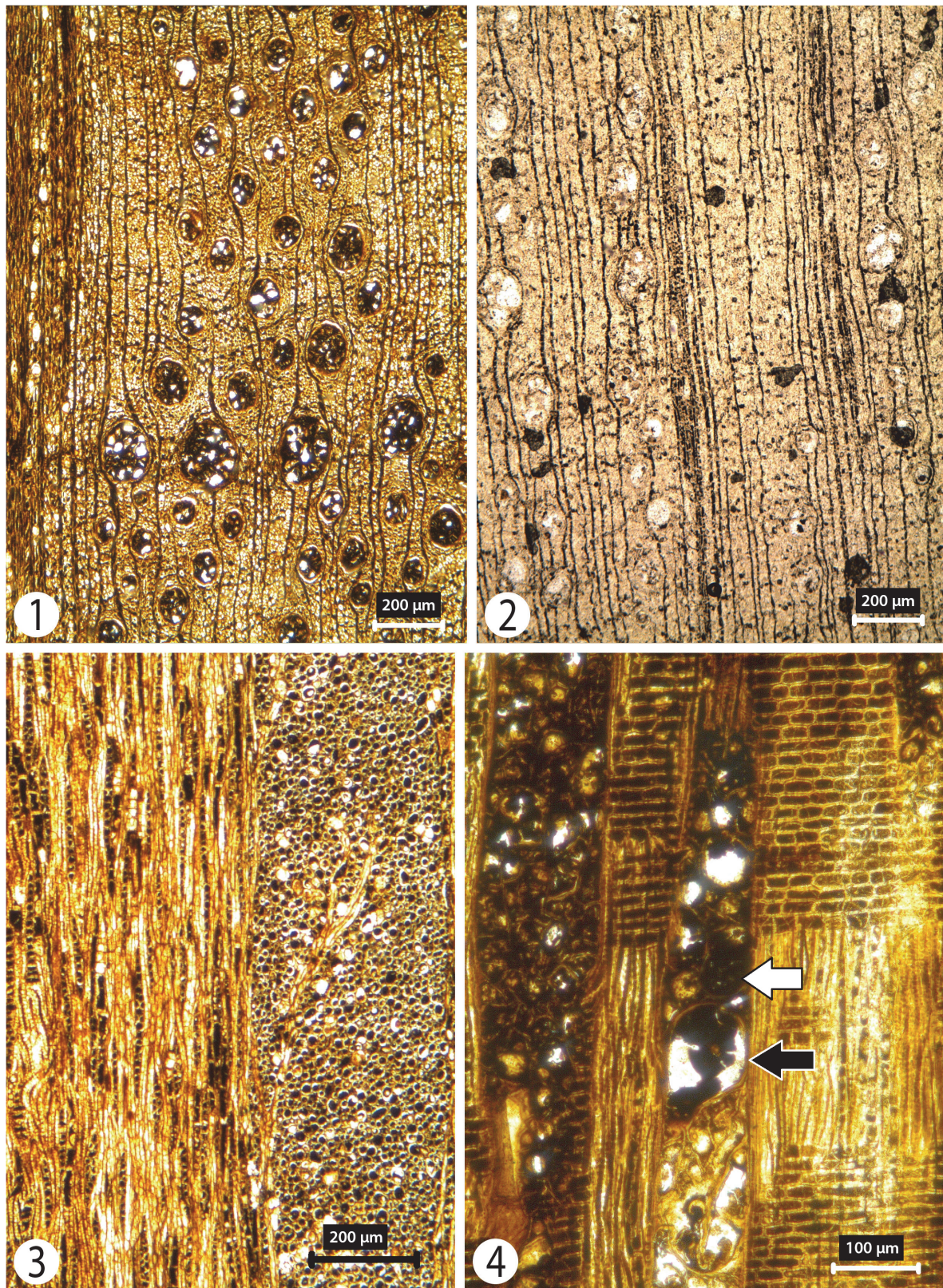
Tangential longitudinal section (TLS). Rays number 4–6 per mm and are of two distinct sizes. Narrow uniseriate and biseriate rays measure 334 (110–805)  $\mu\text{m}$  by 14 (4–26)  $\mu\text{m}$  in width; large multiseriate rays range from 10 to 20 cells, or 258 (67–411)  $\mu\text{m}$ , in width by >1 mm in height (Figure 12).

Radial longitudinal section (RLS). Perforation plates are simple. Rays are homocellular and composed of procumbent ray cells (Figure 12).

**Comments.** The combination of characters outlined above indicate that this wood type belongs to *Quercus* L. and in particular the *Quercus* section *Ilex* wood type given that all these fossil specimens are diffuse porous (as opposed to ring porous, characteristic of *Quercus* Sections *Quercus* and *Cerris*). Furthermore, macrofossils of *Quercus* section *Ilex* (Denk et al., 2017a, 2017c) and pollen (Yavuz-Işık, 2008) have also been found from the Güvern region.

This wood type, which may represent a lineage ancestral to the modern western Eurasia sclerophyllous oaks, has also been found at the Hoçaş site (Figure 1) (Akkemik et al., 2016). Akkemik et al. (2016) concluded that the presence of evergreen oak species found at a total of five fossil sites in GVP suggested that this taxon was not an uncommon element of the early Miocene vegetation. Further studies on modern oaks and sediments around the Mediterranean Basin (Denk and Grimm, 2009, 2010; Hubert et al., 2014; Denk et





**FIGURE 12.** *Quercus* (section *Illex*). 1, TS showing solitary vessels with tyloses and one multiseriate ray. 2, TS showing radial alignment of the solitary vessels and uniseriate rays. 3, TLS showing rays of two sizes. 4, RLS of rays with procumbent cells, vessel with a simple perforation plate (black arrow) and filled with tyloses (white arrow).



al., 2017c) suggest a Eurasian group of evergreen oaks, comprising about 30 species, that occurred across the Mediterranean region and extended into the southern foothills of the Himalayas and across East and Southeast Asia during the early Miocene. The material found on the GVP may belong to this Eurasian group.

Family OLEACEAE Hoffmannsegg and Link, 1809  
Genus FRAXINUS Linnaeus, 1753  
*Fraxinus* sp.

**Material.** AGU13.

**Locality.** Near the village of Beypazarı-Aşağıgüney (AGU).

**Description.** The following description is based on one piece of in situ material measuring 16 cm in length and 8 cm in diameter.

Transversal section (TS). Growth ring boundaries distinct, with ring widths ranging from 2.3 to 5 mm. Wood ring-porous. Earlywood vessels wide becoming narrower in the latewood. Tangential and radial diameters of vessel lumen 69 (26–116)  $\mu\text{m}$  and 90 (33–130)  $\mu\text{m}$ , respectively, in the earlywood, and in the latewood tangential and radial diameters measure 27 (13–46)  $\mu\text{m}$  and 24 (11–52)  $\mu\text{m}$ , respectively. Vessel density 83 (67–104) vessels per  $\text{mm}^2$  with 26 (23–32) vessels per  $\text{mm}^2$  in earlywood and 57 (44–72) vessels in latewood. Vessels solitary and in radial multiples of up to three vessels. Paratracheal axial parenchyma almost lozenge-aliform especially in the latewood (Figure 13).

Tangential longitudinal section (TLS). Rays uniseriate to biseriate sometimes triseriate. Rays number  $\geq 12$  per mm (Figure 13) and measure 216 (58–345)  $\mu\text{m}$  in height and 23 (11–38)  $\mu\text{m}$  in width. Perforation plates simple. Fibres thin- to thick-walled (Figure 13).

Radial longitudinal section (RLS). Homocellular in some rays with all ray cells entirely procumbent others are heterocellular with procumbent body ray cells and square/upright marginals (Figure 13).

**Comments.** The combination of wood characters including ring porous vessels, distinct growth ring borders, clear transition from earlywood to latewood, sparse vessels in the latewood, rays 1–3 seriate with mostly procumbent body cells and slightly lozenge-aliform axial parenchyma shows closest anatomical similarity to *Fraxinus* L. (Akkemik and Yaman, 2012) and has thus been assigned to *Fraxinus* sp. Although *Fraxinus* is represented by four species, *F. excelsior* L., *F. angustifolia* Vahl., *F. ornus* L. and *F. pallisae* Wilmott. In Turkey today, this study represents the first record of *Fraxinus* wood material from the GVP (and from

Turkey as a whole). Other macrofossil and palynological evidence including leaves (Kasaplıgil, 1977), fruits (Denk et al., 2017b) and pollen (Karayığit et al., 1999) supports the presence of *Fraxinus* growing in the Güvem Basin at this time.

Family ROSACEAE Jussieu, 1789  
Genus PRUNUS Linnaeus, 1753  
*Prunus* sp.

**Material.** MEN31.

**Locality.** Beypazarı-Mençeler Valley (MEN).

**Description.** The following description is based on one piece of *ex-situ* material measuring 10 cm in length and 7 cm in width.

Transverse section (TS). Wood ring porous, sometimes more semi-ring porous, with distinct growth ring boundaries and ring widths ranging from 0.4 to 1 mm in width. Tangential and radial diameters of vessels measure 55 (32–75)  $\mu\text{m}$  and 82 (33–136)  $\mu\text{m}$ , respectively, in the earlywood, and 22 (8–43)  $\mu\text{m}$  and 30 (15–50)  $\mu\text{m}$ , respectively, in latewood. Vessels solitary or in short radial files of 2–4. Vessel density 67 (55–78) per  $\text{mm}^2$  (Figure 14).

Tangential longitudinal section (TLS). Rays number 4–12 per mm and are clearly of two distinct sizes (Figure 14); wide rays 4–10 seriate, or 66 (51–78)  $\mu\text{m}$  in width, and  $>1$  mm in height; narrow rays uniseriate (very rarely biseriate), or 7 (4–11)  $\mu\text{m}$  in width, and 117 (67–156)  $\mu\text{m}$  in height.

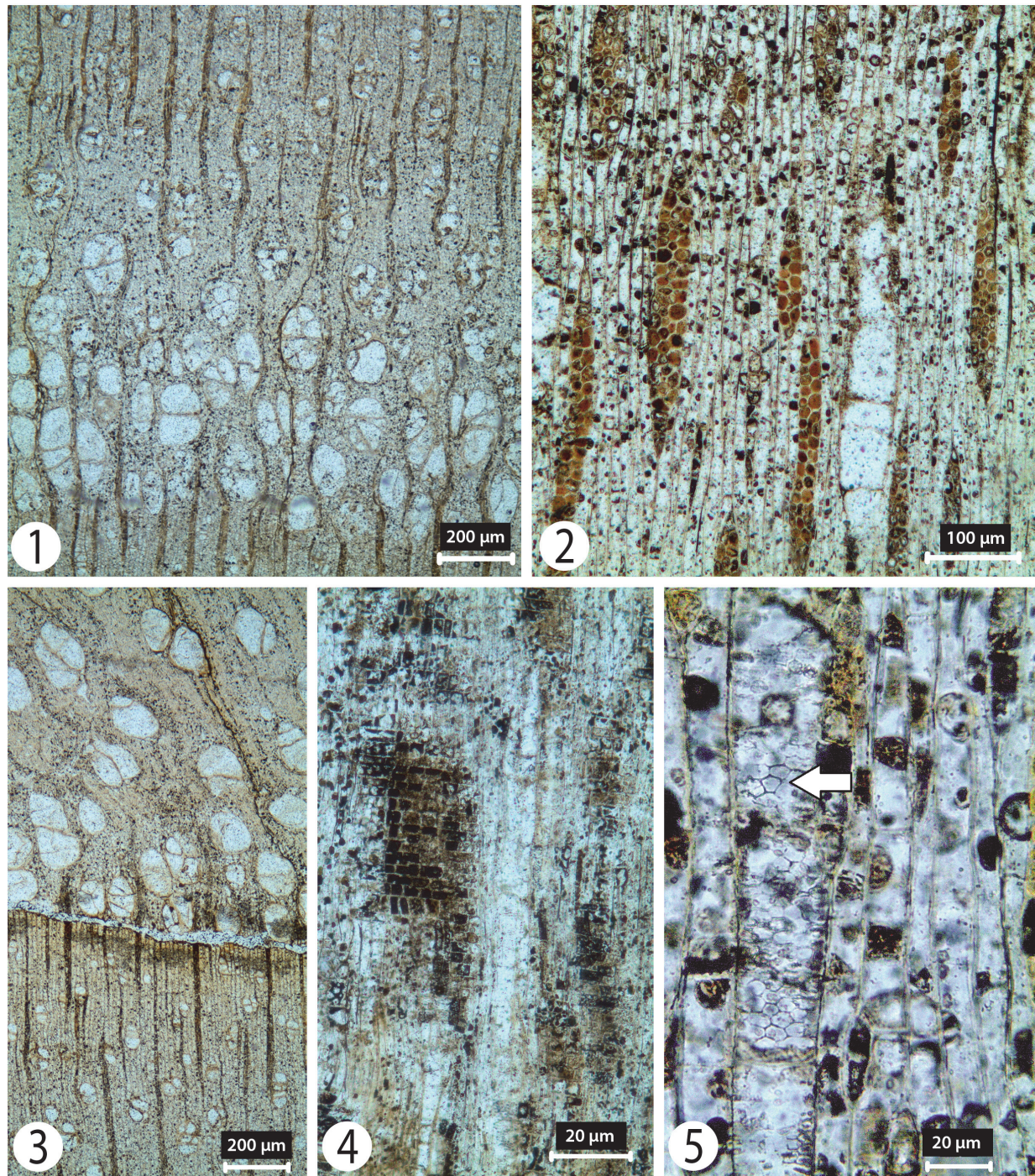
Radial longitudinal section (RLS). Rays heterocellular with procumbent body ray cells and one or more rows of upright and/or square marginal cells. Perforation plates simple. Intervessel pits are alternate or loosely opposite (Figure 14).

**Comments.** The combination of ring porous wood, relatively wide and long heterocellular rays and simple perforation plates suggest greatest affinity with *Prunus*. Today there are four species of *Prunus* growing in Turkey (Akkemik and Eminağaoğlu, 2014). Akkemik and Yaman (2012) described the wood anatomy of *Prunus cocomilla* Ten. from the Mediterranean region. The characters of this silicified wood exhibit strong anatomical similarity to the extant species but since species of *Prunus* are difficult to determine from wood anatomy alone this fossil type is assigned to *Prunus* sp. and represents the first record of fossilized *Prunus* wood from both the GVP and Turkey.

Family SALICACEAE Mirbel, 1815  
Genus SALIX Linnaeus, 1753  
*Salix* sp.

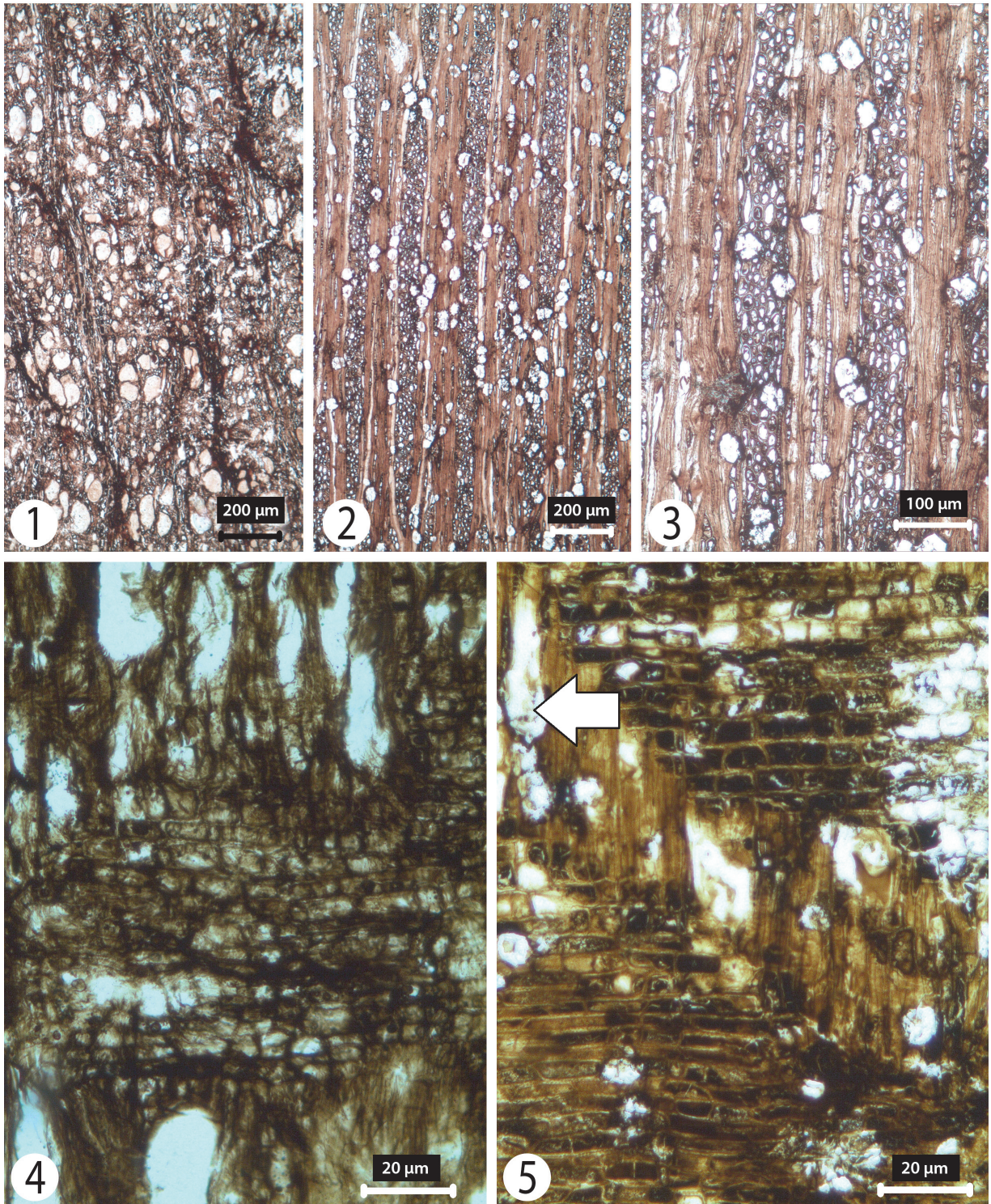
**Material.** KIR59, MEN30 and MEN34 measuring 5–7 cm in diameter and 9–13 cm high.





**FIGURE 13.** *Fraxinus*. **1**, TS showing ring porous nature of the wood. **2**, RLS of triseriate rays. **3**, TS showing distortion in the early wood caused during preservation. **4**, Rays with both procumbent and upright cells (RLS). **5**, Axial parenchyma cells and a vessel with alternate bordered pits.





**FIGURE 14.** *Prunus*. 1, TS showing relatively unclear preservation caused by distortion during preservation and ring-porous vessels. 2, TLS showing the multiseriate ray arrangement. 3, TLS showing the uniseriate rays between the multiseriate rays. 4, RLS showing the poor preservation of the rays but some cellular detail visible. 5, Better preserved rays with procumbent and upright cells, and simple perforation plate (RLS) (arrow) visible.



**Localities.** Near the village of Beypazarı-Nuhhoca (Kıraluç Site) (KIR) and in the Beypazarı-Mençeler Valley (MEN).

**Description.** This description is based on three pieces of in situ material measuring 9–13 cm in length by 5–7 cm in diameter.

Transverse section (TS). Wood diffuse porous, sometimes semi-ring porous. Growth rings measure 1–11 mm wide with distinct or indistinct growth ring boundaries. Vessels solitary or in short radial files of 2–4, elliptic or angular when grouped. Vessel tangential and radial diameters measure 62 (20–165)  $\mu\text{m}$  and 75 (17–193)  $\mu\text{m}$ , respectively. Vessel density is 158 (81–252) per  $\text{mm}^2$  (Figure 15).

Tangential longitudinal section (TLS). Rays exclusively uniseriate, very rarely partly biseriate. Rays 272 (51–994)  $\mu\text{m}$  in height and 10 (4–23)  $\mu\text{m}$  in width, with  $\geq 12$  rays per mm (Figure 15).

Radial longitudinal section (RLS). Rays appear heterocellular with procumbent body ray cells and one to two upright and/or square marginal cells. Intervessel pits are alternate and loosely opposite. Perforation plates simple (Figure 15).

**Comments.** *Salix* L. wood can often be confused with that of *Populus* L., however, the clearly heterocellular rays in this silicified wood identifies this specimen as *Salix* and has therefore been assigned to *Salix* sp. Today *Salix* L. is represented by 33 species in Turkey favouring riparian environments.

Both *Salix* L. and *Populus* L. leaves and pollen have been described from GVP (Kasaplıgil, 1977; Denk et al., 2017b) and *Salix/Populus* wood has been identified from the Hoçaş site (Akkemik et al., 2016). These genera were also common in the Miocene across western Anatolia (Güner, 2016), and support the evidence for the presence of a localized riparian type vegetation occupying lacustrine environments (Toprak et al., 1996 and references cited therein) on the GVP.

Family SAPINDACEAE Jussieu, 1789

Genus ACER Linnaeus, 1753

*Acer* sp.

**Material.** INL02, INL06, KIR02, KIR13, KIR51, KIR52 and MEN29.

**Localities.** Valleys of Beypazarı-Inozu (north side) (INL) and Beypazarı-Mençeler (MEN), and near the village of Beypazarı-Nuhhoca (Kıraluç Site) (KIR).

**Description.** The following description is based on all seven pieces of *ex-situ* petrified wood samples measuring 6.5–10 cm in length by 4–9.5 cm in diameter although the preservation in the samples

collected from Beypazarı-Nuhhoca Village (Kıraluç Site) (KIR) is poor.

Transverse section (TS). Wood diffuse-porous. Ring widths 0.4–1.2 mm. Growth ring boundaries distinct and marked by 1–2 rows of marginal axial parenchyma cells. Tangential and radial diameters of vessels measure 55 (11–158)  $\mu\text{m}$  and 66 (14–181)  $\mu\text{m}$ , respectively. Vessels circular and predominantly solitary with radial multiples of 2–4 or more. Vessel density is 206 (92–343) per  $\text{mm}^2$ . Axial parenchyma is both apotracheal diffuse and marginal. In some of the vessels dark deposits can clearly be seen. Fibers are generally thin walled (Figure 16).

Tangential longitudinal section (TLS). Rays are of two distinct sizes: generally, uniseriate, sometimes biseriate, or multiseriate commonly 4–5 seriate. Rays measure 220 (18–656)  $\mu\text{m}$  in height by 25 (4–78)  $\mu\text{m}$  in width. Intervessel pitting alternate. Helical thickening commonly present on the walls of vessels (Figure 16).

Radial longitudinal section (RLS). Rays are homocellular and composed entirely of procumbent cells. Helical thickenings can also be seen in radial walls of the wood. Intervessel pitting alternate. Perforation plates simple (Figure 16).

**Comments.** This combination of wood features suggests that these samples belong to *Acer*. This genus is relatively common in the Miocene of Anatolia and is represented at five locations across the GVP with fossil wood having been recovered at Hoçaş and Kozyaka (Akkemik et al., 2016). Further micro- and macrofossils of *Acer* L. have been recovered from Güvem (Kasaplıgil, 1977; Denk et al., 2017b).

Today *Acer* is represented by 11 species in Turkey with *Acer monspessulanum* L. being one of the species adapted to the typically xeric conditions of the Mediterranean (Akkemik and Yaman, 2012). Since the wood features of *Acer* L. species in Turkey are very similar (Yaltırık, 1971) and cannot be distinguished by wood anatomy alone, this material is referred to *Acer* sp.

Family ULMACEAE Mirbel, 1815

Genus ULMUS Linnaeus, 1753

*Ulmus* sp.

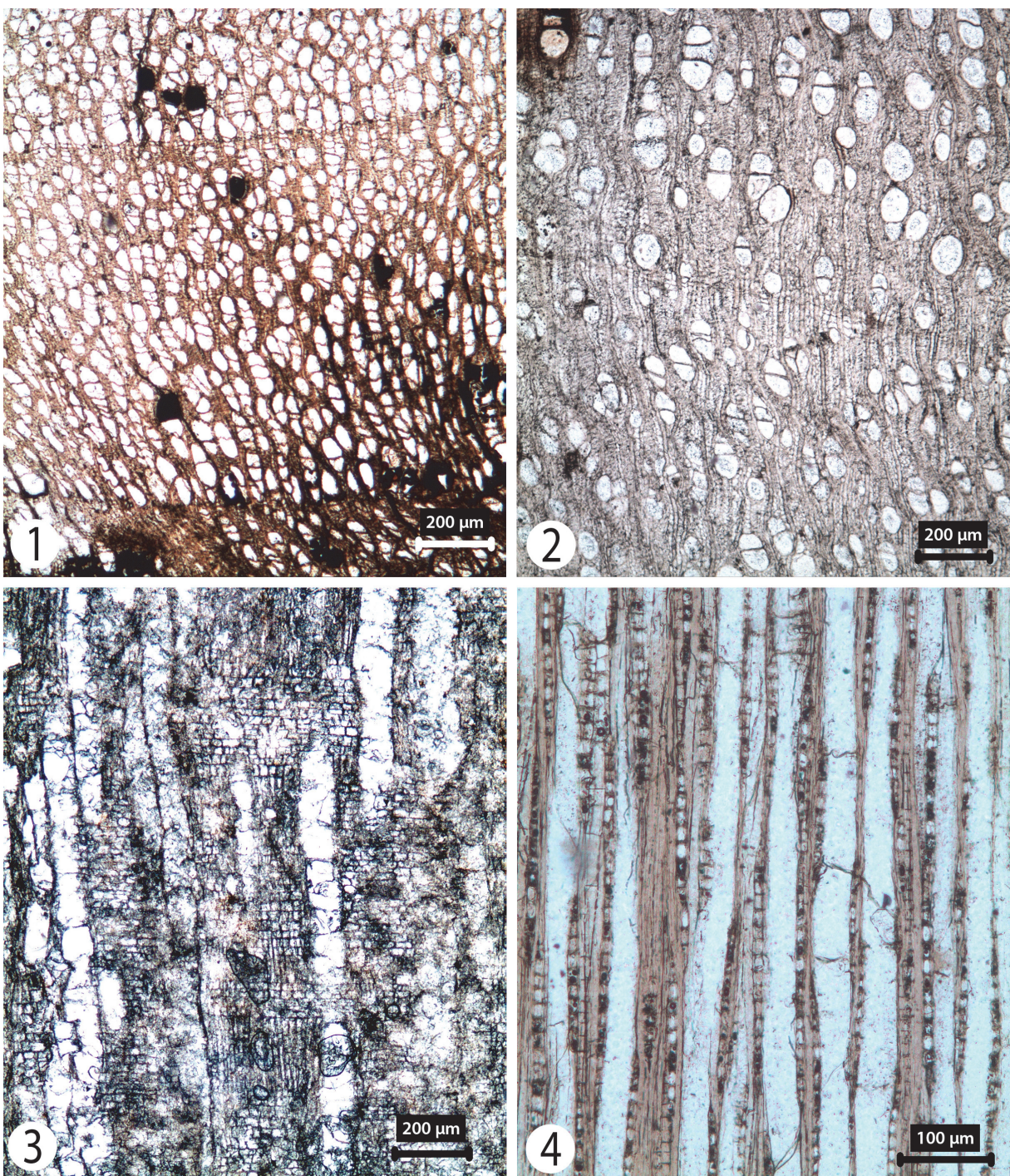
**Material.** KIR14 and KIR65.

**Localities.** Beypazarı-Nuhhoca Village (Kıraluç Site) (KIR)

**Description.** These in situ samples range from 7–9 cm in length by 8–11 cm in diameter.

Transverse section (TS). Wood ring porous; growth ring boundaries distinct; ring width 1–2 mm. Early-wood vessels are arranged in 2–3 rows, solitary or



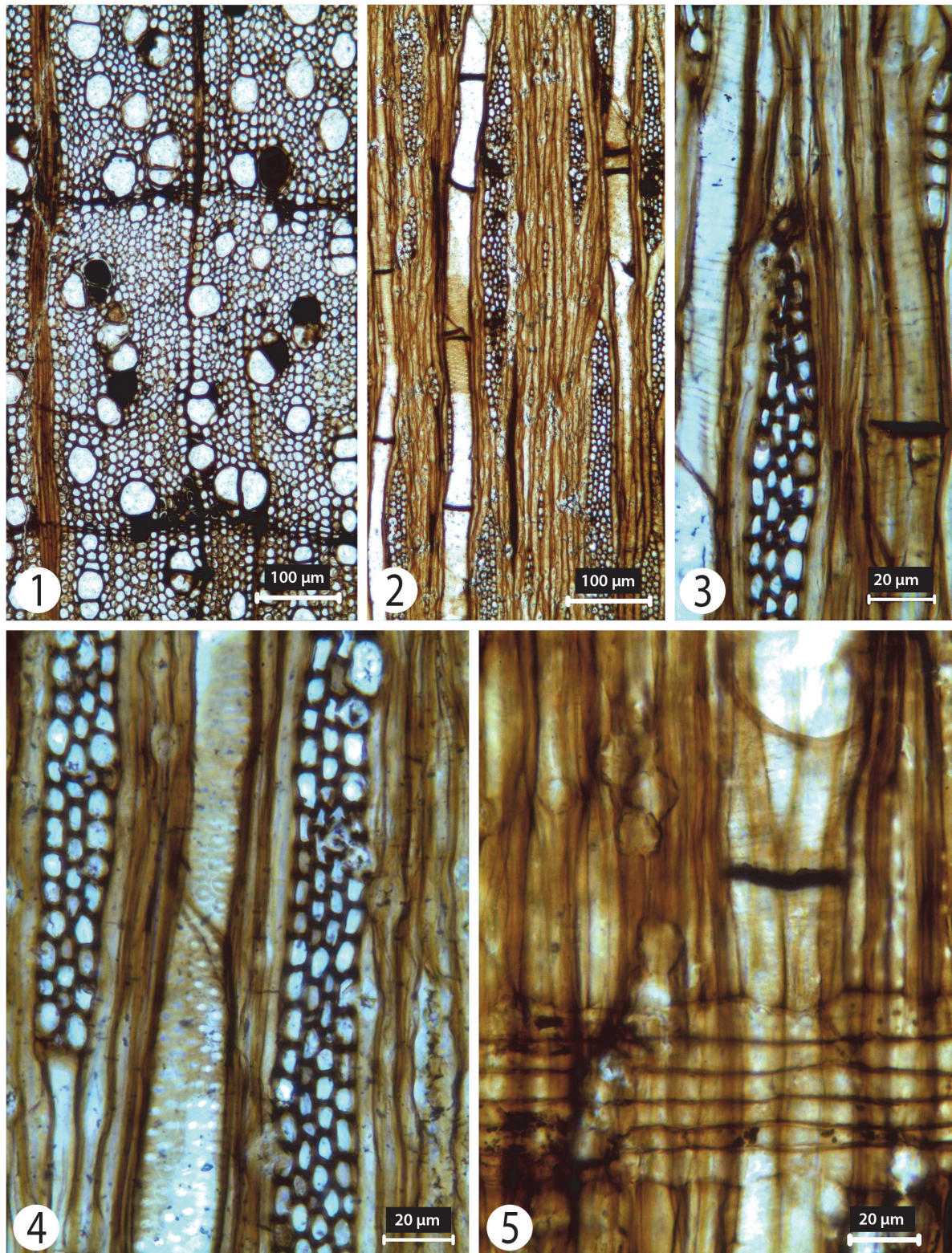


**FIGURE 15.** *Salix*. **1**, TS showing two growth rings and diffuse porosity of the vessels. **2**, TS showing the solitary and paired nature of the vessels. **3**, RLS showing ray with both procumbent and upright cells. **4**, TLS showing uniseriate rays and vessels.

in radial multiples; latewood vessels arranged in tangential bands. Vessel tangential and radial diameters measure 70 (29–117)  $\mu\text{m}$  and 100 (50–169)  $\mu\text{m}$ , respectively, in the earlywood, and 21 (4–37)  $\mu\text{m}$  and 35 (20–46)  $\mu\text{m}$ , respectively, in the

latewood. Vessel density is 11 (8–14) per  $\text{mm}^2$  in the earlywood and 180 (132–244) per  $\text{mm}^2$  in the latewood, overall vessel density is 191 (140–258)





**FIGURE 16.** *Acer*. **1**, TS showing diffuse porous arrangement of the vessels. **2**, TLS showing multiseriate ray arrangement. **3**, TLS showing rays of four cells wide and helical thickenings in the vessel wall. **4**, RLS showing alternate arrangement of bordered pits on the surface of vessels. **5**, RLS showing simple perforation plate and ray with procumbent cells.

per mm<sup>2</sup>. Axial parenchyma scanty paratracheal and diffuse (Figure 17).

Tangential longitudinal section (TLS). Rays 3–4 (up to 6) seriate measuring 34 (13–63) µm in width and 259 (120–445) µm in height (Figure 17).

Radial longitudinal section (RLS). All rays homocellular composed entirely of procumbent ray cells. Intervessel pitting alternate. Perforation plates simple (Figure 17).

**Comments.** The character combination of these two pieces of petrified woods (i.e., ring porous earlywood vessels, latewood vessels arranged in tangential continuous bands, rays 3–4 (up to 6) seriate, simple perforation plates and homocellular rays) identify this wood as *Ulmus* L.

Today four species of *Ulmus* are native to Turkey: *Ulmus canescens* Melville is adapted to xeric conditions within the Mediterranean Basin, the other three species (*U. minor* Mill., *U. glabra* Huds. and *U. laevis* Pall.) grow on more humid sites. However, regardless of their ecological preference, the wood anatomy of *Ulmus* species native to Turkey are all very similar, and it is not possible to separate the species on wood anatomy alone (Akkemik, 1995). Since the wood characters exhibited by the petrified woods studied here are very similar to all four extant species, we refer the fossil material to *Ulmus* sp.

Other micro and macrofossils of *Ulmus* have been identified from the Miocene of Turkey, namely woods from the Hoçaş site and the Kıruluç site to the northwest of Beypazarı-Nuhhoca Village (Akkemik et al., 2016), and leaves and pollen from both west Anatolia (Akgün et al., 2007; Akkiraz et al., 2011; Güner, 2016) and at Güvem (Kasaplıgil, 1977; Denk et al., 2017b). This record, therefore, extends the record of *Ulmus* L. in the local vegetation in this part of Turkey.

Genus ZELKOVA Spach, 1841

*Zelkova* sp.

**Material.** AGU15, AGU28 and KIR01.

**Localities.** Near the villages of Beypazarı-Aşağıgüney (AGU) and Beypazarı-Nuhhoca (Kıruluç Site) (KIR).

**Description.** The following description is based on three pieces of in situ wood measuring 9 cm in length by 6 cm in diameter.

Transversal section (TS). Growth rings 0.5–2.5 mm in width, ring boundaries distinct with transition from earlywood to latewood abrupt. Wood ring porous (Figure 18). Earlywood vessels in 1–2 rows; vessel clusters common in late wood. Vessel tangential and radial diameters measure 76 (23–142) µm and 118 (55–182) µm, respectively, in the

earlywood and 17 (6–36) µm and 19 (6–42) µm, respectively, in the latewood. Vessel density 325 (258–399) per mm<sup>2</sup>. Axial parenchyma unclear.

Tangential longitudinal section (TLS). Rays 1–6 seriate. Wide rays (3–6 seriate) generally 25–35 cells, or 379 (80–854) µm, with a maximum height of 51 cells; commonly >4 seriate or 41 (22–68) µm; narrow rays (1–2 seriate) 14 (6–28) µm in width and generally 11–23 cells or 213 (63–482) µm in height. Vascular/vasicentric tracheids with helical thickening present (Figure 18).

Radial longitudinal section (RLS). Rays homocellular with entirely procumbent ray cells and heterocellular with mostly procumbent body ray cells and one row of upright and/or square marginal cells (Figure 18). Axial parenchyma in strands of 2–4 cells. Prismatic crystals present in enlarged axial parenchyma cells and in the upright cells of rays. Perforation plates simple.

**Comments.** The wood of *Zelkova* Spach. is very similar to that of *Ulmus* L. but these taxa can be distinguished by the presence of heterocellular rays with one or two rows of upright cells together with procumbent body ray cells and prismatic crystals in the axial parenchyma. These characters can be seen in the fossil wood described here and thus we have assigned this material to *Zelkova* sp.

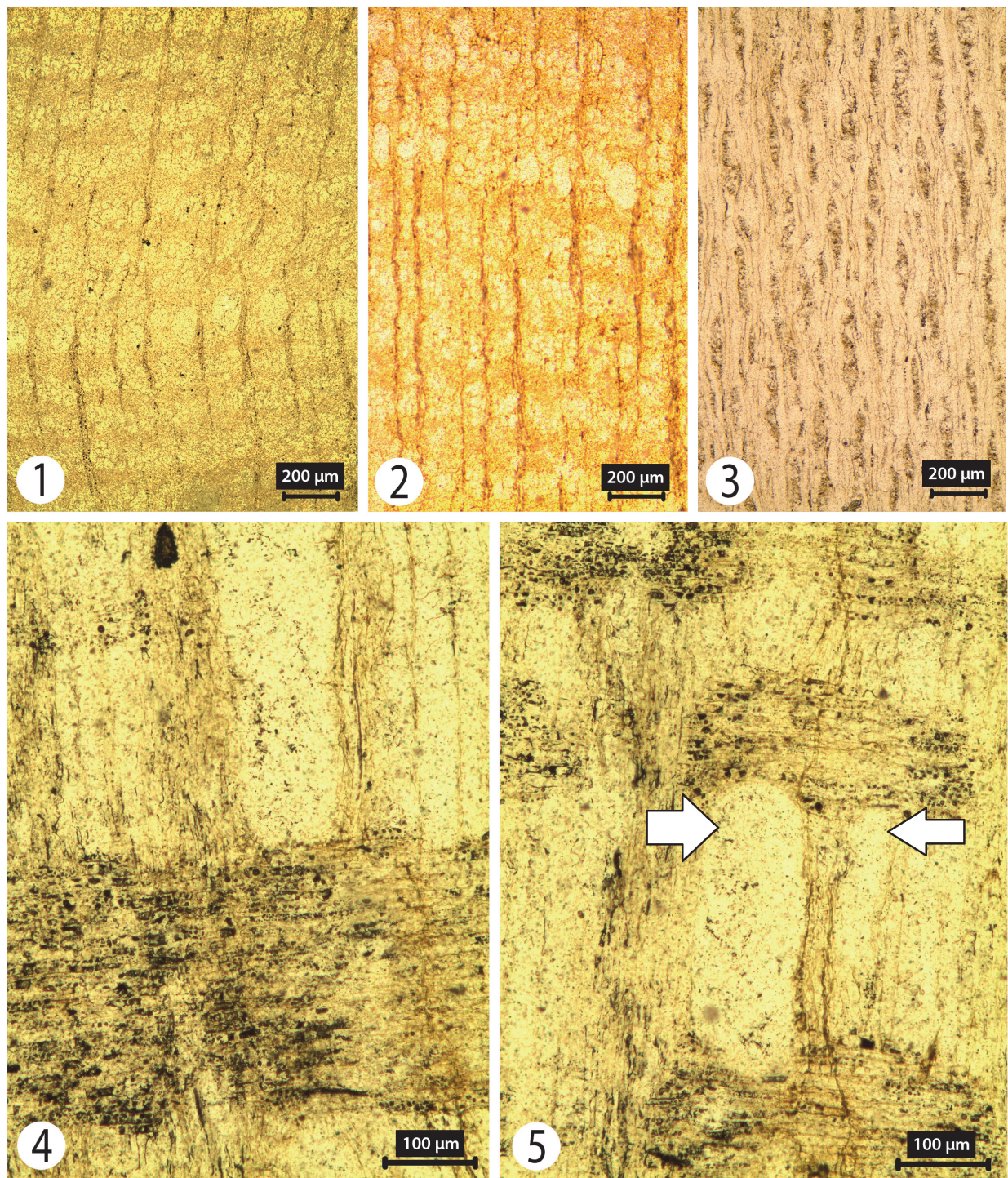
The wood identified here is the first record of *Zelkova* in the fossil wood record of Turkey, although leaf fossils have been identified from both Güvem (Kasaplıgil, 1977; Denk et al., 2017b) and west Anatolia (Güner, 2016; Bouchal et al., 2016, 2017), and pollen has been identified from west Anatolia (Üçbaş Durak and Akkiraz, 2016) and central Anatolia (Akkiraz et al., 2011). Today *Zelkova* is represented by one species, *Zelkova carpinifolia* (Pall.) K. Koch., in the flora of Turkey.

## DISCUSSION

### Forest Associations Across the Galatian Volcanic Province

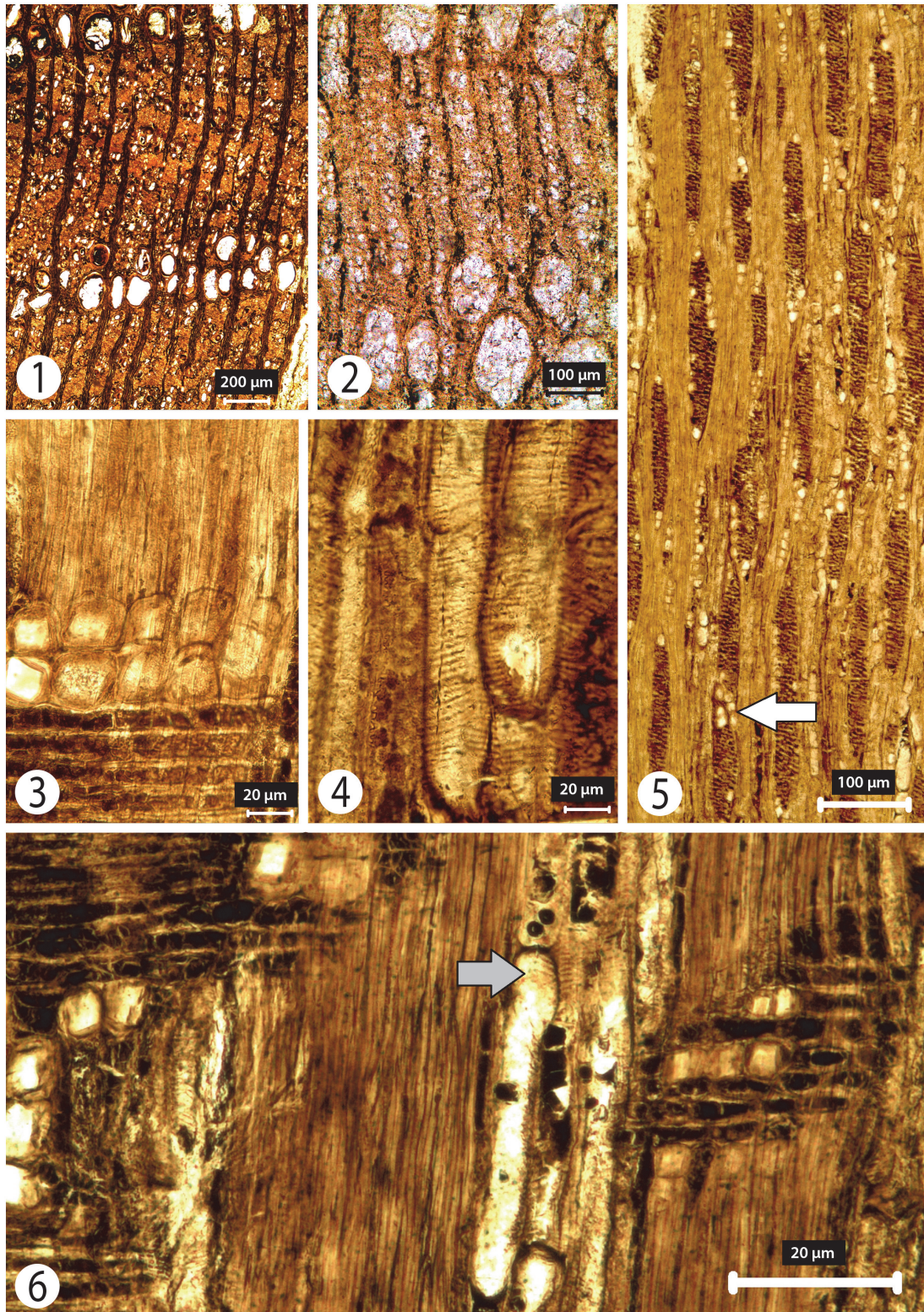
During the early Miocene the climate in western Eurasia was mild and equable, (Rögl, 1999) supporting the growth of rich forest vegetation in Northern Hemisphere as far north as the subarctic (Mai, 1995; Zachos et al., 2008; Denk et al., 2011, 2017a, 2017b, 2017c). Anatolia, along with the Lower Caucasus to the east, formed an island to the north of the eastern Mediterranean and to the south of the Eastern Paratethys (Popov et al., 2004). The Western Pontides Mountains to the north and the Taurus Mountains to the south surrounded the freshwater lakes and marshes that





**FIGURE 17.** *Ulmus* with poor anatomical preservation. **1-2**, TS showing ring porous nature of the earlywood (**2**) and wavy bands of vessels forming more or less tangential rows in the latewood (**1**). **3**, TLS showing rays of up to 4-5 cells in width. **4**, RLS showing rays with procumbent ray cells. **5**, RLS showing simple perforation plates (arrows).





**FIGURE 18.** *Zelkova*. **1**, TS showing the ring porous nature of the wood. **2**, TS showing the vessel arrangement. **3**, RLS of crystals in outer two rows of marginal ray cells. **4**, RLS showing vessels having obvious helical thickening. **5**, TLS showing rays up to 6 cells width and with crystals in their marginal cells (white arrow). **6**, RLS of rays with procumbent ray cells and one row of upright marginal cells, simple perforation plates (grey arrow) in vessels and crystals.



covered much of the Central Anatolian Plateau (Popov et al., 2004; Schildgen et al. 2012a, 2012b). The GVP forms a region in the north-western part of this Central Anatolian Plateau. According to Tankut et al. (1998), the major phase of volcanic activity on the GVP at this time resulted in the deposition of the plant-bearing strata.

Modern analogue techniques (e.g., Utescher et al., 2007; Mustoe and Leopold, 2014) use the composition of floral assemblages present today and compare them with Cenozoic fossil assemblages to provide ecological data of the palaeoenvironments from which the fossils were recovered. Although we do not feel that our palaeoassemblages are complete enough to warrant undertaking such tests just yet, this assemblage approach can provide broad ideas pertaining to the environments in which the trees grew, especially given that these are geologically relatively young assemblages.

The current hypothesis that states the landscape across the GVP were relatively heterogeneous (Denk et al., 2017b) is supported by the results of the local (wood) flora at discrete sites across the GVP found in this study. The 10 sites on the western part of the GVP lie at an altitude of 1000–1400 m whereas those in the central part of the GVP lie at 1000–1250 m (Table 1). Since there is no evidence for any significant change in elevation as a result of tectonic processes since the early Miocene, we cannot use altitudinal differences to explain the different vegetation assemblages. Given that the sites located in the western part of the GVP are geographically distinct and possibly slightly younger than those in the western part (Figure 19 and Table 2), we will, however, consider these two regions separately in the context of previously published work.

**Forests of the central GVP.** Güvem, located in the central part of the GVP, is one of the few precisely dated early Miocene assemblages of plant and animals. The age of this material is 20–18 Ma, or early Burdigalian, and many taxa are more typical of temperate and equatorial monsoon climates of SE Asia (van den Hoek Ostende, 2001). Denk et al. (2017b) undertook a major review of the plant material from this area using evidence from leaves, fruits and flower remains and published data of dispersed pollen and spores (Denk et al., 2017a, 2017c and reference cited therein). A total of 38 of the 40 known taxa identified from the GVP to date have been recorded from Güvem (Table 3). Denk et al. (2017b and reference cited therein) conclude that the landscape was varied and characterized

by a mix of trees, shrubs, lianas, ferns and fern allies. Evergreen broadleaved forests comprising *Ilex*, *Quercus* and other Fagaceae types dominated the areas surrounding lowland lakes and across large parts of both lowland and uplands (Denk et al., 2017b). Restricted open areas probably supported Fabaceae and Poaceae (Yavuz-Isik, 2008; Denk et al., 2017b). Wetland biomes comprising swamp forest communities and wet meadows surrounded the lakes and rivers where taxodiaceous swamp elements (*Glyptostrobus*) and palms flourished (Kasapligil, 1977; Denk et al., 2017b). According to Schroeder (1998) and Denk et al. (2017b), the heterogeneous nature of the landscape is further reflected in well-drained lowland broadleaved evergreen forests, mesic mixed deciduous and coniferous upland forests with *Sequoia*, *Fagus* types and conifers (*Cedrus* and *Cathaya* Chub and Kuang), and the presence of laurel forests. In the pollen flora, a more widespread vegetation assemblage is usually recorded. Yavuz-Isik and Demirci (2009) identified pollen of Pinaceae, *Ulmus*, *Salix*, *Quercus*, *Carpinus* L., *Carya* Nutt. and *Engelhardia* Lesch. ex Blume, which are similar to our findings regarding the composition of the vegetation in the western part of GVP (see below). This suggests that floral associations did indeed vary geographically across this region at this time. These authors also conclude that across the Güvem region the climate was humid and warm temperate, with a somewhat drier season in the winter. This agrees with other studies that suggest the climate in Güvem during the early to mid-Miocene was warm-temperate with a mean annual temperature for the region of 17.2–20.8°C (Benda and Muelenkamp, 1990; Akgün et al., 2007; Kayseri and Akgün, 2008; Yavuz-Isik, 2008).

At other sites across the central GVP region, research undertaken on the wood flora has been used to further knowledge of the prevailing local vegetation. In the central region of the GVP to the west of Güvem near Pelitcik, the Camlidere Fossil Forest, dated to 18.2–16.9 Ma, (Tankut et al., 1995) is thought to have been composed exclusively of '*Taxodium*' and '*Sequoia*' (Akkemik et al., 2009); 80 km to the southeast of Güvem, however, the coeval Gündül Fossil Forest comprised a riparian/wetland environment dominated by a '*Glyptostrobus*' forest (Akkemik et al., 2017). Therefore, the wood flora adds support to the idea of local wet biome floral associations postulated by Kasapligil (1977) and later Denk et al. (2017b).

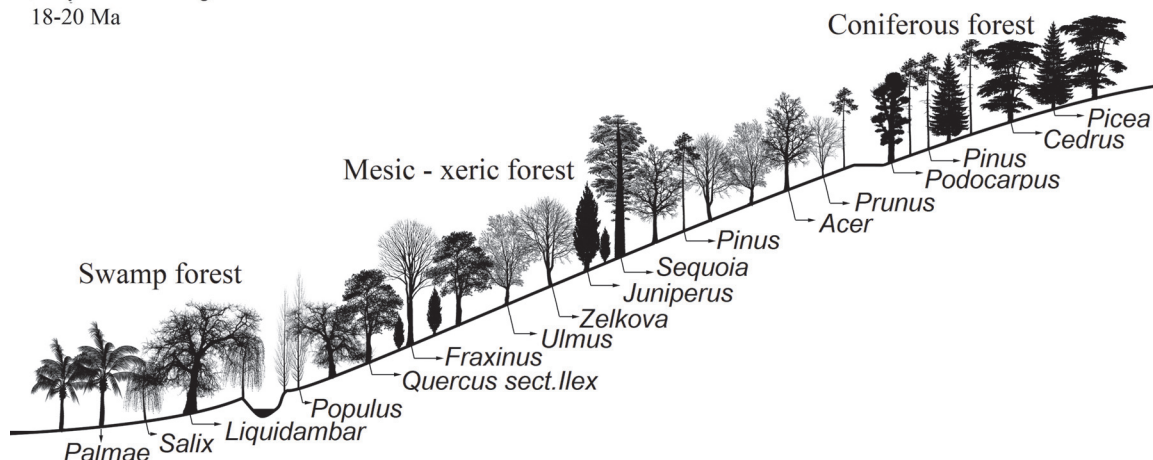
Such swamp forest elements were once widespread across the Northern Hemisphere. Today

1

## The forest trees in the western side of GVP

Early-middle Burdigalian

18-20 Ma

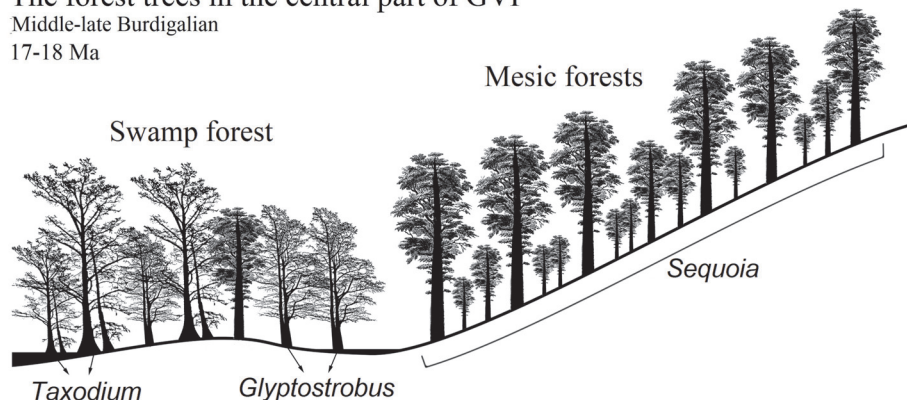


2

## The forest trees in the central part of GVP

Middle-late Burdigalian

17-18 Ma



**FIGURE 19. 1-2,** Summary reconstruction of the two distinct vegetational groups. The elevation line shows possible elevation. The riparian forest vegetation growing on the western side of the GVP at 20-18 Ma based on trees identified from the eight distinct localities (1). Compare this with (2) which summarizes the composition of the more or less contemporaneous swamp forest vegetation (18-17 Ma) growing across the central part of the GVP. (Note that ‘*Sequoia*’ type wood can include one or more of *Sequoia*, *Sequoiadendron* and *Metasequoia*).

they are no longer found in Europe or the Mediterranean, but are restricted to the south-eastern USA and parts of East Asia (Mustoe and Leopold, 2014). *Taxodium* flourishes in regions characterized by poorly drained environments, mild winters, year round rainfall and mean annual precipitation levels of >1000 mm/yr (Hoagland et al., 1996). Today *Sequoia* is restricted to a 725 km strip along the Pacific Coast of North America in moist sheltered habitats where the climate is mild and frost

free, and moisture levels remain relatively high all year (Eckenwalder, 2009). *Glyptostrobus* grows in pure stands across parts of Asia along swamp margins, on river floodplains and in deltas; the buttress base and pneumatophores evidence its adaptation to seasonal inundation (Eckenwalder, 2009). This association, as also suggested by the local wood flora in the central part of the GVP and the presence of palms (at Güvem), supports the hypothesis that the climate was warm temperate,



**TABLE 2.** A summary list of the genera in western and central GVP. The ages and tree compositions in both groups are different.

Site	Genus
Western GVP	
KOZ	<i>Salix/Populus, Cedrus, Pinus, Acer, Ulmus, Quercus sect. Ilex</i>
HOC	<i>Cedrus, Juniperus, Picea, Acer, Quercus sect. Ilex, Salix/Populus, Liquidambar, Palmae</i>
KUZ	<i>Pinus</i>
AGU	<i>Juniperus, Picea, Pinus, Fraxinus, Ulmus, Zelkova</i>
KIR	<i>Salix/Populus, Picea, Acer, Ulmus, Quercus sect. Ilex, Zelkova</i>
MEN	<i>Salix/Populus, Pinus, Acer, Prunus, Quercus sect. Ilex, Zelkova</i>
KAR	<i>Quercus sect. Ilex</i>
INL	<i>Cedrus, Acer, Quercus sect. Ilex</i>
INO	<i>Cedrus, Sequoia</i>
All Western GVP	<i>Cedrus, Juniperus, Picea, Pinus, Sequoia, Acer, Quercus sect. Ilex, Salix/Populus, Liquidambar, Palmae, Prunus, Ulmus, Zelkova</i>
Central GVP	
PEL	<i>Sequoia, Taxodium</i>
BUG	<i>Sequoia, Taxodium</i>
GUD	<i>Glyptostrobus</i>
SOG	<i>Sequoia</i>
All Central GVP	<i>Glyptostrobus, Sequoia, Taxodium</i>

mild and frost free, during the Burdigalian, and wet habitats existed across the region. Cooling of the climate during the later Miocene (Benda and Muelenkamp, 1990; Kayseri and Akgün, 2008; Yavuz-İşık, 2008) probably caused this swamp-for-est association to ultimately disappear from the GVP.

**Forests of western GVP.** Fossil evidence for the vegetation across the western part of the GVP during the Burdigalian is more limited relative to the central region (Figure 19 and Table 2). Here fossil wood floras of Aquitanian-Burdigalian age are currently the main source of evidence and provide insights into the local, rather than a more regional, vegetation composition. Material from two sites in the very western edge of the GVP have already been studied: The Hoçaş Fossil Forest and the Kozyaka Fossil Forest close to the town of Seben. Wood from the Hoçaş site, dated to between 20 and 18 Ma (Burdigalian) (Akgün et al., 2002; Kayseri and Akgün, 2008; Yavuz-İşık, 2008; Akkemik et al., 2016), comprises palms, conifers (*Cedrus*, *Picea*, *Juniperus*,) and angiosperms (*Salix/Populus*, *Acer*, *Quercus* section *Ilex* and *Liquidambar*) and represents the most diverse local flora in this region with nine of a possible 40 taxa identified from the GVP. The Kozyaka Fossil Forest (considered to be coeval with the Hocas Fossil Forest; Akkemik et al., 2016) comprises a similar mixed

flora of six taxa, namely the conifers *Cedrus* and *Pinus*, and the angiosperms *Salix/Populus*, *Acer*, evergreen *Quercus* and *Ulmus*.

The results of this study also suggest that relative to the central GVP the local flora in the western part of the GVP is more diverse (Table 3). *Quercus* section *Ilex* is the most geographically widespread taxon found at six of the nine sites followed closely by *Salix/Populus* and *Acer* found at five of the sites; *Picea* and *Cedrus* have been found at four; *Zelkova*, *Ulmus*, *Juniperus*, *Podocarpus* at two sites; *Prunus*, *Liquidambar*, *Fraxinus*, *Sequoia* and palms each only on one site.

The localities with the most diverse local floras are: The Kıraluç Site at Beypazarı-Nuhhoca Village (KAR) with seven taxa (*Pinus*, *Podocarpus*, *Acer*, *Quercus* section *Ilex*, *Salix/Populus*, *Ulmus* and *Zelkova*), Beypazarı-Aşağıgüney (AGU) with six taxa (*Juniperus*, *Picea*, *Pinus*, *Fraxinus*, *Salix/Populus* and *Zelkova*) and Beypazarı-Mençeler (MEN) with five taxa (*Podocarpus*, *Acer*, *Prunus*, *Quercus* section *Ilex* and *Salix/Populus*). The remaining localities have relatively few taxa identified to date: Beypazarı-Inozu valley (north side) (INL) had *Cedrus*, *Acer* and *Quercus* section *Ilex*; Beypazarı Inozu valley (south side) (INO) had *Cedrus* and *Sequoia*; Beypazarı-Karaşar Village (KAR) and Bolu-Kıbrısçık-Kuzca-Alanhümmet villages

**TABLE 3.** Taxonomic composition of the vegetation that grew across GVP during the early to middle Miocene. See Figure 1 for the site abbreviations and the site names. Data sources are: 1, Akkemik et al. (2016); 2, Akkemik et al. (2017); 3, Akkemik et al. (2009); 4, Kasaplıgil (1977); 5, Denk et al. (2017b); 6, Karayığit et al. (1999); and 7, Yavuz-Işık (2008).

Area	Western part of GVP (Aquitanian–Burdigalian: 24.8–18 Ma)								Central part of GVP (middle to late Burdigalian: 18.2–16.9 Ma)						Kızılcahamam 4,5,6,7
	Seben <sup>1</sup>		Kıbrısçık		Beypazarı				Güdü <sup>2</sup>		Çamlidere <sup>3</sup>		Çamlidere		Kızılcahamam
	KOZ	HOC	KUZ*	AGU*	KIR*	KAR*	MEN*	INL*	INO*	GUD	PEL	BUG*	ELM*	SO*	Güvem (Pollen and macrofossils)
<b>Gymnosperms</b>															
<i>Calocedrus</i>	-	-	-	-	-	-	-	-	-	-	-	-	-	-	+
<i>Cathaya</i>	-	-	-	-	-	-	-	-	-	-	-	-	-	-	+
<i>Cedrus</i>	+	+	-	-	-	-	-	+	+	-	-	-	-	-	+
<i>Glyptostrobus</i>	-	-	-	-	-	-	-	-	-	+	-	-	-	-	+
<i>Juniperus</i>	-	+	-	+	-	-	-	-	-	-	-	-	-	-	-
<i>Picea</i>	-	+	-	+	-	-	-	-	-	-	-	-	-	-	+
<i>Pinus</i>	+	-	+	+	+	-	-	-	-	-	-	-	-	-	+
<i>Podocarpus</i>	-	-	-	-	+	-	+	-	-	-	-	-	-	-	-
<i>Sequoia</i>	-	-	-	-	-	-	-	-	+	-	+	+	-	+	+
<i>Taxodium</i>	-	-	-	-	-	-	-	-	-	-	+	+	-	-	-
<b>Angiosperms</b>															
<i>Acer</i>	+	+	-	-	+	-	+	+	-	-	-	-	-	-	+
<i>Ailanthus</i>	-	-	-	-	-	-	-	-	-	-	-	-	-	-	+
<i>Alnus</i>	-	-	-	-	-	-	-	-	-	-	-	-	-	-	+
<i>Berberis</i>	-	-	-	-	-	-	-	-	-	-	-	-	-	-	+
<i>Betula</i>	-	-	-	-	-	-	-	-	-	-	-	-	-	-	+
<i>Carpinus</i>	-	-	-	-	-	-	-	-	-	-	-	-	-	-	+
<i>Carya</i>	-	-	-	-	-	-	-	-	-	-	-	-	-	-	+
<i>Castanea</i>	-	-	-	-	-	-	-	-	-	-	-	-	-	-	+
<i>Castanopsis</i>	-	-	-	-	-	-	-	-	-	-	-	-	-	-	+
<i>Celtis</i>	-	-	-	-	-	-	-	-	-	-	-	-	-	-	+
<i>Engelhardia</i>	-	-	-	-	-	-	-	-	-	-	-	-	-	-	+
<i>Fagus</i>	-	-	-	-	-	-	-	-	-	-	-	-	-	-	+
<i>Fraxinus</i>	-	-	-	+	-	-	-	-	-	-	-	-	-	-	+
<i>Ilex</i>	-	-	-	-	-	-	-	-	-	-	-	-	-	-	+
<i>Juglans</i>	-	-	-	-	-	-	-	-	-	-	-	-	-	-	+
<i>Laurus</i>	-	-	-	-	-	-	-	-	-	-	-	-	-	-	+
<i>Liquidambar</i>	-	+	-	-	-	-	-	-	-	-	-	-	-	-	+
<i>Magnolia</i>	-	-	-	-	-	-	-	-	-	-	-	-	-	-	+
<i>Mahonia</i>	-	-	-	-	-	-	-	-	-	-	-	-	-	-	+
<i>Myrica</i>	-	-	-	-	-	-	-	-	-	-	-	-	-	-	+
<i>Palmae</i>	-	+	-	-	-	-	-	-	-	-	-	-	-	-	+
<i>Platanus</i>	-	-	-	-	-	-	-	-	-	-	-	-	-	-	+
<i>Pterocaria</i>	-	-	-	-	-	-	-	-	-	-	-	-	-	-	+
<i>Prunus</i>	-	-	-	-	-	-	+	-	-	-	-	-	-	-	+



TABLE 3 (continued).

Area	Western part of GVP (Aquitanian–Burdigalian: 24.8–18 Ma)										Central part of GVP (middle to late Burdigalian: 18.2–16.9 Ma)						
	Seben <sup>1</sup>		Kıbrısçık	Beypazarı						Güdü <sup>2</sup>	Çamlidere <sup>3</sup>	Çamlidere		Kızılcahamam	Kızılcahamam 4,5,6,7		
	KOZ	HOC	KUZ*	AGU*	KIR*	KAR*	MEN*	INL*	INO*	GUD	PEL	BUG*	ELM*	SOG*	Güvem (Pollen and macrofossils)		
<i>Quercus</i> (section <i>Ilex</i> )	+	+	-	-	+	+	+	+	-	-	-	-	-	-	+		
<i>Quercus</i> (deciduous)	-	-	-	-	-	-	-	-	-	-	-	-	-	-	+		
<i>Salix/Populus</i>	+	+	-	+	+	-	+	-	-	-	-	-	-	-	+		
<i>Sorbus</i>	-	-	-	-	-	-	-	-	-	-	-	-	-	-	+		
<i>Tilia</i>	-	-	-	-	-	-	-	-	-	-	-	-	-	-	+		
<i>Ulmus</i>	+	-	-	-	+	-	-	-	-	-	-	-	-	-	+		
<i>Zelkova</i>	-	-	-	+	+	-	-	-	-	-	-	-	-	-	+		

(KUZ) with only one taxon—*Quercus* section *Ilex* and *Pinus* respectively (Figure 19, Tables 2 and 3).

Once again, modern analogue techniques (Utescher et al., 2007; Mustoe and Leopold, 2014) can provide useful information regarding the paleoecology of the western part of the GVP during the Aquitanian–Burdigalian. From comparisons of the assemblages of wood material identified at the sites in the western part of the GVP with the extant flora that grows along the southwestern coasts of Turkey today, we suggest a possible three different forest types may have grown in this region.

A ‘riparian forest’ association. Today this riparian association is fairly widespread along water courses and includes elements such as *Salix*, *Liquidambar*, *Alnus*, *Fraxinus*, *Ulmus*, *Platanus*, *Myrtus*, growing together with palms (*Phoenix theophrastii* Greuter), *Juniperus* and *Pinus* (Efe, 1987). Since this association is restricted to water courses it is relatively intolerant of local conditions and can grow in wet, semi-dry and dry environments. This taxa combination has been identified from the HOC site along with *Acer*, *Picea* and *Cedrus* suggesting that the fossil ‘riparian forest’ flora may have been more taxonomically diverse than its modern counterpart.

A ‘mesic-xeric’ woodland association. Today in the Mediterranean region of Turkey, under a warm temperate climate with mean annual rainfall of c. 500–1000 mm/yr, a semi-dry mesic-xeric macchie-type association (Özalp, 2000) is commonly found growing around the extant riparian forest belt at relatively low altitudes (200–1000 m). For example, in

south-western Turkey, species of *Acer* grow alongside several species of evergreen *Quercus* evergreen oaks (*Quercus ilex* L., *Q. coccifera* L. and *Q. aucheri* Jaub. and Spach) and deciduous oak species in xeric habitats of the Mediterranean region with *Pinus* and *Juniperus* (Yaltırık, 1971; Günel, 2011). During the early Miocene, evidence for *Acer* and *Quercus* can be found together in five of the nine sites in this part of the GVP (Table 3), sometimes with *Pinus* and one or more other species not found in the modern analogue. With more studies of these floras, the composition of this vegetational association and the degree of similarity to modern mesic-xeric woodlands will become clearer.

Upland mixed coniferous-angiosperm forest association. Today a conifer association comprising *Pinus*, *Cedrus*, *Juniperus* and *Abies* can be found at higher altitudes (1000–1800 m a.s.l.) (Yaltırık and Akkemik, 2012). These taxa, except *Abies*, are present in the fossil assemblages (e.g., KOZ, HOC, KUZ) and might suggest a more upland forest association of this type grew, together with *Picea*, *Podocarpus*, *Cedrus*, during the early Miocene. These conifer elements probably occurred alongside angiosperms (e.g., *Acer*, *Quercus*, *Salix/Populus*, *Ulmus*, and possibly *Fraxinus* and *Zelkova*). A link to this extant vegetation is certainly more tenuous at this stage, but the possibility does exist that drier habitats of the surrounding higher ground could have supported a more mixed conifer vegetation. Further work is needed to understand more fully the associations these taxa may have had across the GVP at this time.

The presence of *Sequoia* at the INO site probably reflects a habitat distinct from this association and more similar to the swamp forests in the central part of the GVP. However, the presence of *Cedrus*, more commonly found at higher altitudes today, alongside *Sequoia* may suggest a more diverse habitat and/or vegetation during the Miocene.

*Taxodium* and *Podocarpus* are not represented in the floristically diverse assemblage found at Güvem and constitute new records for the GVP. From nearest living relative studies *Taxodium* would also support a warm temperate environment with humid/damp growing conditions.

Paleoclimate studies undertaken by various workers suggest that a subtropical to warm-temperate climate prevailed over large parts of western Turkey (Güner, 2016), north-western central Anatolia and into Greece during the early Miocene (Velitzelos et al., 2014). The climate was characterized by mean annual temperature of 17.2–21.3°C, a cold month mean of 6.2–13.3°C, a warm month mean of 26.5–27.9°C and a mean annual precipitation of c. 1000–1325 mm/yr and a driest month precipitation level between 40–60 mm (Akgün et al., 2007; Akkiraz et al., 2011). Although the wood associations currently do not provide us with absolute levels of temperature and precipitation at this stage and simply suggest a prevailing warm temperate climate, our findings are in support of the conclusions of Denk et al. (2017a, 2017b, 2017c) and other workers cited above.

## CONCLUSION

The data from fossil wood found at 10 new fossil sites on the GVP (three localities in the central part and seven in the western part) provide a more detailed view of the vegetation associations that grew across the region during the early Miocene. The wood flora suggested evidence of taxonomically distinct local associations. The results provided further support for the current hypotheses for a warm temperate climate prevailing over a heterogeneous landscape comprising a riverine association, a mesic-xeric woodland association and possibly a mixed coniferous forest in the western part of the GVP and swamp forests associations located towards the central part of the GVP. Although this study brings the total of petrified wood genera to 15, studies of other macrofossils and pollen data suggest many more genera were present across the wider area of the GVP. More work is, therefore, needed before we can fully understand the degree of variation in the vegetation across the GVP.

## ACKNOWLEDGMENTS

This study was supported by the Research Fund of Istanbul University Project Number: YÖP-41738 and 22800.

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