

The Volhynian (late Middle Miocene) marine fishes and mammals as proxies for the onset of the Eastern Paratethys re-colonisation by vertebrate fauna

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ABSTRACT

The late Middle Miocene of South-Eastern Europe was the time of major basin-wide physical and biogeographical transformations of marine environments with mass extinction and faunal replacement. We examined new material comprising marine fishes and mammals, re-examined earlier finds, and analysed previously published data on the marine vertebrate assemblage that lived along the western and northern edges of the Eastern Paratethys. The studied materials represent marine fishes that belong to at least four families (Clupeidae, Sparidae, Bothidae, and Scombridae), including *Sarmatella doljeana*, *Sparus brusinai*, *Bothus parvulus*, and indeterminate scombrids; true seals (Phocinae), the earliest seal record in the area; small toothed whales of the family Kentriodontidae, including *Kentriodon fuchsii*, whose presence in the Volhynian is confirmed here, and two other dwarf forms; a bizarre toothed whale *Pachyacanthus*, the first record in the area between the Vienna Basin and Central Asia; baleen whales, including an indeterminate member of the family Cetotheriidae, and "*Archaeocetus fockii*". The material found so far hints to a low taxonomic diversity and biogeographical connections to both the synchronous Sarmatian fauna of the Central Paratethys and the later Bessarabian Eastern Paratethyan fauna. The taxonomic composition of the faunal assemblage and its biogeographical connections indicate that it has Central Paratethyan origin and shows an overall west to east direction of gradual dispersal, from the Central Paratethys through the Fore-Carpathian Basin and further east. The taxonomic composition of described Volhynian fish and mammal assemblages supports the idea of faunal connection and transition between the Central and Eastern Paratethys.

Gol'din, Pavel, Haiduc, Bogdan Stelian, Kovalchuk, Oleksandr, Górká, Marcin, Otryazhyi, Pavlo, Brânzilă, Mihai, Păun, Elena Ionela, Barkaszi, Zoltán, Tibuleac, Paul, and Rățoi, Bogdan Gabriel. 2020. The Volhynian (late Middle Miocene) marine fishes and mammals as proxies for the onset of the Eastern Paratethys re-colonisation by vertebrate fauna. *Palaeontologia Electronica*, 23(3):a43. <https://doi.org/10.26879/1091>

palaeo-electronica.org/content/2020/3152-the-volhynian-late-middle-miocene

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Keywords: Paratethys; Miocene; fishes; seals; cetaceans; dispersal

Submission: 21 April 2020. Acceptance: 25 August 2020.

INTRODUCTION

The late Middle Miocene of South-Eastern Europe was the time of major physical and biogeographic transformations of marine environments (Rögl, 1998; Popov et al., 2004; Palcu et al., 2015, 2017). The Paratethys, a great epicontinental sea of Eurasia, gradually lost its connection with the World Ocean due to the uplift of the Alpine-Dinarides, Carpathians and Pontides, and broke into a number of inner seas (Rögl, 1998; Popov et al., 2004; Palcu et al., 2015, 2017; Sant et al., 2017). At that time, there were two major basins, the Central and the Eastern Paratethys, with changing contours and environmental conditions. These basins were divided by the uplifting landmass of the Carpathians, and by the smaller Fore-Carpathian Basin connected to both basins (Kováč et al., 2017). Their progressing isolation led to the formation of specific local faunas and, respectively, regional stratigraphic units. During the Volhynian age (or Sarmatian s.s.), ca. 12.82 (or 12.65) – 11.8 Ma, which corresponds to the late Serravallian (Palcu et al., 2017), they formed an internally connected basin. Further isolation of the Central Paratethys led to the development of the Pannonian fauna, while in the Eastern Paratethys the endemic

Sarmatian s.l. fauna developed into the assemblages of Bessarabian and, later, Khersonian regional stages.

There are numerous records of marine vertebrates from the Middle/Late Miocene boundary, particularly from the Sarmatian s.s. of the Central Paratethys (Kazár, 2006), and the Bessarabian of the Eastern Paratethys (Brandt, 1873; Koretsky, 2001). However, few localities in the Eastern Paratethys yielded marine vertebrates from pre-Volhynian times. The fossil record of fishes from the Volhynian beds of the Eastern Paratethys is well-documented (Bogachev, 1929, 1938; Paucă, 1938; Ionko, 1954; Bogachev, 1955; Rückert-Ülkümen, 1965; Bannikov, 1986, 1989, 1993; Carnevale et al., 2006; Bannikov, 2009, 2010; Baykina, 2012, 2013a; Bannikov and Kotlyar, 2015; Baykina and Schwarzhans, 2017; Schwarzhans et al., 2017a-d). Bogachev (1955) provided a list of fishes including clupeids and a few other marine fish taxa based on materials from the Volhynian deposits near Amvrosiivka (eastern Ukraine). However, his material is considered problematic (Bannikov, 2010) since it was not described properly and also was probably lost thus unavailable for revision. Macarovici (1970) reported a small fish assemblage from the

Moldavian Platform; those specimens are re-described here. Bannikov (2009) described the early Volhynian fish assemblage from Naslavcea, Republic of Moldova, and noticed that it was different from the contemporary fauna of the North Caucasus.

Contrary to fish records, data on Volhynian marine mammals are scarce. A fish-bearing locality and a few toothed whales ("dolphins") were reported from the Caucasus and the Caspian region (Dombrovsky, 1927; Kirpichnikov, 1954; Mchedlidze, 1964; Bannikov, 2009), whereas the western part of the basin was less studied (Kudrin and Tatarinov, 1966). A few cetaceans were reported from the Moldavian Platform in Romania (Macarovici, 1947; Macarovici and Zaharia, 1968; Ionesi and Galan, 1988; Figure 1).

Here we report newly collected marine vertebrate specimens from Romania and Ukraine with their tentative classification, a review of previous finds, a revision of the localities, and respective biogeographical implications.

GEOLOGICAL FRAMEWORK

The seas of the Central and Eastern Paratethys were interconnected during transgressive events of the Serravallian time, which was repre-

sented by the regional stratigraphic units of the Konkian (or late Badenian in the Central Paratethys) and Volhynian (or Sarmatian s.s. in the Central Paratethys) (Figure 2). During the transgressive phases both basins had fully marine conditions; however, they were separated during regressive phases that led to a gradual decrease in salinity with certain salinity crises (Popov et al., 2004; Ter Borgh et al., 2014; Vernyhorova, 2015). The newly discovered and the revised material mostly comes from the Fore-Carpathian Basin (Podolian Platform): from the Ukrainian localities of Kolubaiivtsi, Hrushivtsi, and Vilkhovets, and from the Romanian localities of the Eastern Carpathian Foreland (Moldavian Platform): Stăuceni, Leucușești, Stâncești, Basarabi, and Drăgușeni (Figures 1, 2).

A number of other findings and localities were reported from these areas (Khotin: Macarovici, 1947; Khonkvitsi: Hofstein, 1965; Brykiv: Kudrin and Tatarinov, 1966; Ghireni, Cordăreni, Rânghilești, Mitoc-Prut, Hănești, Săveni-Ghireni, Drăgușeni: Macarovici and Zaharia, 1967, 1968; Fălticeni, Voinești, Erbiceni, Aroneanu, Românești: Macarovici, 1970; Darabani: Ionesi and Galan, 1988; Basarabi: Macarovici, 1970, Codrea et al., 2014) (Table S1 in Supplementary Material). The

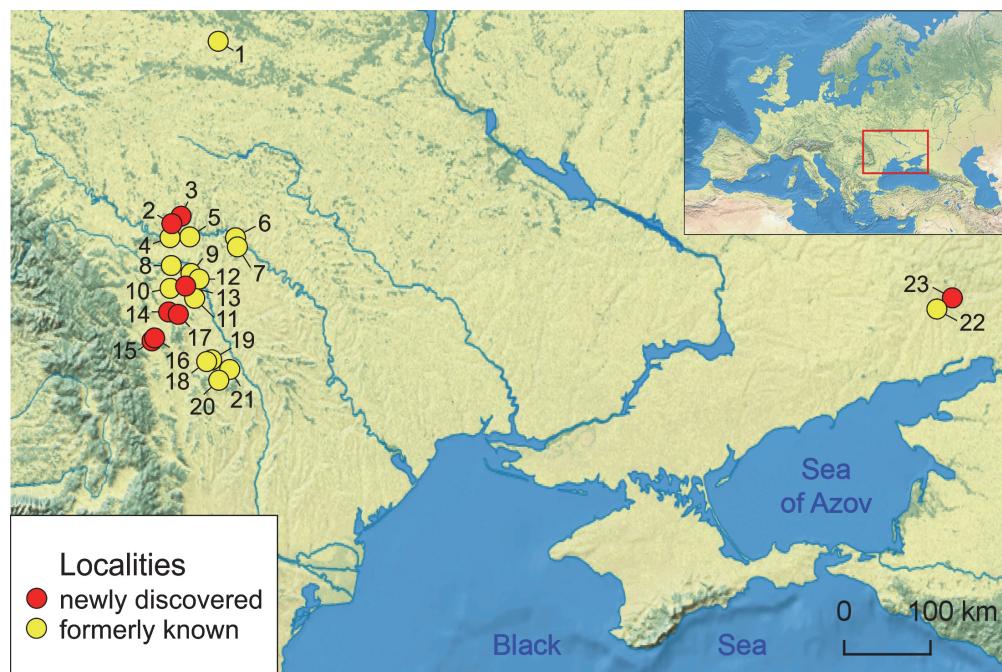


FIGURE 1. Geographic location map of the Volhynian fossil-bearing sites: 1 – Brykiv; 2 – Vilkhovets; 3 – Kolubaiivtsi; 4 – Khotin; 5 – Hrushivtsi; 6 – Khonkvitsi; 7 – Karpov Yar (Naslavcea); 8 – Darabani; 9 – Ghireni; 10 – Cordăreni; 11 – Hănești; 12 – Mitoc; 13 – Drăgușeni; 14 – Stâncești; 15 – Leucușești; 16 – Basarabi; 17 – Stăuceni; 18 – Erbiceni; 19 – Românești; 20 – Aroneanu; 21 – Voinești; 22 – Amvrosiivka; 23 – Saur-Mohyla.

| Time (Ma) | Series | | Standard stages | | | | Lake Pannon Vienna Basin Biozones (Harzhauser and Piller,2004) | Northern Black Sea Biozones (Vernyhorova, 2015) | Carpathian Foredeep Basin Biozones (Poland-Ukraine) (Studencka et al.,1998) | Moldavian Platform Biozones (Ionesi,1991) |
|-----------|---------|--------------|------------------------------|---------------------------|---------------------------|--|---|---|--|---|
| | | | Central Paratethys Substages | | | | | | | |
| | Upper | Tortonian | Pannonian | Central Paratethys Stages | Eastern Paratethys Stages | | | | | |
| 11 | Middle | Serravallian | Badenian | Lower Bessarabian | Middle Bessarabian | Upper Bessarabian | <i>Mytilopsis hoernesii</i> | <i>Mactra caspia</i> <i>Mactra bulgarica</i> | | <i>Mactra bulgarica</i> <i>Mactra orbiculata</i> |
| 12 | Kossian | Sarmatian | Volhyrian | Volhyrian | Volhyrian | Upper Ervilia beds Lower Ervilia beds | <i>Mytilopsis ornithopsis</i> | <i>Plicatiforma fittoni</i> | <i>Plicatiforma fittoni</i> | <i>Plicatiforma fittoni</i> , <i>Gibbula beaumonti</i> <i>Mactra macarovii</i> , <i>Congeria</i> , <i>Hydrobia</i> , <i>Neritina</i> <i>Mactra fabreana</i> |
| 13 | Konkian | | Mohrensternia | | | | <i>Sarmatimactra vitaliana</i> | <i>Mactra vitaliana pallasi</i> | <i>Mactra vitaliana pallasi</i> | <i>Mactra pallasi</i> <i>Tapes gregarius ponderosus</i> |
| | | | | | | | | <i>Plicatiforma plicata</i> <i>Ervilia dissita</i> | <i>Mactra eichwaldi</i> <i>Plicatiforma plicata</i> | <i>Plicatiforma plicata</i> , <i>Potamides mitralis</i> <i>Mactra eichwaldi</i> , <i>Ervilia</i> |
| | | | | | | | | <i>Obsoletiforma lithopodilica</i> <i>Abra reflexa</i> | <i>Mactra eichwaldi</i> <i>Plicatiforma pseudoplicata</i> | <i>Mohernstermia</i> sp. <i>Ocenebrina sublavata</i> , <i>Obsoletiforma lithopodilica</i> |
| | | | | | | | <i>Flabellipecten besseri</i> | <i>Ervilia pusilla</i> <i>Pavivenus konkensis</i> | <i>Flabellipecten besseri</i> <i>Pallolum bittneri</i> <i>Aequipecten elegans</i> | <i>Chlamys elegans</i> |

FIGURE 2. Correlation of the Central and Eastern Paratethyan regional stages with standard chronostratigraphy and magnetostratigraphy modified after Harzhauser et al. (2004), Studencka, (1999), Ionesi (1991), and Vernyhorova (2015).

studied area evolved as a peripheral foreland basin related to the Carpathian frontal movement owing to the flexural subsidence of a craton under an orogenic belt (Oszczypko and Oszczypko-Clowes, 2012). The deposits here are characterised by higher bed tilts or slightly folded structures as compared to the almost horizontal beds found further east (Grasu et al., 2002). The studied area was predominantly covered by shallow marine and brackish water environments within the Galician Gulf (Popov et al., 2004). In addition, Saur-Mohyla locality is situated to the east and represents the northern edge of the Eastern Paratethys (see Figure 1).

GEOLOGICAL SETTING

The geologically youngest marine vertebrate fossil specimens of Volhyrian age were collected from Kolubaivtsi (Ukraine), from south of Stăuceni (Romania), and from Darabani (Romania). The 10 m thick section in Kolubaivtsi is composed of laminated clays with intercalations of limy clays and marls, and with admixture of phosphorite concretions. Its early Volhyrian age was determined from

the presence of the bivalves *Inaequicostata inopinata* and *Abra* spp. (Jasionowski et al., 2006). In Stăuceni, the 6 m high outcrop includes compacted or laminated ash-grey clays with fine sand laminae, dated as early Volhyrian (Brâncilă, 1999). In Darabani and Drăgușeni, there are several bivalve species (including *Inaequicostata inopinata* and *Obsoletiforma lithopodilica*) and foraminifers (*Cycloforina karreri ovata*, *Cycloforina karreri karreri*, *Elphidiella serena*, and *Elphidium regnum*) indicating the early Volhyrian age (Ionesi and Galan, 1988), as well as in the localities reported by Macarovici and Zaharia (1967, 1968). The fossil-bearing site in the abandoned quarry near Khonkvitsi (Vinnytsia Oblast, Ukraine) is possibly of the same age, where a 15 m thick Volhyrian marl layer is overlying 20 cm thick phosphorite-bearing sandstones above the laminated Badenian sandstones. Also, the Volhyrian marl is overlying the Badenian sandstones further northwards.

The late Volhyrian localities include Leucușești, Stâncești, Basarabi, Voinești, Erbiceni, Aroneanu, and Brykiv. A fish specimen comes from a marl bed drilled by the F. 204H Antilești borehole at 67.5 m (ca. 365 m absolute altitude) in the

Dealul Nucului (Leucușești), in the foredeep depo-zone; it is dated as late Volhyanian of the Fălticeni-Boroia Formation (or Fălticeni Formation) based on the presence of the mollusc assemblage (Tibuleac, 2009).

The locality near Basarabi is where a succession of marl, sand, and thin sandstone beds of the latest Volhyanian age are exposed along the Pâraul lui Gheorghe brook (Codrea et al., 2014 and references therein); the vertebrate specimens come from the sandstone layer, which is younger in age than the Dealul Nucului horizon. The Stâncești locality is represented by sandy deposits, dated as late Volhyanian based on the mollusc fauna (Ionesi, 1968).

The fish assemblage described by Macarovici (1970) from the Moldavian Platform was dated as late Volhyanian according to the foraminiferal assemblage identified by Paghida-Trelea (1969). Brykiv in western Ukraine is dated at the same age based on the occurrence of *Mactra eichwaldi* and *Envilia dissita* (Tatarinov, 2000).

Hrushivtsi was reported as a locality where the Volhyanian marl succession is overlying Badenian deposits, without further details on dating (Tatarinov, 2000); the fish-bearing bed is devoid of malacofauna. Also, the deposits from Vilkhovets can only be dated as Volhyanian, without more precise data. The Saur-Mohyla locality is situated in the easternmost part of Ukraine. The marine mammal specimens found here were collected from a borehole at a depth of 5.4–5.6 m in 1951, and their age was identified as Volhyanian in the field, without details. However, as seen from palaeogeographic reconstructions, this area was flooded by the sea during the late Volhyanian time. The nearby locality of Amvrosiivka (Bogachev, 1955) probably is of the same age.

MATERIAL AND METHODS

The examined material (Tables S1-S3) is housed at the Alexandru Ioan Cuza University Department of Geology Collections (AICUPM), National Museum of Natural History, National Academy of Sciences of Ukraine (NMNHU-P), Museum of the Faculty of Geology, University of Warsaw, Poland (MWGUW), and Chernivtsi National University, Ukraine (CNU). The newly found material was collected in 2018 by one of the authors (BH) from an open pit in Stăuceni and from Basarabi, as well as by another author (PT) in Leucușești at Dealul Nucului (all above-mentioned localities are from Romania). The comparative materials were examined in the Hungarian Natural

History Museum (HNHM) and the Georgian National Museum (GNM).

The material is represented by complete and partial skeletons, isolated bones, and scales. It was identified based on comparative observations in reference collections and published data (Carnavale et al., 2006; Kazár, 2005, 2006; Baykina, 2013a, b; Schwarzhans et al., 2017d). The map of localities was created in QGIS 3.6 using open domain layers (Natural Earth).

SYSTEMATIC PALAEONTOLOGY

Class ACTINOPTERYGII sensu Goodrich, 1930
 Order CLUPEIFORMES Bleeker, 1859
 Family CLUPEIDAE Cuvier, 1817
 Genus *SARMATELLA* Menner, 1949
Sarmatella doljeana (Gorjanović-Kramberger, 1884)

Figure 3A-B

Material and localities. Two incomplete articulated skeletons; Leucușești, Hrushivtsi.

Description. The anterior part of the body (Figure 3A) is well preserved; however, many bones were destroyed and are represented by fragments. The body is elongated and laterally compressed. Its dorsal profile is straight, while the abdomen is moderately convex. The head is large (head length 14.1 mm, head depth 9.8 mm, preorbital distance 4.3 mm, postorbital distance 6.5 mm). The skull roof is almost straight, and the snout is pointed. The relatively large orbit (orbital diameter 3.3 mm, vertical diameter 2.6 mm) is oval. There are two auditory capsules: bulla prootica is 2.3 times larger than bulla pterotica. The mouth is small and terminal. The lower jaw is 7.2 mm long. It projects beyond the upper jaw and is articulated with the skull behind the anterior orbital margin. The vertebral column extends above the longitudinal axis of the body and consists of elongated vertebrae. The dorsal fin is preceded by 10 supraneuralia. Another specimen (Figure 3B) is a caudal part of the body. The minimum body depth is 3.5 mm. The caudal fin is deeply forked and has 19 principal rays and 10 additional ones. Its skeleton consists of two epurals, six hypurals, and one parhypural fused with the first preural vertebra.

Remarks. The studied specimens share a set of features diagnostic for *Sarmatella doljeana* (Gorjanović-Kramberger, 1884): the number and relative size of auditory capsules, projection of the lower jaw and the place of its articulation with the skull, the number of supraneuralia, as well as the structure of the caudal fin (Baykina, 2013a). Also, there are a few complete and subcomplete skele-



FIGURE 3. Fish remains from the Volhynian beds of Romania and Ukraine: **A-B** – *Sarmatella doljeana* (Kramberger, 1884), anterior part of the body (**A**), and caudal part (**B**), Leucuşeşti; **C** – Clupeinae gen. et sp. indet., isolated scale, Voineşti; **D-E** – Scombroidei indet., caudal part (**D**), Erbiceni, and middle part of the body (**E**), Aroneanu; **F** – *Sparus brusinai* (Kramberger, 1882), skeleton, Hrushivtsi; **G-H** – *Sparus* cf. *brusinai* (Kramberger, 1882), right dentary in lateral (**G**) and dorsal view (**H**), Pârâul lui Gheorghe; **I** – *Bothus parvulus* (Kramberger, 1883), body imprint, Româneşti. Scale bars equal 2 mm in C, 5 mm in **A-B**, **D-E**, **G-I**, and 2 cm in **F**.

tons provisionally assigned to *Sarmatella doljeana* in the collection from Hrushivtsi (Ukraine).

Clupeinae indet.

Figure 3C

1970 *Clupea* sp.; Macarovici, pl. 1, fig. 7.

Material and locality. Scale; Voineşti.

Description. A single isolated scale from Voineşti (depth 398–400 m) is scutate in shape; it is 7.0 mm wide and 6.2 mm long. The edges of the scale are rounded except for the posterior field, which is moderately crenulated. The anterior field is smooth. There are five vertical primary and secondary grooves arranged symmetrically on the surface of the scale.

Comparison. The scale described is similar to those associated with the posterior body part of "*Clupea*" *sardinites* (Szymczyk, 1978, p. 390, table 1; p. 395, figs. 4d, 5c) in the number and arrangement of grooves, as well as the shape and length-width ratio of the scale plate.

Remarks. The scale resembles that in herrings in shape and morphology (Lagler, 1947; Szymczyk, 1978). It was assigned to Clupeinae based on the absence of nodules at the middle and lateral fields, unlike those in Alosinae (Peabody, 1928; Szymczyk, 1978). We suggest that the diagnostic importance of clupeid squamation should be limited by subfamily level since comparative studies in case of joint findings of scales and skeletons are absent.

Order PERCIFORMES Bleeker, 1859

Family SPARIDAE Bonaparte, 1832

Genus *SPARUS* Linnaeus, 1758

Sparus brusinai (Gorjanović-Kramberger, 1882)

Figure 3F

2000 *Serranus* sp.; Tatarinov.

Material and locality. Almost complete skeleton; Hrushivtsi.

Description. The body is 29 cm long; its dorsal profile is convex, while the ventral one is rather flattened. Maximum body depth is about 9.5 cm (38% standard length (SL)) at the beginning of the dorsal fin. The head is noticeably shorter than body depth comprising 30% SL. The orbital diameter equals 22% head length. There are small caniniform and molariform teeth arranged in a few rows on dentary. Three large supraneurals are seen anterior to the dorsal fin, which consists of nine preserved spines and 12 soft rays. The number of rays in the pectoral and ventral fins cannot be counted. The anal fin has three spines, the anteriormost of which is the shortest, and the second one is the longest

and robust. The forked caudal fin consists of 17 principal rays.

Remarks. The general body shape, the presence of various tooth morphotypes, multi-row dentition, smooth posterior margin of the preopercle, three large supraneurals, the structure of dorsal, anal, and caudal fins, as well as a number of metric characters (maximum body depth, head length, orbital diameter) allowed us to identify this specimen as *Sparus brusinai*.

Sparus cf. brusinai (Gorjanović-Kramberger, 1882)

Figure 3G-H

Material and locality. Dentary; Basarabi.

Description. The specimen is represented by a broken robust right dentary 26.9 mm long and 17.8 mm high. The dental field is convex with traces of sub-circular tooth crown bases on its surface. Teeth are arranged at least in three rows; their diameter varies within 1.5–2.0 mm. The diameter of the teeth in the symphyseal part is 1 mm. The symphyseal part is slightly isolated, with an oval pore on its outer edge. The symphysis is 4.6 mm high. The ascending part of the dentary is broad and convex, whereas the descending part is narrow and concave.

Comparison. The studied specimen is generally similar to the dentary of *Sparus brusinai* (NMNHUP 53/4219, stored in the Department of Palaeontology of the National Museum of Natural History, National Academy of Sciences of Ukraine) described by Tarashchuk (1957) from the Badenian deposits of Dobryvody (western Ukraine), although it is slightly larger.

Remarks. The dentary was assigned to Sparidae, and more precisely to *Sparus*, due to its shape and multi-row dentition. Proportions of the specimen as well as characteristic tooth bases allowed us to attribute it tentatively to *Sparus brusinai* (Gorjanović-Kramberger, 1882).

Suborder SCOMBROIDEI Bleeker, 1859

Scombroidei indet.

Figure 3D-E

1970 *Caranx* sp. (? *gracilis* Kramberger); Macarovici, pl. 1, fig. 5.

1970 *Caranx* sp.; Macarovici, pl. 1, fig. 6.

Material and localities. Incomplete skeletons; Aroneanu, Erbiceni, Kolubaivtsi.

Description. There are two incomplete fish body imprints obtained from boreholes. One of them, 36 mm long, comes from Aroneanu (depth 280–282 m) and corresponds to the middle part of the body (Figure 3E). The body is 14.8 mm high. The vertebral column is located above the body midline, and

its preserved part consists of 18–20 subrectangular vertebrae associated with long and thin arcuate ribs almost reaching the ventral margin of the abdomen. There are two deep impressions behind the head. The entire body is covered by large scales which are twice as high as long, with rounded edges. The specimen from Erbiceni (depth 230–232 m) is represented by the caudal part of the body (Figure 3D). Its dorsal profile is almost straight, whereas the ventral margin is sharply narrowing caudally. The depth ratio of dorsal and ventral portions of the body changes in anteroposterior direction from 1:2 to 1:1. The minimum body depth is 3.2 mm. Vertebrae are connected with long neural and haemal spines oriented backward at 40–45° to the longitudinal body axis. The anal-fin pterygiophores are robust. Also, the specimen from Kolubaivtsi is represented by a similar, although less preserved, caudal part of the body.

Remarks. All the examined specimens were diagnosed as Scombroidei according to the body shape, vertebrae and distinct squamation represented by high subrectangular scales (Monsch and Bannikov, 2012). We identified them only to suborder level because of poor preservation.

Order PLEURONECTIFORMES Bleeker, 1859

Family BOTHIDAE Regan, 1910

Genus *BOTHUS* Rafinesque, 1810

Bothus parvulus (Gorjanović-Kramberger, 1884)

Figure 3I

1970 *Caprovesposus* sp.; Macarovici, pl. 1, fig. 4.

Material and locality. Incomplete skeleton; Românești.

Description. An incomplete body of a sinistral flatfish (SL ca. 16.5 mm) with a reduced and partly broken caudal peduncle. The body is oval, 7.6 mm wide. The head is deep (5.9 mm or 128% head length) and relatively short (4.6 mm; 28% SL). The skull is asymmetrical, and its neurocranial part is roughly triangular. The small mouth (ca. 5% SL) is oblique, directed 60° to central body axis. The opercle is large and ovoid, and the preopercle is slender and boomerang-shaped. The preserved vertebral column is composed of 35 subrectangular vertebrae, including 10 abdominal and roughly 25 caudal ones. This assumption is based on the length of a vertebra and the total length of the vertebral column. The neural spines are long, and a few anterior abdominal spines bear conspicuous kinks at the bases. Imprints of myorhabdoi are visible at the distal tips of neural and haemal spines. Anterior pterygiophores of the dorsal fin are hourglass shaped. There are 17 principal rays in the

caudal fin. The number of rays in other fins is unknown because of incompleteness of the specimen.

Comparison. The specimen described is almost identical to those of *Bothus parvulus* from the Sarmatian s.s. beds of Dolje (Croatia) in size, body and head proportions (Schwarzans et al., 2017d, fig. 3, table 2), and in the total number of vertebrae. This specimen differs from the Middle Miocene *Bothus* sp. from the North Caucasus (Carnevale et al., 2006) in having a larger size, a greater number of precaudal vertebrae, a shorter and deeper head, in the shape of the preopercle, and in the structure of anterior neural spine bases.

Remarks. The specimen is assigned to Pleuronectoidei based on body shape, asymmetrical skull and the dorsal fin extension over the orbital region (Chapleau, 1993; Carnevale et al., 2006). The presence of myorhabdoi and 17 rays in the caudal fin are diagnostic for Bothidae (Hensley, 1977; Hensley and Ahlstrom, 1984; Carnevale et al., 2006; Schwarzans et al., 2017d). Relationships between separate taxa within this family based on osteological characters remain not completely resolved (Chapleau, 1993; Chanet and Schultz, 1994). However, the shape of anterior dorsal fin pterygiophores and their insertion into cavities of the supraoccipital crest and the presence of robust rectangular haemal spines (Chanet and Sorbini, 2001; Carnevale et al., 2006) support its assignment to the genus *Bothus*. A boomerang-shaped preopercle and a kink at the base of anterior neural spines of the dorsal fin (Schwarzans et al., 2017d) indicate that it is *Bothus parvulus*.

Class MAMMALIA Linnaeus, 1758

Order CARNIVORA Bowdich, 1821

Family PHOCIDAE Gray, 1821

Subfamily PHOCINAE Gray, 1821

Phocinae indet. 1

Figure 4

Material and locality. Vertebrae, ribs, scapula, and forelimb bones; Kolubaivtsi.

Specific features of anatomy. The material was identified as a member of the Phocinae by the presence of an epicondylar foramen in the humerus (Dewaele et al., 2018a). Meanwhile, it shares a high trochlear crest of the humerus with Monachinae (Koretsky and Grigorescu, 2002; Dewaele et al., 2018b). It is characterised by a smooth distal termination of the long deltopectoral crest reaching the condyle, and a relatively straight posterior margin of the ulna.

Comparison. The specimen has a body size comparable to phocine seals found in later (Bessara-

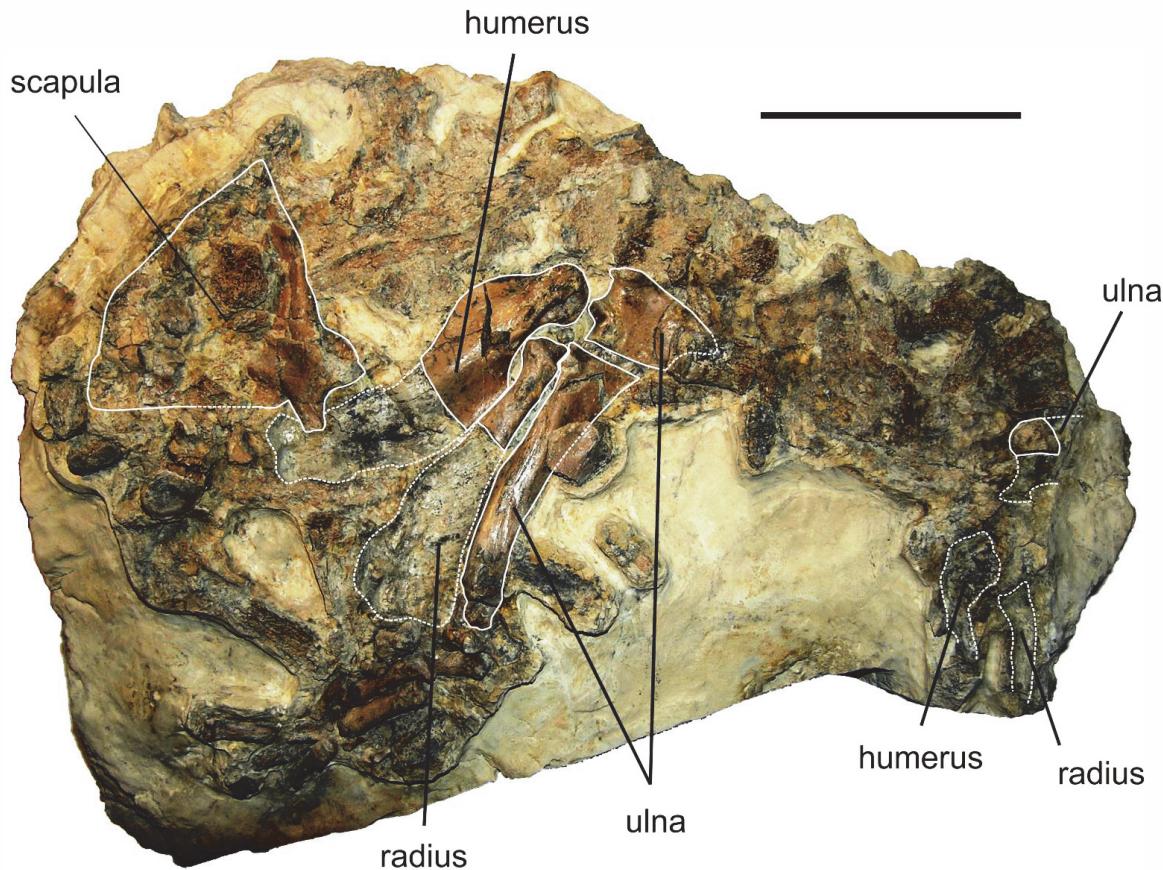


FIGURE 4. The partial skeleton of a true seal (*Phocinae* indet. 1) from the Volhyanian beds of Kolubaivtsi (Ukraine). Scale bar equals 10 cm.

bian) deposits of the Eastern Paratethys. It is larger than *Monachopsis pontica*, and similar in size to *Cryptophoca maeotica* (Koretsky, 2001). The humerus has a well-defined trochlear crest larger than that in most of other Paratethyan phocines. The deltoid crest is as long as that in *Sarmatonectes sintsovi* (Koretsky, 2001).

Phocinae indet. 2
Figure 5A

Material and locality. Scapula bearing the glenoid fossa, and a caudal vertebra; Stăuceni.

Description. A slender neck of the glenoid portion, a high glenoid fossa, and the slenderness of the scapula support the tentative identification of the fragmentary specimen as *Phocinae*. The glenoid fossa is deep and the glenoid tubercle is high, the angle between the anterior part and the scapular neck is smaller than in many other seals. The scapular neck is short, and the size of the fragment

is small, which may indicate that this specimen represents a young individual.

Comparison. The specimen resembles *Praepusa* sp. from the Sarmatian s.s. of Hungary and Austria (Koretsky, 2001, 2003) by the structure of glenoid fossa and the anterior edge of the scapula. The glenoid fossa in *Pachyphoca chapskii* is flatter and the glenoid tubercle is low. This specimen also differs from *Nanophoca* and *Monachopsis* by a greater angle between the scapular neck and the anterior edge of the bone.

CETACEA Brisson, 1762
ODONTOCETI Flower, 1867
Family KENTRIODONTIDAE Slijper, 1936
Genus KENTRIODON Kellogg, 1927
Kentriodon fuchsii (Brandt, 1873)
Figures 5B, 6

Material and locality. Partially preserved periotic bone, lumbar vertebra; Stăuceni, Basarabi, Stâncești.

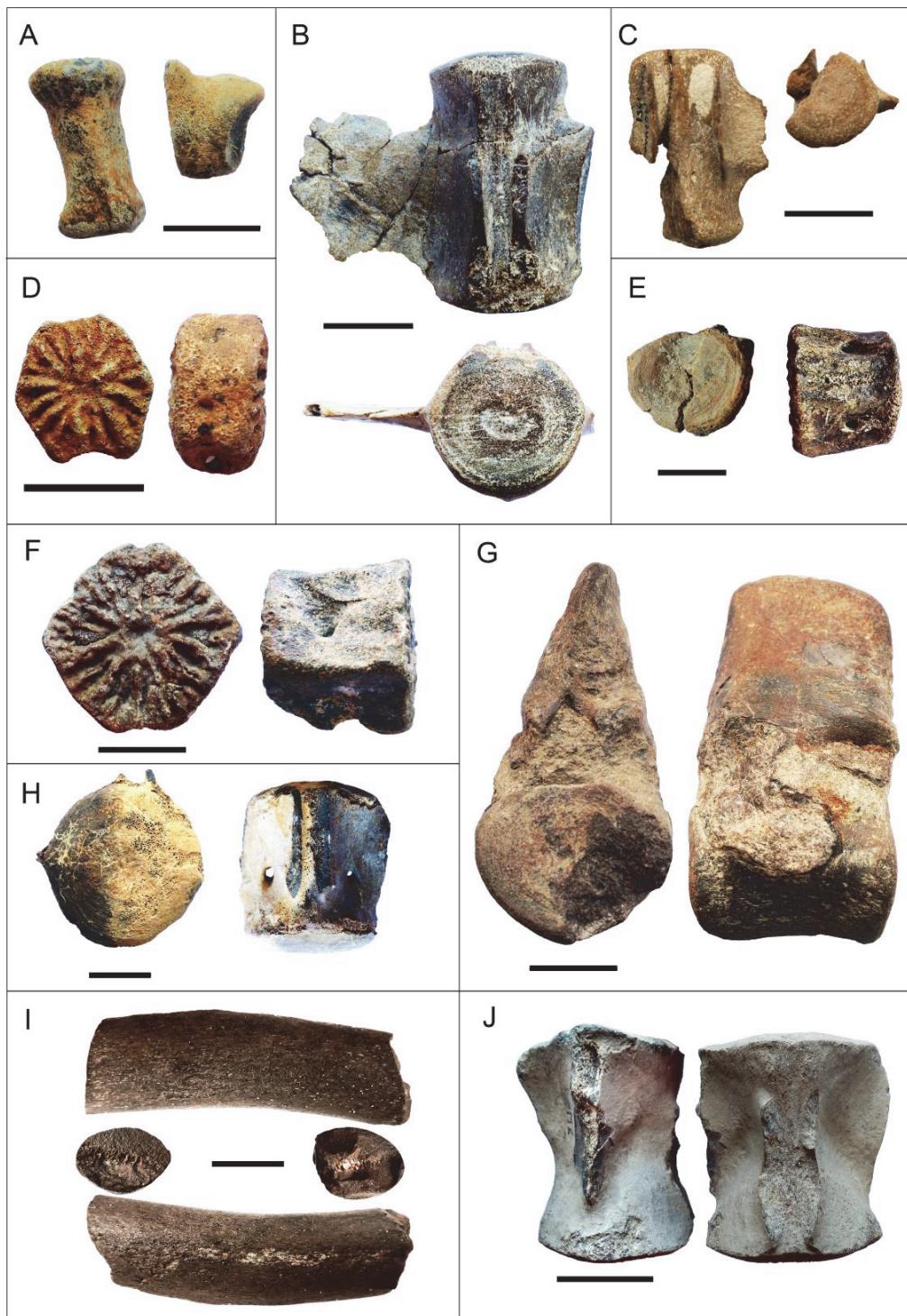


FIGURE 5. Marine mammals from the Volhynian beds of the Moldavian Platform: **A** – Phocinae indet. 2, scapula and caudal vertebra, Stăuceni; **B** – *Kentriodon fuchsii*, a lumbar vertebra, dorsal and posterior view, Basarabi; **C** – Kentriodontidae indet. 1 (cf. *Imerodelphis thabagarii*), lumbar vertebra, dorsal and anterior view, Saur-Mohyla; **D** – Kentriodontidae indet. 2, caudal vertebra, anterior and lateral view, Stăuceni; **E** – Kentriodontidae indet. 2, thoracic vertebra, anterior view, Stâncești; **F** – Kentriodontidae indet. 3, caudal vertebra, anterior and lateral view, Vilkhovets; **G** – *Pachyacanthus* sp., thoracic vertebra, anterior and lateral view, Vilkhovets; **H** – Cetotheriidae indet., caudal vertebra, dorsal and lateral view, Stăuceni; **I-J** – ? Mysticeti indet. ("*Archaeocetus fockii*"), rib fragment, lateral view and cross-section (**I**), caudal vertebra (**J**), dorsal and lateral view, Drăgușeni. Scale bars equal 2 cm in **A-I** and 5 cm in **J**.

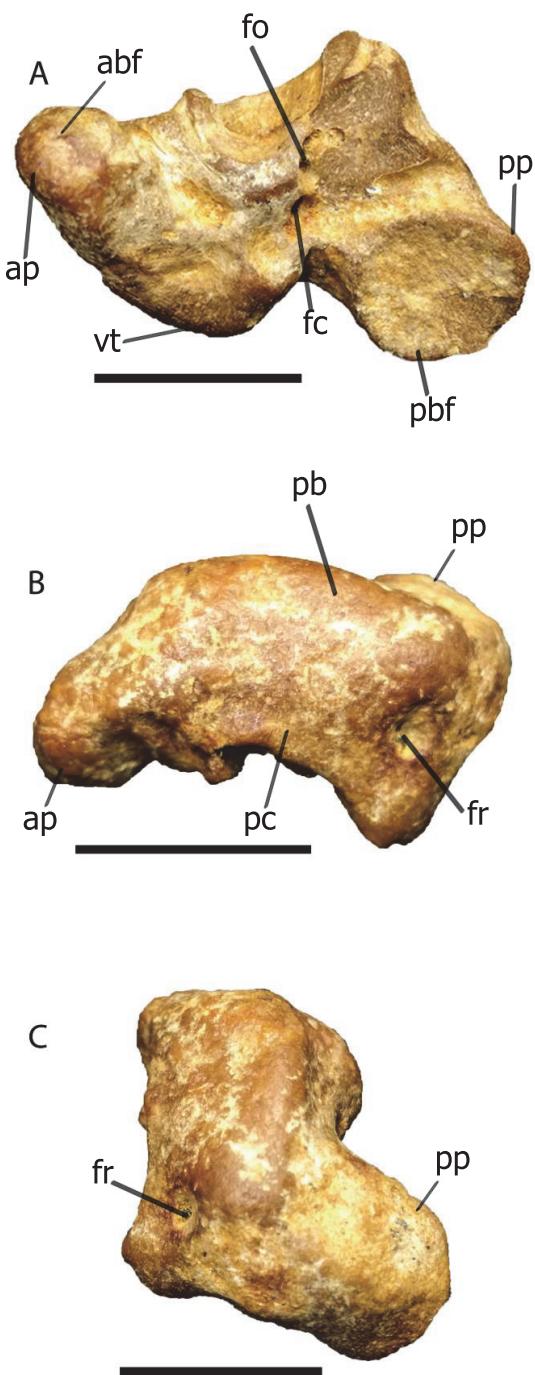


FIGURE 6. The periotic bone of *Kentriodon fuchsii* from the Volhynian of Stăuceni (Romania) in ventral (**A**), lateral (**B**), and posterior view (**C**). Abbreviations: abf, anterior bullar facet; ap, anterior process; fc, ventral foramen of the facial canal; fo, fenestra ovalis; fr, fenestra rounda; pbf, posterior bullar facet; pc, pars cochlearis; pb, periotic body; pp, posterior process; vt, ventrolateral tuberosity. Scale bars equal 2 cm.

Remarks and comparison. The periotic bone from Stăuceni (Figure 6) shares diagnostic features of *Kentriodon fuchsii* defined by Kazár (2006): medial bending of the anterior process and lateral bending of the posterior process; short posterior process with pentagonal posterior bullar facet; pars cochlearis (as preserved) broadly joined to the body of the periotic; a large ventrolateral tuberosity; a large eminence dorsal to fenestra rotunda. Also, a few more fragments can be tentatively referred to this taxon. They include elongated vertebrae, which were reported from Ghireni and Cordăreni by Macarovici and Zaharia (1967), Ionesi and Galan (1988), and newly found lumbar and caudal vertebrae, matching the original description of *Kentriodon fuchsii* (Brandt, 1873). The ulna reported by Codrea et al. (2014) from Basarabi also corresponds to the original description of *Kentriodon fuchsii* (Brandt, 1873). There are a few issues preventing clear attribution of vertebrae: size variation across the vertebrae; all the specimens are disarticulated and none of them is associated with the periotic bone; *Kentriodon fuchsii* itself needs redescription and at the end can be split into a few taxa (Kazár, 2006). Another dwarf kentriodontid of this morphotype but smaller than *K. fuchsii* is *Imerodelphis thabagarii*.

Kentriodontidae indet. 1
(cf. *Imerodelphis thabagarii* Mchedlidze, 1964)
Figure 5C

Material and locality. Three small and long lumbar vertebrae; Saur-Mohyla.

Comparison. A kentriodontid of this morphotype and size is *Imerodelphis thabagarii* from the Volhynian of Georgia (Mchedlidze, 1964). A vertebra identified earlier by Macarovici and Zaharia (1967) as *Phocaena euxinica* has similar size and proportions and, therefore, may represent this or a closely related taxon. However, the type material of *Phocaena euxinica* does not include vertebrae (Nordmann, 1860) and requires further revision.

? Kentriodontidae indet. 2
Figure 5D–E

Material and localities. Two thoracic and caudal vertebrae; Stăuceni, Stâncești.

Description and comparison. This is an unusually derived morphotype characterised by extremely short vertebral centra, mostly recorded in modern delphinids (Buchholtz, 2001). The dwarf size of the specimens resembles *Microphocaena podolica* from the Volhynian of Brykiv (Ukraine), which is known only from its cranium. Notably, *M. podolica* is the only previously known Volhynian

cetacean from the Fore-Carpathian Basin. Periotic bones from a few dwarf kentriodontids were also reported earlier from the Vienna Basin (Kazár, 2006).

? Kentriodontidae indet. 3

Figure 5F

Material and localities. Two relatively large and short caudal vertebrae; Stăuceni, Saur-Mohyla.

Description and comparison. Both specimens are characterised by a derived shape of centra, which are as short as high (Figure 5F), and, therefore, are similar to modern delphinids (Buchholtz, 2001). They resemble those of *Sophianaeetus commenticus* (Kazár, 2005) or "*Heterodelphis*" *leiodontus* (Papp, 1905), both from the Sarmatian s.s. of the Central Paratethys. Another similar vertebra from Mitoc-Prut was reported by Macarovici and Zaharia (1967).

Family PLATANISTIDAE Gray, 1846

Genus PACHYACANTHUS Brandt, 1871

Pachyacanthus sp.

Figure 5G

Material and locality. Thoracic vertebra and a partial rib; Vilkhovets.

Description. The vertebra shows lateral pachyosteosclerotic swelling of the spine distinct for this taxon, and the rib is strongly pachyosteosclerotic. This is the first record for *Pachyacanthus* in the Fore-Carpathian Basin: before it was only recorded in the Vienna Basin and adjoining areas of the Central Paratethys and at the eastern edge of the Eastern Paratethys (Mangyshlak Peninsula, Kazakhstan) (Dombrovsky, 1927; Kazár, 2010).

MYSTICETI Flower, 1864

Family CETOTHERIIDAE Brandt, 1872

Cetotheriidae indet.

Figure 5H

Material and localities. Caudal vertebra; Stăuceni, Kolubaivtsi.

Description. The new material from Stăuceni is represented by a caudal vertebra, which is assigned to Cetotheriidae due to its slight pachyosteosclerosis, a specific trait, which is shared only by that lineage during the late Middle Miocene. In dorsal view, the conical short vertebra has two lateral vascular foramina. The vertebra is similar in size to those of *Cetotherium riabinini*, a 3 m long whale (Gol'din et al., 2014). The specimen from Kolubaivtsi is a rib fragment, 3 cm wide, with developed pachyostosis and strong osteosclerosis, a distinct feature of the Cetotheriinae from the Eastern Paratethys (Gol'din and Startsev, 2017), also

shared by *Otradnocetus*, a stem member of Cetotherioidea (Gol'din, 2018).

Comparison. There is a vertebra reported by Macarovici (1947) from Săveni-Ghireni, and two more vertebrae reported by Macarovici and Zaharia (1967) from Mitoc-Prut, identified by them as *Cetotherium priscum*: all of the specimens share pachyosteosclerosis and small size, both typical for Cetotheriinae. Also, a similar vertebra was reported by Codrea (2006) from the Volhynian deposits of Țaga (Romania) in the Central Paratethys. A small, short pachyosteoclerotic lumbar vertebra of Cetotheriidae indet. also comes from Kaisersteibruch (Austria), a Badenian-Sarmatian locality in the Central Paratethys (Kazár, 2006). A small baleen whale from Khonkvitsi reported by Hofstein (1965) (temporarily unavailable for study) could belong to this family based on its extremely small size, short and high vertebrae: the centrum of the longest vertebra, the first caudal one, is 70 mm long, 70 mm wide and 70 mm high. Also, it is different from *Ciuciulea davidi* (Gol'din, 2018) and ? *Parietobalaena* sp. from the Central Paratethys (Gol'din and Radović, 2018) in having short lumbar vertebrae, a derived trait observed in Late Miocene cetotheriids. Meanwhile, its vertebrae are not pachyosteosclerotic, unlike those of the Paratethyan cetotheriines. Thus, it could represent a hitherto unknown dwarf mysticete taxon.

? Mysticeti indet.

Figure 5I, J

Material and locality. An osteosclerotic rib fragment without signs of pachyostosis; Drăgușeni.

Remarks. Long caudal vertebrae reported as "*Archaeocetus fockii*" by Macarovici (1947) from Khotin, Macarovici and Zaharia (1967) and by Macarovici and Zaharia (1968) from Hănești and Ghireni belong to this taxon. Similar specimens were described from the Sarmatian s.s. of the Central Paratethys. None of the postcranial elements can be firmly associated with cranial specimens and, therefore, the specimens cannot be assigned to any particular taxonomic unit. Besides, this type of vertebral anatomy is plesiomorphic for cetaceans and does not have any taxonomic value. Among Sarmatian cetaceans, a similar anatomy was recorded for "*Mesocetus*" *agrami* (Van Beneden, 1884).

DISCUSSION

The new records of the Volhynian fish fauna from Romania and Ukraine include five taxa representing four families (*Sarmatella doljeana*, *Sparus*

brusinai, *Bothus parvulus*, Clupeinae indet., and Scombroidei indet.). *Sarmatella doljeana* was a common member of the Sarmatian s.s. fauna of Croatia (Gorjanović-Kramberger, 1884), Serbia (Andelković, 1962, 1989), and Turkey (Rückert-Ülkümen, 1960), and it was recorded from Romanian localities earlier (Baykina, 2013a). *Sparus brusinai* was previously reported from the Sarmatian s.s. deposits of Croatia, Serbia, Turkey, Moldova, and Azerbaijan (Bannikov, 2010). Tarashchuk (1957) described it from the Badenian of western Ukraine as *Pagellus erythrinus* (revised by Bannikov, 2010; see Šoster and Kovalchuk, 2016). The flatfish *Bothus parvulus* was also recorded from the Sarmatian s.s. of Croatia (Schwarzans et al., 2017d). Therefore, all the fish taxa are shared with the Central Paratethys, and most of them are also shared with other sub-areas of the Eastern Paratethys where they are represented by more recent records (Kovalchuk et al., 2019).

The Volhynian marine mammals of the region are represented by cetaceans (at least five toothed and two baleen whales) and true seals. All of the cetaceans had small body size similar to those reported from the region (Gol'din et al., 2014; Gol'din and Radović, 2018). The cetacean assemblage from the Fore-Carpathian Basin is similar to that from the Sarmatian s.s. of the Central Paratethys (Codrea, 2006; Grigorescu and Kazár, 2006; Kazár, 2006; Gol'din and Radović, 2018), sharing its characteristic taxa such as *Kentriodon fuchsii* and *Pachyacanthus* sp., although less diverse, and is different from Volhynian faunas of the other parts of the Eastern Paratethys. In addition to cetaceans, the assemblage described here also includes seals, which represent a remarkable element of the fauna. True seals (Phocidae) are known from the Early Miocene of the Mediterranean region and from the Middle Miocene of the North Atlantic (Koretsky, 2001; Dewaele et al., 2018a, 2018b). In the Central Paratethys, seal remains were found both in Badenian (*Devinophoca*; Koretsky and Holec, 2002) and Sarmatian s.s. deposits (*Praepusa*; Koretsky, 2003). However, in the Eastern Paratethys, numerous seals were recorded in the later, Bessarabian, deposits but not in the earlier Miocene strata (Koretsky, 2001). In total, the Volhynian marine mammalian assemblage is similar to that from the subsequent Bessarabian stage of the Eastern Paratethys, but it is less diverse.

Grigorescu and Kazár (2006) suggested that the Central Paratethys was recolonised by the Sarmatian mammalian fauna from the Eastern Para-

tethys after the Badenian/Sarmatian extinction event. As seen from this study, the earliest Sarmatian fauna of the Eastern Paratethys was less diverse and less abundant than the Central Paratethyan fauna of the same age, and in the early Volhynian it was mostly concentrated in the Fore-Carpathian Basin, the region connected both with the Central and Eastern Paratethys (Kováč et al., 2017). Only the easternmost localities in the Central Paratethys yielded a different marine vertebrate fauna, e.g., in the Caucasus (Bannikov, 2009), or in younger deposits (Mchedlidze, 1964). Also, the Sarmatian fauna studied here includes taxa absent from the earlier Middle Miocene deposits of the Eastern Paratethys but present in the Mediterranean or Atlantic, e.g. *Kentriodon* or *Phocinae*.

Sparus brusinai and Cetotheriidae whales, on the contrary, are known from the Badenian of the Fore-Carpathian Basin (Bannikov, 2010; Gol'din, 2018). Meanwhile, the younger, Bessarabian fauna of the whole Eastern Paratethys is abundant, diverse, and similar to the earlier fauna of the Central Paratethys and Fore-Carpathian Basin. Importantly, the faunal assemblage described here is typical for a marine environment. Most of the fishes studied here (*Sarmatella*, scombrids, *Sparus*) are euryhaline, although they require normal marine conditions for spawning (Bannikov, 2009; Baykina, 2013). *Bothus* is a stenohaline genus intolerant of low water salinity (Carnevale et al., 2006). Therefore, most of the Volhynian fauna can be only of marine origin and came from an area with stable marine conditions. Thus, some elements of the Volhynian marine vertebrate fauna found in the Eastern Paratethys had likely originated in the Badenian of the Central Paratethys (Figure 7) (Palcu et al., 2015, 2017). Some of its representatives (e.g., Cetotheriidae) could survive the Badenian/Sarmatian extinction event in the Fore-Carpathian Basin and subsequently disperse to the Sarmatian seas of both the Central and Eastern Paratethys. The Volhynian vertebrate fauna can be regarded as transitional between the Sarmatian s.s. fauna of the Central Paratethys and the subsequent Bessarabian fauna of the Eastern Paratethys. However, this was not the sole trend in the development of the Volhynian fauna. For example, some elements of the endemic fish fauna are recorded in the Eastern Paratethys from the Konkian and then the endemism accelerated in the Volhynian (Carnevale et al., 2006; Bannikov, 2010; Schwarzans et al., 2017a-d).

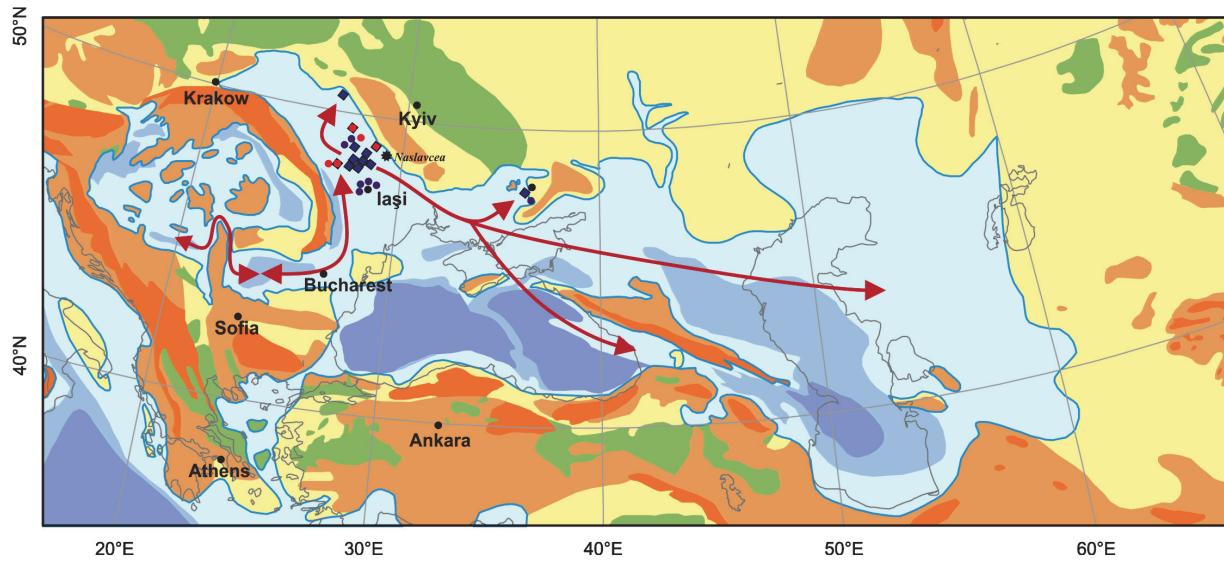


FIGURE 7. Suggested scheme of marine vertebrate fauna dispersal in the Eastern Paratethys during the Volhynian age (modified after Schneider et al., 2013).

Both marine fish and mammal taxa show the tendency toward a decrease in body size. In particular, marine mammals are represented only by small-sized taxa including a dwarf toothed whale. Notably, *Microphocaena podolica* (Kudrin and Tatařinov, 1966) from the Volhynian of the Eastern Paratethys (Brykiv) was among the smallest cetaceans in history, ca. 80–90 cm in total length (Gol'din, 2016). In addition, unusual reduction of the number of vertebrae in some fishes, mentioned by Bogachev (1955) and Carnevale et al. (2006), is also seen in our material. This dwarfism is shared by Sarmatian s.s. marine mammals of the Central Paratethys (Gol'din and Radović, 2018). Diminishing body size of top predators is a well-known effect in isolated and marginal seas (Perrin, 1984; Shackell et al., 2010). This is an additional argument for the Central Paratethyan origin of some members of the Volhynian Eastern Paratethyan fauna. During the Late Miocene, a restricted marine environment prevailed in the Eastern Paratethys (Bratishko et al., 2017), while the Central Paratethys developed into the brackish Pannonian Sea (Popov et al., 2004). The connection between different parts of the Paratethyan realm was interrupted at the end of the Middle Miocene, leading to the development of the endemic mammalian megafauna in the Eastern Paratethys.

CONCLUSION

The Volhynian vertebrate faunal assemblage of the western and northern edges of the Eastern

Paratethys, including the adjoining Fore-Carpathian Basin, is a marine fauna with strong affinities to the synchronous earliest Sarmatian fauna of the Central Paratethys. It shares characteristic Central Paratethyan taxa, although it is less diverse and less abundant. Therefore, the Sarmatian marine fauna of the Eastern Paratethys at least partly originated from the Mediterranean and Paratethyan faunas during Badenian. Marine taxa initially concentrated in the Fore-Carpathian Basin, from where they dispersed eastwards into the Eastern Paratethys during Volhynian and evolved under increasingly isolated conditions. Some of them (e.g. cetotheriid whales) could survive the Badenian/Sarmatian extinction event in the Fore-Carpathian Basin and disperse both westwards and eastwards from that region.

ACKNOWLEDGEMENTS

All authors have contributed equally. We are thankful to B. Ridush (CNU), M. Bukhsianidze (GNM) and M. Gasparik (HNHM) for their kind help with allowing an access to the collections. L. Godlevska, M. Savchenko and P. Vorobey are acknowledged for help in the field work by PG. The authors thank W. Schwarzhans, D. Palcu, Ş. Vasile, B. Studencka, and an anonymous reviewer for their critical comments to the early draft and suggestions that significantly improved the manuscript.

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SUPPLEMENTARY MATERIAL

Tables available as PDF for download.

TABLE S1. Age summary of the studied localities.

TABLE S2. Marine vertebrate records from the Volhynian of the Eastern Paratethys (western and northern part) by taxon, locality and age.

TABLE S3. List of specimens used in this study.