



The oldest Asian *Sivaonyx* (Lutrinae, Mustelidae): a contribution to the evolutionary history of bunodont otters

Camille Grohé, Louis de Bonis, Yaowalak Chaimanee, Cécile Blondel,
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ABSTRACT

We describe here specimens of the otter *Sivaonyx* cf. *gandakasensis* from the late Middle Miocene hominoid-bearing basin of Chiang Muan, northern Thailand. Asian *Sivaonyx* species were previously recorded in the Late Miocene of the Siwalik Hills and Southern China. The new Thai remains could constitute the oldest occurrence of the genus in Asia (12.4–12.2 m.y.) and, more generally, in the Old World. *S. gandakasensis*, also occurring in the early Late Miocene and possibly Middle Miocene deposits of the Siwalik Hills, is the most primitive species of the genus. Based on the m1 morphology, we suggest a close relationship of this species with *S. hessicus* from the Late Miocene of Europe. In this study, we discuss the primitive morphological characters of Eurasian *Sivaonyx* compared to African ones, using notably the previously unknown M1 of *S. cf. gandakasensis* (e.g., developed M1 parastyle, high protoconid and narrow talonid basin on m1). Moreover, *S. bathygnathus*, the third and most derived Eurasian species known from the Late Miocene of the Siwaliks and Southern China, seems to be the most probable ancestor of Mio-Pleistocene African *Sivaonyx*. We therefore propose a dispersal of *Sivaonyx* from Asia to Africa during the Late Miocene, probably between ca. 10 and 7 m.y.

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INTRODUCTION

The Lutrinae belongs to the Mustelidae, the most species-rich family of Carnivora. They are currently represented by 7 genera and 13 species occurring both in freshwater and coastal marine habitats throughout much of the continents (except Australia and Antarctica). Although otters exhibit a taxonomically diversified and worldwide fossil record since the Miocene (Willemse, 1992; Pickford, 2007; Werdelin and Peigné, 2010), their phylogenetical relationships remain up to now controversial when based on morphological data (e.g., de Muizon, 1982; van Zyll de Jong 1987, 1991; Willemse, 1992), and in conflict with the results suggested by molecular analyses (Koepfli and Wayne, 1998; Koepfli et al., 2008). For example, in these last analyses, the basal position of the giant otter *Pteronura* within the subfamily and the diphyle between the river otters *Lutra* and *Lontra* notably contribute to split the tribes used by morphologists. Morphological classifications of otters into different tribes are mainly based on dental characters, because of a fragmentary fossil record suffering from the lack of cranial and postcranial specimens. These classifications are therefore closely related to diet preferences. Bunodont otters (tribe Enhydrini after Willemse, 1992 and tribes Enhydrini+Enhydriodontini after Pickford, 2007) constitute a diverse group of fossil and extant forms for which dentition seems to be adapted to hard item consumption (e.g., extant sea otter *Enhydra*). Some important generic controversy exists for this group and is covered later in this paper (see discussion). Fossil bunodont otters successfully occupied continental ecosystems of Eurasia, Africa and North America from the Miocene to the Pleistocene. One of the most charismatic and abundant genus of this group is *Sivaonyx*, known to include lion-sized gigantic predators. *Sivaonyx* is recorded in the Mio-Pliocene and possibly Pleistocene of Africa (Kenya, Ethiopia, Uganda, Chad, South Africa, Egypt), and in the Miocene of Asia (Siwalik Hills of Pakistan and India, Southern

China) and Europe (Germany, Turkey). In this study we describe new material of *Sivaonyx* recovered from the late Middle Miocene Ban Sa locality of the Chiang Muan Basin, northwestern Thailand. This material could represent the oldest occurrence of the genus in Asia and brings a new opportunity to discuss the primitive morphological characters of Asian *Sivaonyx* compared to African ones and to propose possible Miocene dispersal events between Southern Asia and Africa. This study aims to improve our knowledge on both the evolutionary history of otters and the faunal dispersals occurring between Africa and Asia during the Miocene.

Geological Setting

The Middle Miocene Ban Sa locality is located in the Chiang Muan Basin, one of the numerous intramontane Miocene fossiliferous basins of northern Thailand. The sediments of the Chiang Muan Basin are mainly composed of mudstones and sandstones intercalated with palaeosols and two main lignite beds (i.e., Lower and Upper Lignite Members), both providing fossil remains (see Benammi et al., 2004 and Coster et al., 2010 for lithological details and stratigraphic section of Ban Sa locality). The Chiang Muan coal mine is famous for exposing hominoid-bearing fossiliferous levels, yielding dental remains of the large-bodied genus *Khoratpithecus*, closely related to extant orangutans (Chaimanee et al. 2003, 2004). Hominoid associated mammal fauna includes the gomphotheriid proboscidean *Tetralophodon* cf. *xialongtanensis*, the suoids *Paracheleuastochoerus sinensis*, *Hippopotamodon* cf. *hyotherioides*, *Conohyus sindensis*, *Pecarichoerus sminthos*, the rhinocerotid *Chilotherium* (*Subchilotherium*) *intermedium*, the ruminants *Dorcatherium*, *Eotragus*, and *Siamotragulus*, and the castorid *Steneofiber siamensis* (Chaimanee et al., 2003; Pickford et al., 2004; Suganuma et al., 2006; Chavasseau et al., 2009; Suraprasit et al., 2011). Biochronological data suggested an age comprised between ca. 12 and 10

m.y. for the Chiang Muan sequence based on the faunal comparison with Xiaolongtan locality, Yunnan Province of China (Pickford and Liu, 2001; Pickford et al., 2004). More recent paleomagnetical studies (Benammi et al., 2004; Coster et al., 2010) and correlations established with Mae Moh coal mine sediments in Thailand rather indicate an age of 12.4-12.2 m.y. for the Upper Lignite Member. In 2003, a Thai-French paleontological expedition found the first fossils of Carnivora from the Upper Lignite Member of the Ban Sa locality. This material constitutes the focus of our study.

MATERIAL AND METHODS

The material is composed of two isolated and well-preserved molars stored at the Department of Mineral Resources of Bangkok, Thailand (**DMR**). The proportions and morphology of the teeth as well as their stratigraphic location permit to attribute them to the same species. Abbreviations for other institutions and specimens cited in the text and tables are: **AM**, American Museum of Natural History, New York, NY; **BAR**, fossils from Baringo District, Community Museums of Kenya; **CM**, Chiang Muan fossils, Thailand; **BMNH**, Natural History Museum, London, England; **BSPG**, Bayerische Staatssammlung für Paläontologie und historische Geologie, Munich, Germany; **GSI**, Geological Survey of India, Calcutta; **GSP-Y**, Geological Survey of Pakistan, Quetta – Yale University, New Haven, CO; **IVPP V**, Institute of Vertebrate Paleontology and Paleoanthropology, Beijing, Vertebrate collection; **KNM**, Kenya National Museums, Nairobi: -**KP** Kanapoi; -**LU** Lukeino; **NK**, fossils from Nkondo, Uganda National Museum, Kampala; **SAM PQ-L**, Iziko South African Museum, Cape Town, collections from Langebaanweg; **TM**, fossils from Toros-Menalla, Chad; **YPM**, Yale Peabody Museum, New Haven, CO; **YV**, Yunnan Vertebrate, China.

Measurements were taken with a digital caliper to the nearest 0.1 mm. Abbreviations for description and measurements are: **L**, maximum mesiodistal length; **Llab** and **Lling**, labial and lingual maximum mesiodistal length of M1; **m**, lower molar; **M**, upper molar; **MlabDling**, maximum distance between mesiolabial corner and distolingual corner of M1; **MlingDlab**, maximum distance between mesiolingual corner and distolabial corner of M1; **W**, maximum linguolabial width; **Wdist**, linguolabial width of M1 at the distal border.

SYSTEMATIC PALAEONTOLOGY

Order CARNIVORA Bowdich, 1821
Suborder CANIFORMIA Kretzoi, 1943
Infraorder ARCTOIDEA Flower, 1869
Family MUSTELIDAE Fischer, 1817
Subfamily LUTRINAE Bonaparte, 1838
Genus *SIVAONYX* Pilgrim, 1931

Diagnosis (Pilgrim, 1932; pp.88). “Lutrinae of large size; having a P4 of an approximately quadrate shape with a moderately trenchant paracone-metacone blade; paracone somewhat stoutly built; metacone very slightly elongate; parastyle well-developed but low; with two internal cusps, protocone and hypocone, lower than the outer cusps, somewhat stoutened and expanded so that the basin-shaped area is contracted; with a strong internal cingulum; mandible with a ramus deeper than the length of m1; canine large; p1 single-rooted; p4 broad posteriorly, with a strong, outwardly situated posterior cusp and a broad cingulum, no anterior accessory cusp; m1 very broad; paraconid situated so as to be bisected by the median axis of the tooth; metaconid as high as paraconid; talonid broader than trigonid and approximately equal in length to it, basin-shaped, surrounded by a ring of cusps of which the entoconid is as well developed as the hypoconid; m2 small, oval.”

Comments. Because a revision of all *Sivaonyx* species is beyond the scope of this paper and because some generic attributions are still debated (see discussion), we do not emend the diagnosis of *Sivaonyx* until more material could allow it. However, we revise some morphological features for the m1 and mention other ones for the M1 in light of the recent specimens assigned to the genus. The trigonid of *Sivaonyx* is low with approximately equal cusps, the protoconid is however higher than the paraconid and metaconid in the most primitive species (see description and comparisons). A ring of cusplets surrounding the talonid, and a similarly developed hypoconid-entoconid are not observed in all species (e.g., hypoconid higher than the entoconid in *S. gandakasensis*, entoconid higher than the hypoconid in *S. hendeysi*). As underlined by Peigné et al. (2008) and not mentioned in the original diagnosis, the labial cingulid is marked on m1, and the M1 shows a strong metaconule and a distinct notch between the paraconule and the protocone.

Type species. *Sivaonyx bathygnathus* (Lydekker, 1884)

Other included species. *S. africanus* (Stromer, 1931), *S. beyi* Peigné et al., 2008, *S. ekecaman* (Werdelin, 2003), *S. gandakasensis* Pickford,

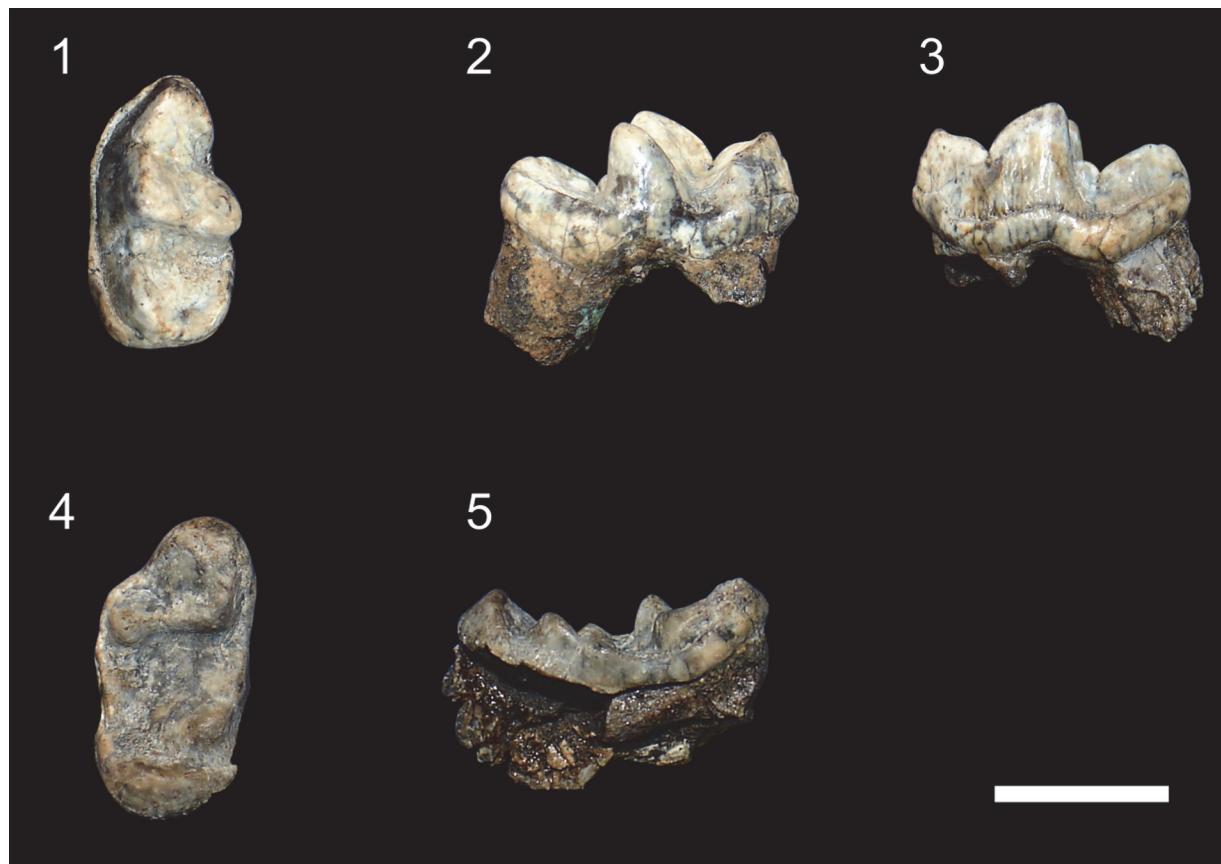


FIGURE 1. *Sivaonyx* cf. *gandakasensis* from the late Middle Miocene of Chiang Muan, Thailand: CM-03-A, left m1 in occlusal (1), lingual (2) and labial (3) views; CM-03-B, right M1 in occlusal (4) and mesial (5) views. Scale = 1cm.

2007, *S. hendeyi* (Morales et al., 2005), *S. hessicus* (Lydekker, 1890), *S. kamuhangirei* Morales and Pickford, 2005, *S. soriae* Morales and Pickford, 2005

Distribution. Late Miocene-Early Pleistocene of Africa (Egypt, Chad, Ethiopia, Kenya, Uganda, South Africa), Middle-Late Miocene of Asia (India, Pakistan, Thailand, China), Late Miocene of Europe (Germany, Turkey).

Sivaonyx cf. *gandakasensis* Pickford, 2007

Figure 1; Tables 1 and 2

Referred specimens. CM-03-A, left m1; CM-03-B, right M1 from Chiang Muan Basin, Thailand

Holotype locality. Locality Y182, Gandakas, Potwar Plateau of northern Pakistan – early Late Miocene (9.2 m.y. after Barry et al., 2002, same age as locality Y317 after Barry et al., 1982)

Distribution. Upper Lignite Member of the Chiang Muan Basin, northern Thailand – late Middle Miocene (12.4-12.2 m.y. after Benammi et al., 2004 and Coster et al., 2010)

Description. The left m1 shows a low trigonid as wide as the talonid and longer than it. The protoconid is higher than the paraconid and metaconid. The metaconid is slightly higher and is more lingual than the paraconid. The protoconid and the paraconid are separated by a well-marked carnassial notch. The paraconid has a nearly vertical mesial face and its distolingual crest does not extend more than the distolabial one and does not reach the lingual cingulid. Distal to this crest, a deep V-shaped valley separates the paraconid from the metaconid, contributing to a lingually opened trigonid. The metaconid is displaced distally with respect to the protoconid and is curved backwardly in lingual view. A small cuspule is present at the distal base of the protoconid. The talonid exhibits a tall hypoconid crest but lower than the paraconid, a weaker and curved entoconid crest, and a narrow and shallow talonid basin. The lingual border of the talonid appears oblique: it is lower and thicker near the metaconid relative to the distal end of the tooth. This height difference creates an acute angle between the distal face of the metaconid and the

TABLE 1. Lower dental measurements of *Sivaonyx* species (in mm). Measurements are from first hand and from ¹Pilgrim 1932; ²Colbert 1935; ³Qi 1983; ⁴Zong 1997; ⁵Werdelin 2003; ⁶Morales et al. 2005; ⁷Morales and Pickford 2005; ⁸Pickford 2007; ⁹Peigné et al. 2008.

	L	W
<i>Sivaonyx cf. gandakasensis</i>		
CM-03-A	13.7	7.6
YPM 20206	15.4	8.5
<i>S. gandakasensis</i>		
GSP-Y 4225 (holotype) ⁸	15	7.9
GSP-Y 11703 ⁸	13.9	7.9
GSI D 249 ⁸	14.4	8.6
BMNH M 13175	15.4	9.1
?<i>S. gandakasensis</i>		
BMNH G 4	12.8	7.6
<i>S. bathygnathus</i>		
GSI D 33 (holotype) ⁸	17.1	9.7
GSI D 156 ¹	16	9.4
GSI D 250 ⁸	16.4	9.1
AM 19509 ²	16.5	9.5
IVPP V 6886.12 ³	15.9	9
<i>S. cf. bathygnathus</i>		
YV 2517 ⁴	14	8
<i>S. hessicus</i>		
BMNH 27486 (holotype)	15.9	8.6
<i>S. beyi</i>		
TM 171-01-033 (holotype) ⁹	20,3	
TM 172-05-001 ⁹	22,8	13,4
TM 355-02-002 ⁹	20	11,6
TM 247-01-005 ⁹	21,5	12,7
<i>S. ekecaman</i>		
KNM-KP 10034 (holotype) ⁵	21,2	13,5
BAR 720'03 ⁷		12,8
BAR 567'05 ⁷	20,1	13
<i>S. soriae</i>		
KNM-LU 337 & 338 (holotype) ⁶	17,6	10,5
BAR 1984'05 ⁷	17,5	10,6
<i>S. hendeyi</i>		
SAM PQ-L 5000 (holotype) ⁹	21,3	13,1
<i>S. kamuhangirei</i>		
Unnumbered (holotype) ⁹	26	15,9
NK 1988'89 ⁷		14,4
<i>S. africanus</i>		
BSPG 1930 XI 1 (holotype) ⁹	22,2	

TABLE 2. Upper dental measurements of *Sivaonyx* species (in mm). Measurements are from first hand and from ¹Qi 1983; ²Peigné et al. 2008.

	Llab	Lling	Wdist	MlingDlab	MlabDling
<i>Sivaonyx cf. gandakasensis</i>					
CM-03-B	7.6	7.1	14	10.7	14.2
<i>S. bathygnathus</i>					
IVPP V 6886.3 ¹	9.3		14		
IVPP V 6886.5 ¹	9.5		14.6		
<i>S. beyi</i>					
TM 90-00-0662	11.7	11.7	17.3	16.5	20
<i>S. soriae</i>					
BAR 1082'012	11	12.2		15.6	18.2
<i>S. ekecaman</i>					
KNM-KP 10034 (holotype) ²	12.1	15.8	19		
<i>S. africanus</i>					
BSPG 1930 XI 1 (holotype) ²	13	14.1	18.8	17.1	

lingual talonid rim. The cingulid is thick on the labial side of the carnassial. It extends until the lingual base of the paraconid where it appears thinner. A shearing facet is visible on the upper part of the labial faces of the paraconid and protoconid.

The right M1 is broken at its most mesiolingual part so that the cingulum is interrupted at the paraconule-protocone level. The tooth is transversely widened. The mesial and distal borders are roughly parallel. The metacone is higher and mesiodistally shorter than the paracone. The paracone is more crestiform than the metacone. Both cusps are separated by a shallow notch. The labial shelf of the paracone is much wider than the metacone labial shelf. The parastyle is very long and extends from the paracone to join a wide cingulum surrounding the mesiolabial corner of the tooth. The paraconule and the protocone are crestiform, separated by a small notch, and constitute a blade along the mesial border of the tooth that bends distally at the protocone level. The protocone is as high as the paracone. The paraconule is located at a similar level as the metaconule but the latter is more prominent. Distally, the metaconule is not connected to the metacone. The labial crest descending from the metaconule joins the distal cingulum at the distolingual base of the metacone. A crescent-shaped cingulum is present at the distolingual corner of the tooth. It reaches approximately the height of the metacone. The thickest part of this cingulum is displaced mesially relative to the metacone. Both m1 and M1 show finely wrinkled enamel.

COMPARISONS

The wide talonid, the presence of a cuspule at the distal base of the protoconid, the high metacoenid of m1, and its abrupt connection with the lingual talonid rim recalls the carnassial morphology of otters. Within Miocene otters, the morphology of the Chiang Muan material is closer to that of *Sivaonyx* than to that of any other genera (e.g., low trigonid cusps, high hypoconid, moderately deep talonid basin and rounded talonid rim on m1, presence of a metaconule on M1). Until now, nine species of *Sivaonyx* have been recorded from the Miocene of Eurasia and from the Mio-Pliocene of Africa. Several specimens of *Sivaonyx* sp. are also reported in the Plio-Pleistocene of Chad, Ethiopia and Kenya but are undescribed yet (Werdelin and Peigné, 2010). Here we compare the material from Chiang Muan to that of Eurasian and then African species of *Sivaonyx*. Dental measurements of m1 and M1 of *Sivaonyx* species are presented in Tables 1 and 2. Their stratigraphic and geographic distributions are illustrated in Figure 2.

The type species *S. bathygnathus* has long been the only *Sivaonyx* species recorded in Asia. Originally described from the Late Miocene of the Siwalik Hills of Pakistan (Hasnot, Potwar Plateau; Lydekker, 1884; Pilgrim, 1932), the species is also reported in the Late Miocene of Haritalyangar, India (Prasad, 1968), of Lufeng and possibly Yuanmou, Yunnan Province of China (Qi, 1979, 1983; Zong, 1997; Qi et al., 2006). The holotype of *Sivaonyx bathygnathus* (GSI D 33) is a mandible fragment with a poorly preserved m1, but the additional

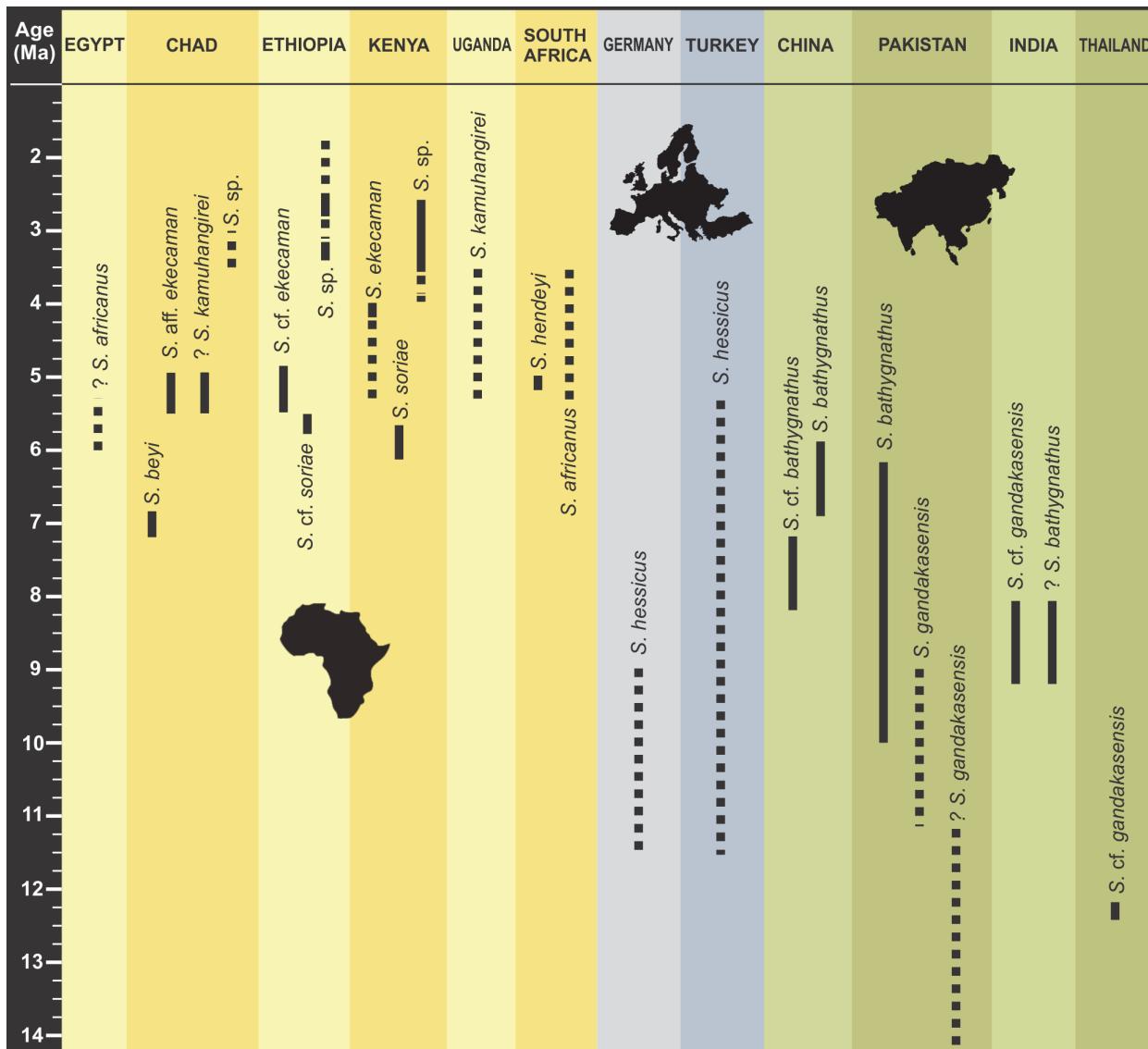


FIGURE 2. Stratigraphic and geographic distribution of *Sivaonyx* species in Africa, Europe and Asia. We used Barry et al. 2002, Qi et al. 2006, Pickford 2007, Coster et al. 2010 and Patnaik et al. 2013 for the datation of Asian localities and Werdelin and Sanders 2010 and references therein for African ones. Dotted lines indicate imprecise datations.

material described by Pilgrim (1932) and Colbert (1935) permits to compare the morphology of the lower carnassial with our material: CM-03-A differs from GSI D 156 and AM 19509 by smaller proportions, a narrower and longer paraconid, a higher protoconid compared to the paraconid instead of a similar height of these cusps and a narrower talonid basin. Moreover, CM-03-A differs from the m1 of *Sivaonyx* cf. *bathygnathus* from Yuanmou (YV 2517, figured by Zong, 1997) in being smaller and exhibiting a narrower talonid basin. The only M1 described as *S. bathygnathus* comes from Lufeng (IVPP V 6886.5, figured by Qi, 1983). A very similar M1 coming from Pakistan (loc.D013 dated to

6.3 m.y.) and stored in the Harvard Peabody Museum collections (GSP-Y16790) has been directly observed and thus allows more comparisons. CM-03-B differs from these specimens in being narrower, and in having a more developed paracone labial shelf, a less prominent metaconule, a less trenchant paraconule-protocone blade and a less extended distolingual shelf.

The more recent revision of bunodont otters from the Siwaliks provided by Pickford (2007) lead to the erection of a new species, *S. gandakasensis*, for the oldest *Sivaonyx* forms of Pakistan. CM-03-A is very close to the m1 of *S. gandakasensis* by comparisons with morphological features of the

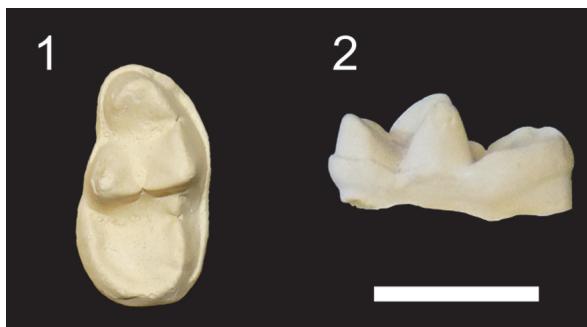


FIGURE 3. *Sivaonyx* cf. *gandakasensis* from the Late Miocene of Haritalyangar, India: cast of YPM 20206, right m1 in occlusal (1) and lingual (2) views. Scale = 1cm.

holotype GSP-Y 4225 and with the dimensions of the referred material. The cuspule at the distal base of the protoconid is less distinct than in CM-03-A, but this feature is certainly variable within *Sivaonyx* species, as observed for the Potwar Plateau material of the Harvard Peabody Museum collections referred to *S. bathygnathus*. Moreover, the specimens of *S. gandakasensis* exhibit generally worn cusps, which prevent us from checking whether the protoconid is higher than the paraconid and metaconid. An isolated m1 from Haritalyangar observed in the YPM collections (YPM 20206, cast illustrated here for the first time, Figure 3) shows striking morphological similarities with the holotype of *S. gandakasensis*. This Indian lower carnassial displays a higher protoconid compared to the metaconid, as in the Chiang Muan material. The closer morphology of YPM 20206 with the holotype of *S. gandakasensis* rather than with the specimens referred to the type species (e.g., shorter paraconid, higher protoconid, narrower talonid basin) and the need of taxonomic revision of the specimens published by Prasad in 1968, which some do not belong to otters (Morales and Pickford, 2011), sow doubts on the presence of *S. bathygnathus* in Haritalyangar. Finally, Pickford (2007) attributed a mandible (BMNH G 4) from the Middle Miocene of Chhorwali Katas, northern Pakistan, to *Vishnuonyx chinjiensis*. However, the holotype of this species is an upper carnassial and the only dental material bearing m1 (GSI D 245, figured by Pilgrim, 1932) shows a lower carnassial with a narrower talonid basin and a less rounded outline of the distal talonid compared to that of BMNH G 4. This specimen is compared here because its morphology rather recalls that of *Sivaonyx gandakasensis* than that of *Vishnuonyx chinjiensis*. In comparison with the lower carnassial of our Thai material, the m1 of BMNH G 4 is slightly

smaller and exhibits a less distinct cuspule at the base of the protoconid, a more distinct hypoconid, and a lingual cingulid extending until the metaconid base. Both Indian and Thai specimens compared in this paper can be referred to *S. cf. gandakasensis* until additional material from Haritalyangar and Chiang Muan are found.

The European species *S. hessicus* is represented by a mandible (BMNH M 27486) from the Late Miocene of Eppelsheim, in Germany (Lydekker, 1890) and listed from the Miocene of Kutchuk-Tchekmedjé locality, in Turkey (Nicolas, 1978). The holotype of *S. hessicus* shows an m1 very close morphologically to CM-03-A (e.g., protoconid higher than the paraconid, narrow talonid basin, distinct cuspule at the distal base of the protoconid). Only the trigonid lingual valley is narrower and the lingual cingulid extends under the metaconid, unlike in our Thai material. Moreover, the m1 has larger dimensions than CM-03-A. Originally attributed to *Lutra* (Lydekker, 1890), Pilgrim (1931) moved the holotype of the European species to *Sivaonyx*, considering the strong morphological similarities with *S. bathygnathus* from the Indian Subcontinent. Pickford (2007) agreed with Pilgrim's attribution and even suggested a synonymy with the Asian species. However, *S. hessicus* is clearly closer morphologically to *S. gandakasensis* than to *S. bathygnathus* (e.g., narrower paraconid, protoconid higher than the metaconid and paraconid, narrower talonid basin).

Sivaonyx beyi from the Late Miocene of Toros-Menalla, Chad, is the oldest representative of the genus in Africa (Peigné et al., 2008). The m1s (TM 247-01-005 and TM 355-02-002) differ from CM-03-A by a deeper valley between the paraconid and protoconid with a more oblique distolingual crest of the paraconid, a protoconid lower than the metaconid (TM 355-02-002) or approximately equal to it (TM 247-01-005), a less distinct cuspule at the distal base of the protoconid, a nearly horizontal lingual rim of the talonid, and a deeper talonid basin. The M1 (TM 90-00-066) differs from CM-03-B in being longer and wider, in having a poorly developed paracone labial shelf and a shorter parastyle, an oblique and a more trenchant paraconule-protocone blade, a crest descending mesiolingually from the metaconule. The highest part of the distolingual cingulum is also directly lingual to the metacone rather than mesial to it.

Sivaonyx soriae is recorded in the Late Miocene Lukeino Formation, Kenya (Morales and Pickford, 2005). The occurrence of this species in

the Lukeino beds has been originally based on lower dental specimens. Upper isolated teeth from the same locality initially attributed to *S. senutae* are now considered as synonym based on their suitable dental morphology and on the range of intraspecific size variation within *Sivaonyx* species (i.e., *S. hendeyi* and *S. ekecaman*; see Peigné et al., 2008 for details). *S. cf. soriae* is also reported by a fragment of P4 from the Late Miocene of the Asa Koma Member, Adu Asa Formation, in Ethiopia (Haile-Selassie, 2008). Moreover, *S. aff. ekecaman* occurs in the locality of Kossom Bougoudi, Chad, at the Mio-Pliocene boundary (Bonis et al. 2008). The m1s of *S. soriae* (KNM LU 337 and BAR 1984'05) differ from CM-03-A by a wider metaconid, resulting in a less opened trigonid and a more oblique protocristid, a metaconid higher than the protoconid, a wider talonid relative to the trigonid, and a deeper talonid basin. The M1 (BAR 1082'01) differs from CM-03-B in being wider and mostly longer, in having an oblique paraconule-protocone blade with a protocone closer to the lingual rim of the tooth, and a metaconule closer to the metacone.

Sivaonyx ekecaman has been recovered from the Early Pliocene of the Kanapoi Formation (Werdelin, 2003) and of the Magabet Formation, in Kenya (Morales and Pickford, 2005). *S. cf. ekecaman* is also reported from the latest Miocene and earliest Pliocene of the Kuseralee and Haradaso Members, Sagantole Formation, in the Middle Awash of the Afar rift, in Ethiopia (Haile-Selassie, 2008). Moreover, *S. aff. ekecaman* occurs in the locality of Kossom Bougoudi, Chad, at the Mio-Pliocene boundary (Bonis et al. 2008). The m1s differ from CM-03-A by very low cusps (hypoconid-protoconid-paraconid of similar height on KNM-KP10034b), an oblique distolingual crest of the paraconid which joins the mesial crest descending from the metaconid to close the trigonid lingual valley, a nearly complete fusion of the basal cuspule of the talonid with the distal face of the protoconid, and a very wide and more basined talonid bordered by cuspules. The M1 (KNM-KP10034c) is distinct from that of the Chiang Muan otter by a narrower paracone labial shelf, a bulbous parastyle, an oblique paraconule-protocone blade, a protocone close to the lingual rim of the tooth, a more extended distolingual cingulum resulting in a distal position of its highest part relative to the metacone, and in unparallel mesial and distal borders of M1.

Sivaonyx africanus is recorded from the Early Pliocene of Klein Zee, South Africa, and possibly

the Late Miocene of Wadi Natrun, Egypt (Stromer, 1931; Morales and Pickford, 2005; Pickford, 2007). The mandible and the associated M1 (BSPG 1930 XI 1) from South Africa are compared here. The m1 shows a wider talonid basin, a metaconid higher than the protoconid, protoconid and paraconid of similar height, a less distinct cuspule at the distal base of the protoconid, and a less opened trigonid lingual valley compared to CM-03-A. The M1 is again wider and mostly longer, and shows an oblique paraconule-protocone blade, a protocone closer to the lingual rim of the tooth, a metaconule closer to the metacone, and an extended distolingual cingulum unlike CM-03-B.

Sivaonyx hendeyi from the earliest Pliocene of Langebaanweg, in South Africa (Morales et al., 2005), shows on its mandible (SAM PQ-L 50000) a more bunodont m1 than CM-03-A with very low cusps, a thicker mesial cingulid, a metaconid higher than the protoconid, a narrower trigonid lingual valley, a poorly distinct cuspule at the distal base of the protoconid, and an entoconid crest higher than the hypoconid one.

Sivaonyx kamuhangirei is reported in the Early Pliocene Kazinga Beds and Warwire Formation of Nkondo area, Uganda (Morales and Pickford, 2005). It is possibly present in the latest Miocene-earliest Pliocene Chadian locality of Kossom Bougoudi (Bonis et al. 2008). The holotype (unnumbered specimen from Kazinga, figured in Morales and Pickford, 2005) differs from CM-03-A by a narrower trigonid lingual valley, a more oblique hypoconid crest, and a stronger entoconid.

Overall, all African species show a more derived dental morphology compared to *Sivaonyx cf. gandakasensis* and Eurasian forms by displaying more bunodont teeth with notably larger proportions, a lower m1 protoconid, and a shorter and wider paraconid. Within Eurasian species, *S. bathygnathus* displays a more derived dental morphology compared to *S. gandakasensis* and *S. hessicus* (e.g., wider m1, reduction of the parastyle and distolingual extension on M1) and thus appears closer to the African species of *Sivaonyx*.

DISCUSSION

Miocene bunodont otters include the genera *Vishnuonyx*, *Djourabus*, *Paludolutra*, *Sivaonyx*, and possibly *Enhydriodon* (Willemsen, 1992; Pickford, 2007; Peigné et al., 2008). Much generic confusion has been experienced for the last three taxa. In the European genus *Paludolutra*, *P. lluecai* was considered as a species of *Enhydriodon* (Villalta and Crusafont-Pairó, 1945) then *Sivaonyx* (Alcalá,

1994), and *P. maremmana* and *P. campanii* were assigned to *Enhydriodon* (Forsyth Major, 1873), a suggestion followed later by many authors. Now these species belong with a general acceptance to *Paludolutra*, mainly based on the morphology of their upper dentition. The attribution of fossil specimens to either *Sivaonyx* or *Enhydriodon* is problematic because the fragmentary material shows poor morphological differences, especially the lower carnassial, and because the stratigraphic occurrence of *Enhydriodon* type species in the Siwalik Hills is uncertain (Pilgrim, 1932). Unlike *Paludolutra*, the confusion between *Sivaonyx* and *Enhydriodon* is still ongoing, mostly for African forms: the species "*S.*" *africanus*, "*S.*" *ekecaman*, and "*S.*" *hendeyi* are assigned either to *Sivaonyx* (e.g., Morales and Pickford, 2005; Peigné et al., 2008) or to *Enhydriodon* (e.g., Geraads et al., 2011). The closer morphological resemblance of the Chiang Muan otter with the type species *S. bathygnathus* rather than with African otters, and its older age, allow us to consider this taxon as a *Sivaonyx* species with confidence. Even if some African forms could be related to *Enhydriodon*, it is likely that they belong to the same bunodont lineage whose primitive features are exhibited by our Thai specimens. These features are a long and narrow paraconid, a high protoconid, a narrow talonid basin and a well-distinct cuspule at the distal base of the protoconid of m1, associated with a wide paracone labial shelf and a more rectangular shape of M1 (i.e., much wider than long and with parallel mesial and distal borders).

The lutrine material from Chiang Muan Basin brings new support for the hypothesis of a unique biogeographical province during the Middle Miocene comprising Southeast Asia and the Siwalik Hills (Ducrocq et al., 1994; Chavasseau et al., 2006). This has been previously revealed by the common contemporaneous and exclusive presence of the rhinocerotid *Chilotherium intermedium* and the suoids *Pecarichoerus sminthos* and *Conohyus sindiensis* in Chiang Muan and in the Potwar Plateau, the latter species also occurring in several localities of the Mae Moh Group of northern Thailand (Pickford et al., 2004; Chavasseau et al., 2006; Suganuma et al., 2006). The geographical distribution of *S. gandakasensis* could also be larger than those of all known African *Sivaonyx* species. Currently living areas of otters are variable among species. For instance, *Aonyx capensis* is present in Southern Africa, northwards to Ethiopia to the east and in Western Africa (Larivière, 2001a) whereas *Aonyx congicus* occurs in central

Africa (Larivière, 2001b). The presence of a unique otter species in the Middle-Late Miocene of Southern Asia may traduce similar environmental conditions between Pakistan, India, and northern Thailand. The sediments from the Ban Sa Mudstone Formation of the Chiang Muan Basin indicate a fluviolacustrine and swamp environment (Chaimanee et al., 2007). Moreover, pollen sampled from the Upper Lignite Member of this locality reveal warm and moist climate conditions with the presence of a mosaic of tropical freshwater swamps dominated by the lowland forest taxon *Syzygium* (Chaimanee et al., 2003; Sepulchre et al., 2010). Similarly, Haritalyangar deposits suggest a river flood plain setting including an open forest vegetation with patches of bamboo (Pillans et al., 2005), and the Potwar Plateau of Pakistan shows during the Early and Middle Miocene extensive river flood plain systems that might have contained a wide variety of habitats, from woodlands to bush and grasslands near the large streams (Barry and Flynn, 1990).

Data from postcranial remains of African bunodont otters, such as *Sivaonyx* species, indicate different locomotion patterns, ranging from semi-aquatic to fully aquatic (Lewis, 2008). The study of the postcranial anatomy of *S. beyi*, the oldest African species, suggests that this otter was a poorly specialized swimmer and probably a terrestrial predator (Peigné et al., 2008). Even if no postcranial remains of Asian *Sivaonyx* species are known, one can assume the capacity of long distance travelling on land, as in modern otters, following freshwater system, as *Sivaonyx* fossil remains are always associated with lacustrine and/or fluvial paleoenvironments. Moreover, the derived dental morphological features of African *Sivaonyx* species compared to Eurasian ones and the closer similarity of *S. bathygnathus* with African *Sivaonyx* species could suggest a dispersal event from Asia to Africa probably during the Late Miocene (between ca. 10 and 7 m.y.). Other bunodont otters could have immigrated to Africa from Asia during the Mio-Pliocene: the genus *Vishnuonyx* (considered as the ancestor of bunodont otters) during the Middle Miocene, and possibly the genus *Enhydriodon* during the Late Miocene/Pliocene. The Middle Miocene carnivoran faunas from Southeast Asia (e.g., Peigné et al., 2006; Grohé et al., 2010) along with further fieldwork efforts in the Miocene of Southern Asia promise to improve our comprehension of the evolutionary history of Carnivora as well as the relevance of Neogene biogeographic scenarii between the Old World continents.

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