



An Early Neogene Elasmobranch fauna from the southern Caribbean (Western Venezuela)

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ABSTRACT

The Cantaure Formation (Burdigalian to ?early Langhian) is located in the Falcón Basin, North Western Venezuela, and includes one of the most diverse Neogene teleostean and benthonic invertebrate faunas in Tropical America. The paleoenvironmental preferences of the members of this fauna, as well as published paleogeographic reconstructions, suggest that the Cantaure Formation was deposited in a highly-productive shallow water environment, associated with coastal upwelling. We documented a paleodiversity of 39 shark and ray species, including 15 previously unreported taxa for Venezuela and six for Tropical America. We performed a bathymetric analysis of the fossil assemblage based on the distribution of closely-related extant chondrichthyan relatives of fossil taxa and discuss the ecological role and stratigraphic significance of the latter. Our results support the hypothesis that the Cantaure Formation was deposited in an insular inner-middle shelf environment. The elasmobranch fauna is characterized by a predominance of benthopelagic sharks with piscivorous feeding preferences (e.g., †*Paratodus*, *Galeorhinus*, *Hemipristis*, *Rhizoprionodon*, *Carcharhinus*, *Isogomphodon*, *Negaprion*, †*Physogaleus* and *Sphyrna*) followed by durophagous/cancritrophic feeders (e.g., *Heterodontus*, *Nebrius*, *Mustelus*, *Rhynchobatus*, *Pristis*, *Dasyatis*, cf. *Pteroplatytrygon*, cf. *Taeniurus*, *Aetobatus*, *Aetomylaeus* and *Rhinoptera*). Filter (e.g., *Mobula* and †*Plinthus*), eurytrophic/sarcophagous (e.g., †*Carcharocles* and *Galeocerdo*) and teuthitrophic (e.g., *Alopias*) feeder species were also found. Teeth of *Carcharocles megalodon* found in Burdigalian sediments of the Cantaure Formation support the presence of this species already in the early Miocene. Some taxa (*Nebrius*, *Carcharhinus* cf. *C. maculoti* and *Rhynchobatus*) are absent from the extant Caribbean and Western Atlantic fauna, but were present in the region before the closure of the Central American Seaway.

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INTRODUCTION

Sharks and rays are essential and diverse elements of tropical marine ecosystems (Cortés et al., 2008; Klimley, 2013), and their fossils are ubiquitous in many localities (e.g., Cappetta, 2012). In “Tropical America” – the geographic area of the Western Hemisphere located between the Tropic of Cancer and the Tropic of Capricorn, Neogene chondrichthyan assemblages are known from Barbados (Casier, 1958, 1966), Brazil (Santos and Travassos, 1960; Santos and Salgado, 1971; Reis, 2005; Costa et al., 2009), Cuba (Iturralde-Vinent et al., 1996; MacPhee et al., 2003), Southern Mexico (González-Rodríguez et al., 2013), Panama (Pimiento et al., 2013a), Peru (Alván et al., 2006), Trinidad (Leriche, 1938), The Grenadines (Portell et al., 2008) and Venezuela (Leriche, 1938; Sánchez-Villagra et al., 2000; Aguilera, 2010; Aguilera and Rodrigues de Aguilera, 2004; Aguilera and Lundberg, 2010). Despite all previous efforts, the chondrichthyan fossil record from the Caribbean is still fragmentary and poorly known, making it difficult to reconstruct paleofaunas that could help our understanding of the major ecological and biogeographical changes that happened in the area during the Neogene. Among those changes are those associated with the uplift of the Central American land bridge (Coates and Stallard, 2013; Montes et al., 2015), which interrupted the interoceanic corridor known as the Central American Seaway (CAS), and influenced the water circulation patterns of Pacific and the Caribbean (O’Dea et al., 2007, Leigh et al., 2014). The geographic position of Venezuelan deposits renders them as relevant for investigating the effects of those major paleogeographic rearrangements.

More than 20 years of paleontological expeditions in the Paraguaná Peninsula, North Western Venezuela, produced a collection of 1836 fossil elasmobranch specimens from the Cantaure Formation. This geological unit is thought to have

been deposited under fully marine conditions (Díaz de Gamero, 1974; Rey, 1996), exhibiting a well-known high diversity of invertebrates (Ingram, 1947; Jung, 1965; Gibson-Smith, 1974; Gibson-Smith and Gibson-Smith, 1974; Aguilera et al., 2010; Landau et al., in press) and teleostean fishes (Nolf and Aguilera, 1998; Aguilera, 2010; Aguilera and Lundberg, 2010). We hereby present a new taxonomic study of the elasmobranch fauna from the Cantaure Formation and a discussion of the paleoenvironmental and feeding preferences of these elasmobranchs, based on comparisons with extant species.

GEOLOGICAL AND STRATIGRAPHIC SETTING

The Cantaure Formation (Jung, 1965) has been dated as early Miocene, based on biozones of planktonic foraminifera (N7-8) and calcareous nannofossils (NN4-5) (Díaz de Gamero, 1974; Rey, 1996). According to Gradstein et al. (2012), bio-zones N7 and N8 correspond to the Lower to Middle Miocene transition (upper Burdigalian to lower Langhian). Griffiths et al. (2013) assigned a Burdigalian age (16.5 ± 0.4 Ma.) to a horizon of the Cantaure Formation using Sr isotopes from corals (Figure 1.2). Landau et al. (in press) suggested that the gastropod assemblage from The Cantaure Formation is typical for early-middle Miocene deposits. The stratotype of the Cantaure Formation is located approximately 10 km west of Pueblo Nuevo on the Paraguaná Peninsula, Falcón State, Venezuela. Outcrops are found south of Casa Cantaure (Figure 1.1) and are composed of fossiliferous silty shales interbedded with thin algal limestones and shell beds (Hunter and Bartok, 1974). Jung (1965) estimated a thickness of 75 m for the composit section of the Cantaure Formation. Hunter and Bartok (1974) divided the Cantaure Formation into three sections: 1) fossiliferous basal breccia unconformably overlaying the granite basement, 2) lower section comprised mainly of

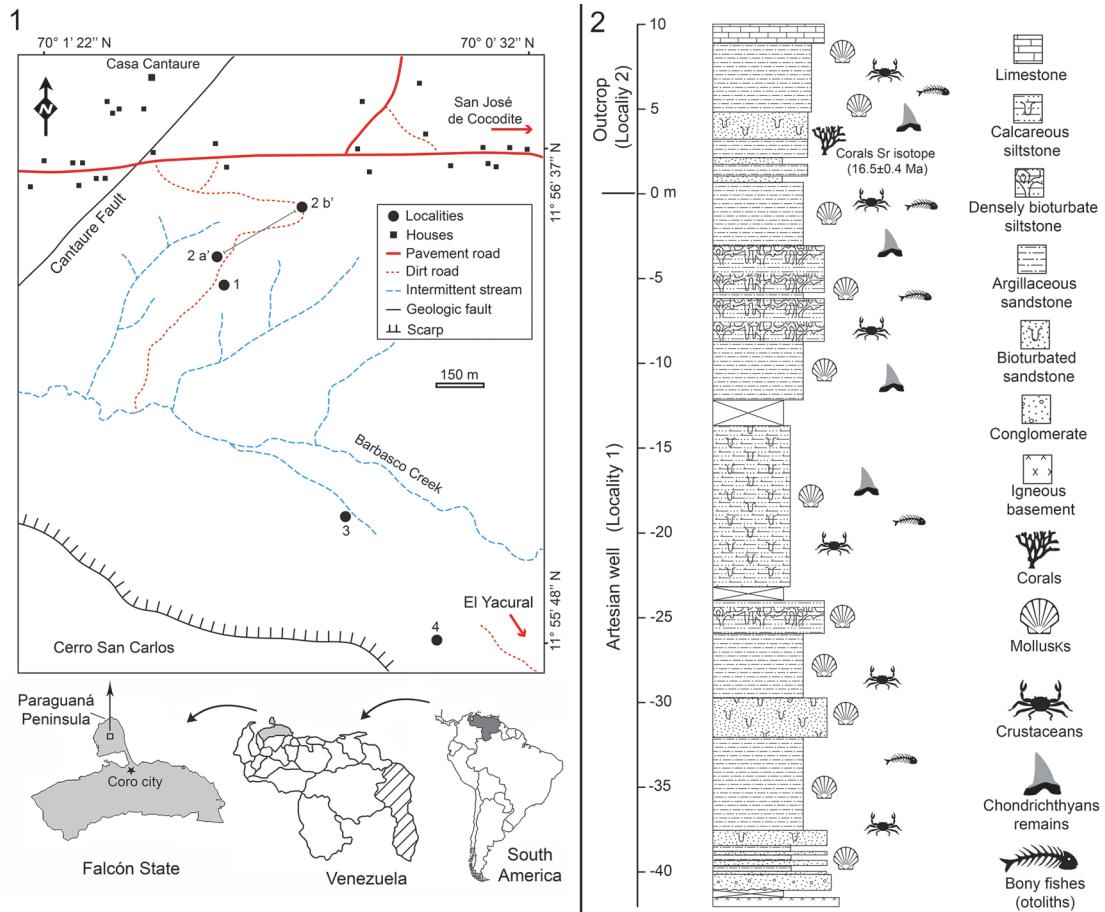


FIGURE 1. Location and stratigraphy of the Cantaire Formation. 1. Fossiliferous localities. 2. Stratigraphic section of locality 1 (artesian well) and locality 2 (outcrop). Stratigraphic column modified after Aguilera (2010) and Aguilera et al. (2013).

shales and some sandy horizons [including the mollusk rich horizons illustrated by Jung (1965)], and 3) upper section represented by silty shales interbedded with thin algal limestones and shell beds. However, Hunter and Bartok (1974) did not provide a stratigraphic column, while they mentioned that the landscape of the area prevents a detailed measurement of the complete section. In contrast, Rey (1996) constructed a stratigraphic column of the suggested basal and middle sections, and also assumed that the mollusks described by Jung (1965), Gibson-Smith (1974) and Gibson-Smith and Gibson-Smith (1974), were derived from the upper section.

The top of the Cantaire Formation is conformably overlain by limestone layers that have been interpreted by Jung (1965) and Hunter and Bartok (1974) as post-Cantaire and pre-Pliocene sedimentation. Rey (1996) suggested that the upper contact of the Cantaire Formation represents an unconformity with the Amuay Member

(early Pliocene) of the Paraguana Formation. Our field observations (OAA, JDJB, pers. obs.) indicate possible lateral variations in a short distance, with fossiliferous outcrops (e.g., thick layers of barnacles) that are not exposed in the previously known and described fossiliferous localities (e.g., Jung, 1965; Hunter and Bartok, 1974; Rey, 1996, among others). All the above clearly indicates that the Cantaire Formation is a more complex sequence than previously thought, in need of a new, detailed stratigraphic redescription.

MATERIAL AND METHODS

The fossil elasmobranch fauna described here (Table 1; Appendix 1) consists of 1836 specimens from four localities in the Cantaire Formation (Figure 1.1), which have been collected by the authors and other collaborators during several expeditions since 1992. Locality 1 (11° 56' 20" N, 70° 1' 2" W) corresponds to an artesian well ~44 m deep (Figure 1.2) and locality 2 (a': 11° 56' 23" N,

TABLE 1. Elasmobranchii paleodiversity of the Cantaure Formation.

Superorder	Order	Family	Genus	Taxon
Galeomorphii	Heterodontiformes	Heterodontidae	<i>Heterodontus</i>	<i>Heterodontus</i> sp.
	Orectolobiformes	Ginglymostomatidae	<i>Nebrius</i>	<i>Nebrius</i> sp.
	Lamniformes	†Otodontidae	† <i>Paratodus</i>	† <i>Paratodus benedenii</i> (Le Hon, 1871)
			† <i>Carcharocles</i>	† <i>Carcharocles megalodon</i> (Agassiz, 1843)
		Alopiidae	<i>Alopias</i>	<i>Alopias</i> cf. <i>A. vulpinus</i> (Bonnaterre, 1788)
	Carcharhiniformes	Triakidae	<i>Galeorhinus</i>	<i>Galeorhinus</i> cf. <i>G. galeus</i> (Linnaeus, 1758)
			<i>Mustelus</i>	<i>Mustelus</i> sp.
		Hemigaleidae	<i>Hemipristis</i>	† <i>Hemipristis serra</i> (Agassiz, 1835)
			<i>Galeocerdo</i>	† <i>Galeocerdo aduncus</i> Agassiz, 1843
		<i>Rhizoprionodon</i>		† <i>Galeocerdo mayumbensis</i> Darteville and Casier, 1943
			<i>Carcharhinus</i>	† <i>Rhizoprionodon</i> sp.
				† <i>Carcharhinus ackermannii</i> Santos & Travassos, 1960
				<i>Carcharhinus brachyurus</i> (Günther, 1870)
				<i>Carcharhinus falciformis</i> (Müller and Henle, 1839)
				<i>Carcharhinus</i> cf. <i>C. galapagensis</i> (Snodgrass and Heller, 1905)
				† <i>Carcharhinus gibbesii</i> (Woodward, 1889)
				<i>Carcharhinus</i> cf. <i>C. limbatus</i> (Müller and Henle, 1839)
				<i>Carcharhinus</i> cf. <i>C. maculot</i> (Müller and Henle, 1839)

70° 1' 3" W and b': 11° 56' 2" N, 70° 0' 55" W) is an outcrop (overlying the section of locality 1) with a ~300 m long collecting surface (Figures 1.1-1.2). The corals used by Griffiths et al. (2013) for Sr isotope analysis were collected at locality 2 (Figure 1.1-1.2). This suggests that localities 1 and at least the lower section of locality 2 are of Burdigalian age (Figure 1.2). Localities 3 (11° 55' 59" N, 70° 1' 1" W) and 4 (11° 55' 48" N, 70° 00' 40" W) correspond to surface collecting areas along the south of Barbasco Creek, on the north escarpment of the Cerro San Carlos (Figure 1.1). Without a detailed stratigraphical knowledge of the outcrops south of Barbasco Creek, it is not possible to establish if localities 3 and 4 are different vertical strata (facies) along the Cantaure section or horizontal variations of localities 1 and 2. Field observations (OAA, JDCB and GAF pers. obs.) suggest that localities 3 and 4 are stratigraphically older than

the limestones layers attributed to post-Cantaure sedimentation (Jung, 1965; Hunter and Bartok, 1974).

Large specimens were surface-collected directly from the outcrop, while microsamples were collected after screen-washing approximately 1000 and 380 kg of sediments (mesh sizes: 0.5 and 2 mm) from locality 1 and 2, respectively. Approximately 50 kg of sediment was processed from locality 3, but it did not contain otoliths or elasmobranch micro-teeth. Specimens from locality 4 were only surface collected. The Cantaure specimens (Appendix 2) are housed in the paleontological collections of the Alcaldía Bolivariana de Urumaco (AMU-CURS), Centro de Investigaciones Antropológicas, Arqueológicas y Paleontológicas of the Universidad Experimental Francisco de Miranda (CIAAP, UNEFM-PF), and the Museo de Ciencias de Caracas (MCNC), all in Venezuela.

TABLE 1 (continued).

Superorder	Order	Family	Genus	Taxon
				<i>Carcharhinus perezi</i> (Poey, 1876)
				<i>Carcharhinus</i> sp. 1
				<i>Carcharhinus</i> spp.
			<i>Isogomphodon</i>	† <i>Isogomphodon acuaricus</i> (Probst, 1879)
			<i>Negaprion</i>	† <i>Negaprion eurybathrodon</i> (Blake, 1862)
			<i>†Physogaleus</i>	† <i>Physogaleus contortus</i> (Gibbes, 1849)
			Gen. indet.	sp. indet.
		Sphyrnidae	<i>Sphyrna</i>	† <i>Sphyrna arambourgi</i> Cappetta, 1970
				<i>Sphyrna</i> cf. † <i>S. laevissima</i> (Cope, 1867)
Batomorphii	Rajiformes	Rhynchobatidae	<i>Rhynchobatus</i>	<i>Rhynchobatus</i> sp.
		Pristidae	<i>Pristis</i>	<i>Pristis</i> sp.
	Myliobatiformes	Dasyatidae	<i>Dasyatis</i>	<i>Dasyatis</i> sp.
			cf. <i>Pteroplatytrygon</i>	cf. <i>Pteroplatytrygon</i> sp.
			cf. <i>Taeniurus</i>	cf. <i>Taeniurus</i> sp.
				Dasyatidae Indet.
		Myliobatidae	<i>Aetobatus</i>	<i>Aetobatus</i> sp.
			<i>Aetomylaeus</i>	<i>Aetomylaeus</i> sp.
		Rhinopteridae	<i>Rhinoptera</i>	<i>Rhinoptera</i> sp.
		Mobulidae	<i>Mobula</i>	† <i>Mobula fragilis</i> (Cappetta, 1970)
				<i>Mobula</i> cf. † <i>M. loupiensis</i> (Cappetta, 1970)
				<i>Mobula</i> sp.
			† <i>Plinthiscus</i>	† <i>Plinthiscus stenodon</i> Cope, 1869
				Myliobatiformes indet.
				Chondrichthyes indet.

The taxonomy follows Cappetta (2012) and Compagno (2005), with the exception of the extinct genus *Carcharocles* Jordan and Hannibal, 1923, for which we follow the nomenclature of Pimiento et al. (2010). Taxonomic identification was based on an extensive literature review, and also comparative analyses between fossil and recent specimens from the following collections: Mapuka Museum of Universidad del Norte (MUN-STRI), Barranquilla, Colombia; Museu Paraense Emilio Goeldi (MPEG-V), Belem, Brazil; Natural History Museum of Basel (NMB), Switzerland; Paleontological collection of the Institut des Sciences de l' Evolution, University of Montpellier (UM), France; Palaeontological Institute and Museum at the University of Zurich (PIMUZ), Switzerland and René Kindlimann private collection, Uster, Switzerland.

We analyzed the abundance using percentages of specimens by order, families and species

using the 39 securely recognized taxa that are represented by 1711 specimens (out of the 1836 collected in total) (Appendix 1). For conducting the paleoecological interpretation, we compiled information on habitat preference (benthic, benthopelagic, pelagic, neritic and bathyal) and feeding ecology (dietary composition and behavior) of extant taxa recorded in the Cantare Formation (Appendix 3), following Compagno (1984a, 1984b); Cortés (1999); Compagno et al. (2005); Musick et al. (2004); Kiraly et al. (2003); Voigt and Weber (2011); Cortés et al. (2008); Ebert and Stehmann (2013); and the FishBase website (Froese and Pauly, 2015). The paleobathymetric analysis was performed following the methodology of Nolf and Brzobohatý (1994), adapted to fossil sharks (Carrillo-Briceño et al., 2015a, 2016). For this analysis we included only species/genera with closely related extant taxa. Extinct species and taxa with-

out clear identification to genus were removed from the analysis. A total of 34 (out of 39) taxa, for which ecological information is available, were analyzed.

Taking into consideration the limitations associated with paleoecological interpretations of fossil assemblages, we tried to approach the ecological role that members of the Cantaure elasmobranch fauna played in this ancient marine environment. We use the categorization of elasmobranch feeding preferences suggested by Cortés (1999) and Cortés et al. (2008). Living sharks and rays as a whole exhibit a wide range of diets, but each species has specific preferences. The latter are used to infer the diet of taxa we documented from the Cantaure assemblage (Appendix 3).

RESULTS

Taxonomical Composition

The taxonomical composition of the four study localities includes at least 39 species of galeomorphs and batoids (Table 1; Appendix 1, 4). Galeomorphs are represented by 26 species, 15 genera and eight families of Heterodontiformes, Orectolobiformes, Lamniformes and Carcarhiniiformes (Table 1; Figure 2). Batoids include 13 species attributed to 10 genera and six families of Rajiformes and Myliobatiformes (Table 1; Figure 2). Pictures of the chondrichthyan teeth, vertebrae and caudal and rostral denticles in Figures 3-12, were illustrated following the taxonomic classification shown in Table 1.

Heterodontiformes Berg, 1937. The record of this group (Heterodontidae Gray, 1851) consists of a single and undetermined specimen of *Heterodontus* Blainville, 1816. The tooth was collected from locality 2 (Appendix 1, 4), and was previously reported by Aguilera and Lundberg (2010) and Aguilera (2010). The specimen is missing from the UNEFM collection and it is not illustrated here.

Orectolobiformes Applegate, 1972. The orectolobiform specimens are characterized by a long apron and a short main cusp flanked by more than seven pairs of lateral cusplets (Figure 3.1-3.9). This condition is diagnostic for *Nebrius* Rüppell, 1837, a genus recorded from North America going back to the Eocene (e.g., Kent, 1994; Müller, 1999). The specimens from the Cantaure assemblage (Figure 3.1-3.9; Appendix 1, 4) were previously assigned to *Ginglymostoma delfortriei* (Daimeries, 1889) (Aguilera and Rodrigues de Aguilera, 2004; Aguilera, 2010; Aguilera and Lundberg, 2010). Fossils of *Nebrius* in the Americas have not been thoroughly studied, and there is no

clear consensus about their taxonomy at the species level or stratigraphic range (e.g., Kent, 1994; Cicimurri and Knight, 2009). Therefore we refrain from assigning species to our specimens. Cicimurri and Knight (2009) suggested that some fossil teeth of *Nebrius* (from North America) have been misidentified as *Ginglymostoma* Müller and Henle, 1837. The specimens referred to *G. delfortriei* from Panama (Pimiento et al., 2013b) and *Ginglymostoma* sp. from Brazil (Costa et al., 2009) resemble our *Nebrius* teeth from the Cantaure Formation.

Lamniformes Berg, 1937. These sharks are represented by one species of Alopiidae Bonaparte, 1838 and two genera and two species of †Otodontidae Glikman 1964 (Figure 2). *Alopias* cf. *A. vulpinus* (Bonnaterre, 1788) is represented by three specimens (Figure 3.22-3.26; Appendix 1, 3) while the otodontid †*Paratodus benedenii* (Le Hon, 1871) is the least abundant lamnid (Figure 3.10-3.11; Appendix 1, 4). The most abundant lamnid is the otodontid †*Carcharocles megalodon* (Agassiz, 1843), with more than 100 specimens (Figure 3.12-3.21) from all studied localities (Appendix 1, 4). Many *C. megalodon* teeth are broken; however, most of them, especially large specimens, preserve the wide triangular shape (crown without lateral cusplets), a v-shaped 'neck', a fine serrations, and the large size that characterize this species (Purdy et al., 2001; Cappetta, 2012; Pimiento et al., 2010, 2013b). Small teeth having a pair of short lateral cusplets may belong to juvenile and/or subadult individuals of *C. megalodon* (Figure 3.12-3.16), a feature not present in adult teeth (Figure 3.17-3.21). Morphologically, the younger species *Carcharocles chubutensis* (Ameghino, 1901), differs from *C. megalodon* by the presence of lateral cusplets not separated from the crown on the teeth of juveniles, subadults and adults (Cappetta, 2012; Pimiento et al., 2013a). In more derived populations as *C. megalodon*, lateral cusplets disappear, except sometimes on very lateral files or in juvenile individuals (Cappetta, 2012). This absence of lateral cusplets in *C. megalodon* adult teeth has been interpreted as the result of heterochronic processes in the *Carcharocles* clade through geologic time (Pimiento et al., 2010, 2013a, 2013b; Pimiento and Balk, 2015). The absence of lateral cusplets in all well-preserved teeth interpreted as adults (including also some teeth that may belong to juvenile/subadults specimens) allows us to assign our specimens to *C. megalodon*. In addition, adult teeth of *C. megalodon* referred here, are very similar to those of the

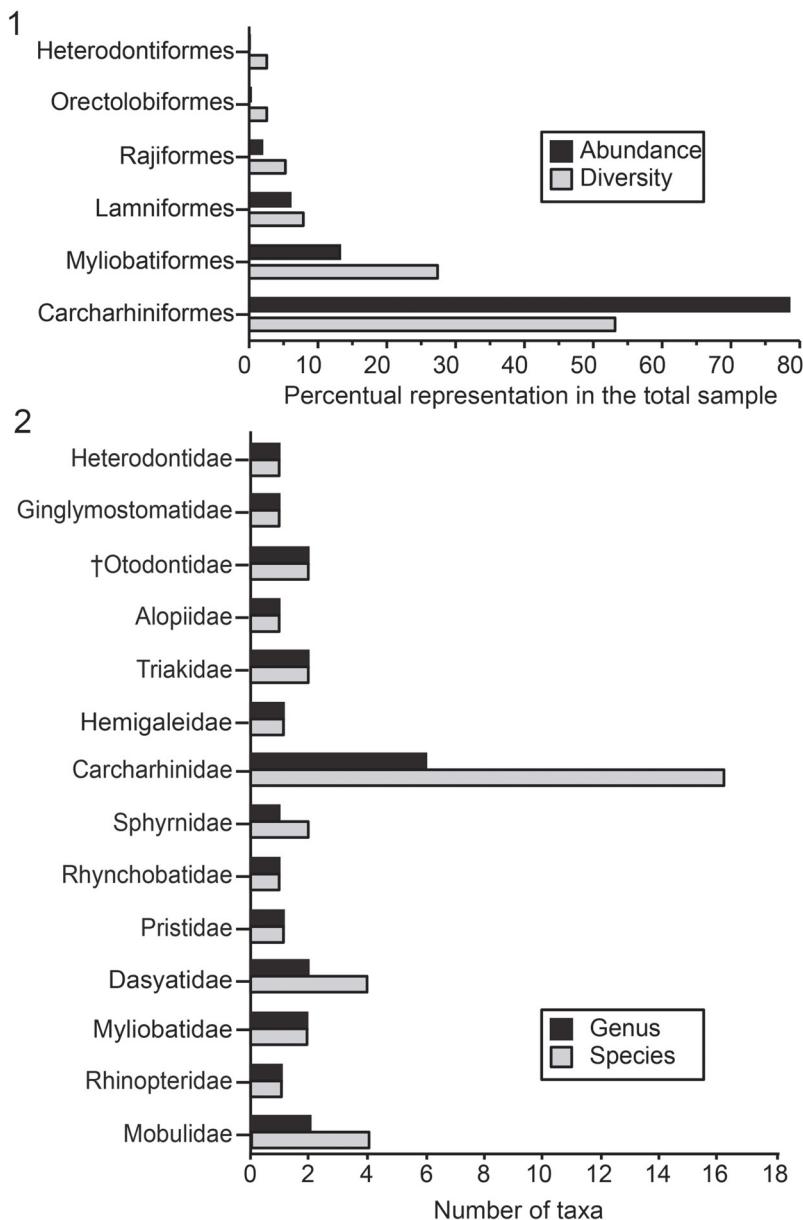


FIGURE 2. Elasmobranch paleodiversity of the Cantaire Formation. 1. Orders and 2. Families and genera.

same species, found in the late Burdigalian section (~16.2 Ma.) of the Castilletes Formation at North East Colombia (under study by JDGB).

Carcharhiniformes Compagno, 1973. This order is represented by four families, 10 genera and 21 species, thus being the most abundant and diverse one of the Cantaire assemblage (Figures 2, 4-8; Table 1; Appendix 1, 4). Triakidae Gray, 1851 [two genera and two species: *Galeorhinus* cf. *G. galeus* (Linnaeus, 1758) (Figure 4.1-4.2), and *Mustelus* Linck, 1790 (*Mustelus* sp.) (Figure 4.3-4.7)], Hemigaleidae Hasse, 1879 [*†Hemipristis serra* (Agassiz, 1835) (Figure 4.8-4.13)], and Sphyrnidae Gill,

1872 [*†Sphyrna arambourgi* Cappetta, 1970 (Figure 7.21-7.25), and *Sphyrna* cf. *†S. laevissima* (Cope, 1867) (Figure 8.1-8.5)], are the less diverse families of this order in the Cantaire assemblage (Figure 2; Appendix 1, 4). Nevertheless, *Hemipristis serra* with 425 specimens (Appendix 1, 4), is the second most abundant species of the assemblage, representing 23% of the total specimens collected. The specimen referred by Aguilera and Rodrigues de Aguilera (2004) to *Paragaleus* sp. corresponds to a juvenile posterior ?lower tooth of *H. serra* (Figure 4.8).

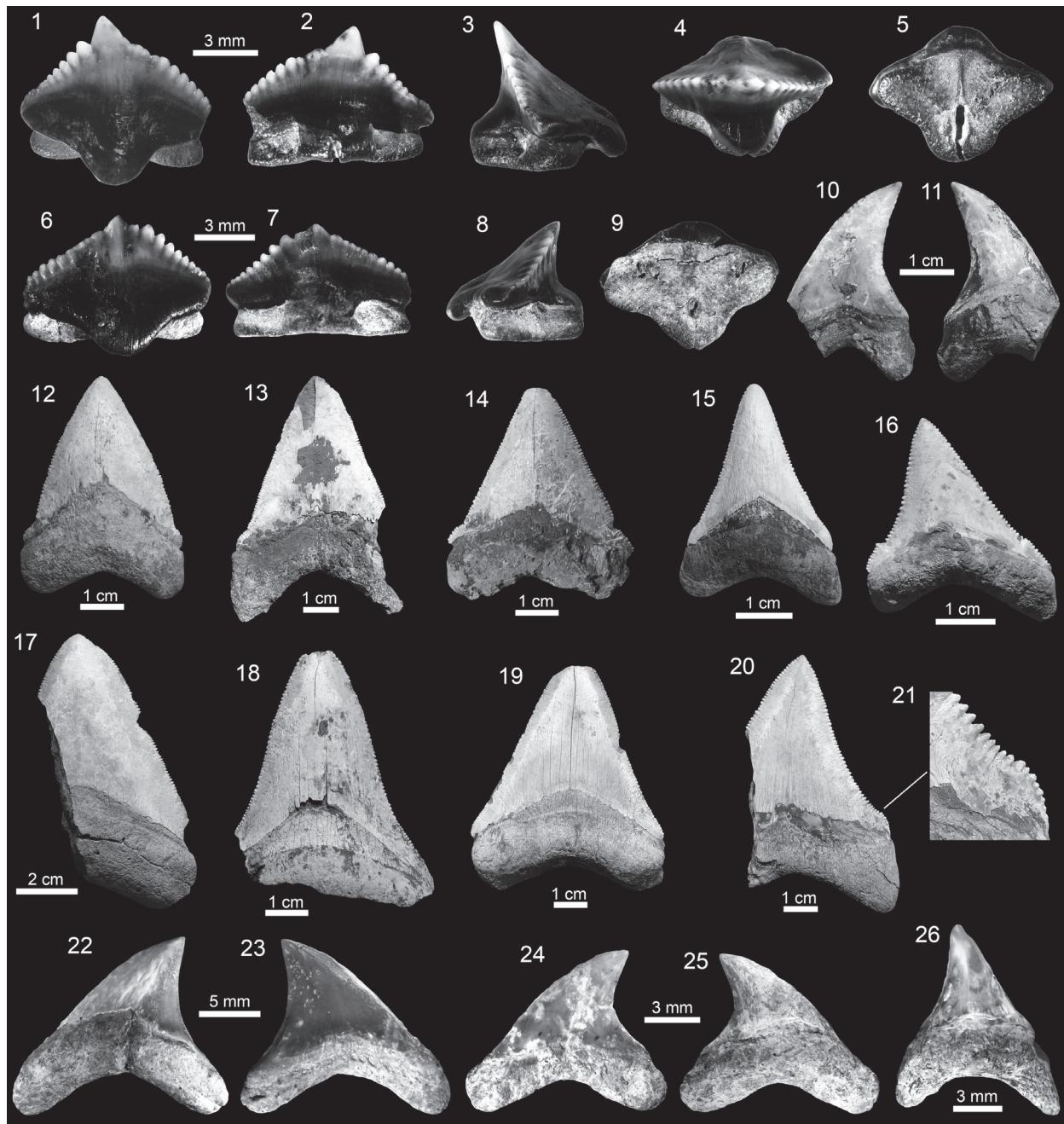


FIGURE 3. Orectolobiformes and Lamniformes of the Cantaure Formation. 1-9. *Nebrius* sp. (AMU-CURS-986); 10-11. †*Paratodus benedenii* (AMU-CURS-645). 12-21. †*Carcharocles megalodon* (12: AMU-CURS-997; 13: AMU-CURS-512; 14: UNEFM-PF-349; 15: AMU-CURS-996; 16: AMU-CURS-518; 17-18: UNEFM-PF-s/n; 19: AMU-CURS-512 and 20-21: AMU-CURS-515). 22-26. *Alopias* cf. *A. vulpinus* (22-23: AMU-CURS-985; 24-25: AMU-CURS-983 and 26: AMU-CURS-984). Jaw position: upper (10-11? 12, 14? 17, 22-25), lower (13, 15, 18) and indet. (1-9, 16, 19-21, 26). View: labial (1, 6, 10, 13, 20-21, 23, 24), lingual (2, 7, 11-12, 14-19, 22, 25-26), profile (3, 8), occlusal (4) and basal (5, 9).

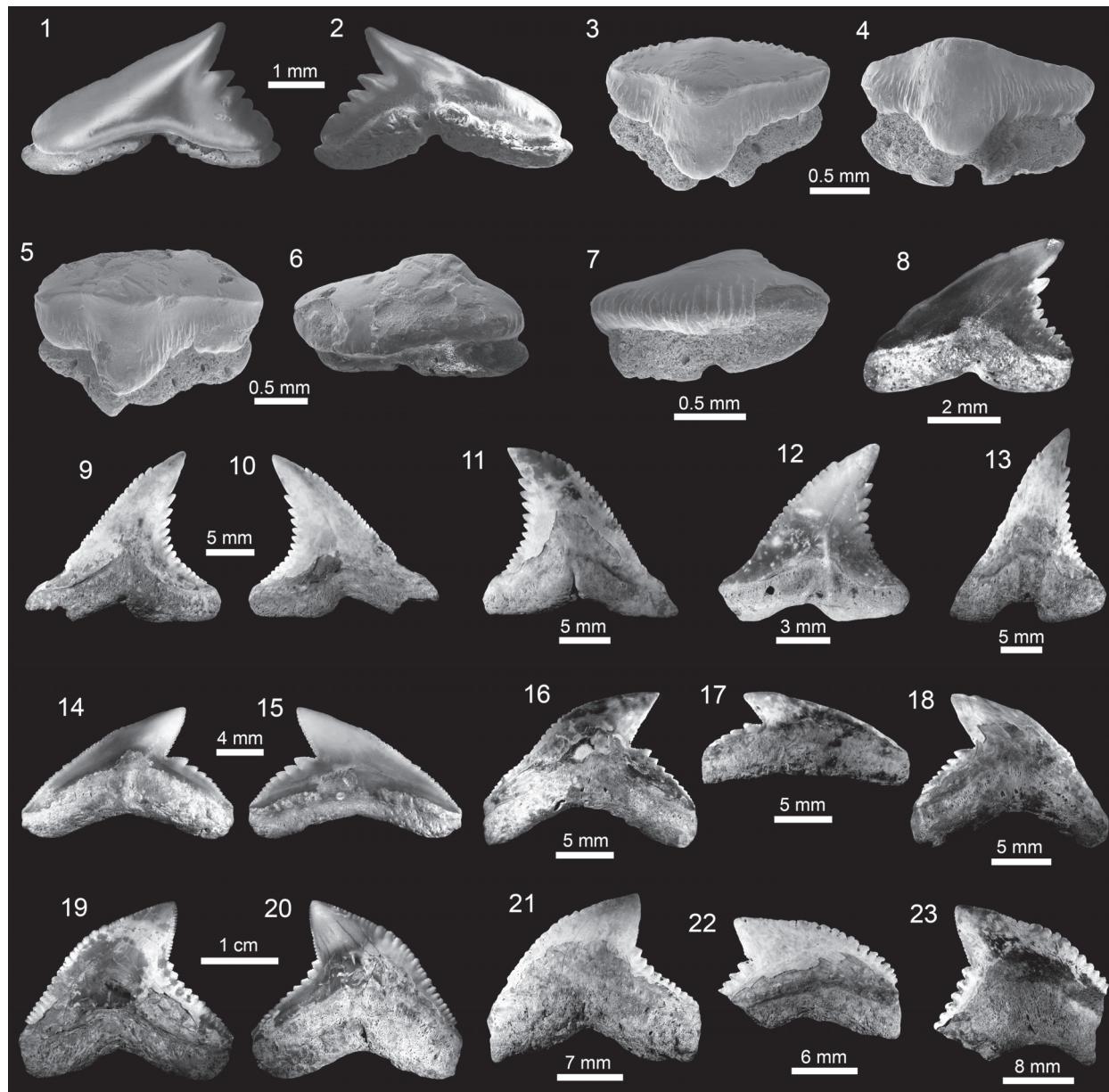


FIGURE 4. Carcharhiniformes of the Cantaure Formation. 1-2. *Galeorhinus* cf. *G. galeus* (AMU-CURS-974). 3-7. *Mustelus* sp. (AMU-CURS-975). 8-13. †*Hemipristis serra* (8: AMU-CURS-731 and 9-13: AMU-CURS-644). 14-18. †*Galeocerdo aduncus* (14-15: AMU-CURS-730 and 16-18: AMU-CURS-647). 19-23. †*Galeocerdo mayumbensis* (19-20: AMU-CURS-995 and 21-23: AMU-CURS-646). Jaw position: upper (9-12), lower (8? 13) and indet. (1-7, 14-23). View: labial (1, 6, 7, 10, 15, 19, 21-22), lingual (2, 4, 8-9, 11-14, 16-18, 20, 23) and occlusal-lingual (3, 5).

The Carcharhinidae Jordan and Evermann, 1896, with six genera and 16 species, is the most diverse and abundant family of the assemblage (Figure 2; Table 1; Appendix 1, 4). The genera *Rhizoprionodon* Whitley, 1929 [*Rhizoprionodon* sp. (Figure 5.1-5.5)], *Isogomphodon* Gill, 1862 [†*Isogomphodon acuarius* (Probst, 1879) (Figure 7.1-7.5)], *Negaprion* Whitley, 1940 [†*Negaprion eurybathrodon* (Blake, 1862) (Figure 7.6-7.10)], and †*Physoga-*

leus Cappetta, 1980 [†*Physogaleus contortus* (Gibbes, 1849) (Figure 7.11-7.15)], are represented by one species each. *Galeocerdo* Müller and Henle, 1837 is represented by two species [†*Galeocerdo aduncus* Agassiz, 1843 (Figure 4.14-4.18), and †*Galeocerdo mayumbensis* Darteville and Casier, 1943 (Figure 4.19-4.23)]. *Negaprion eurybathrodon* with 586 specimens (Appendix 1, 4) is the most abundant species of the assemblage,

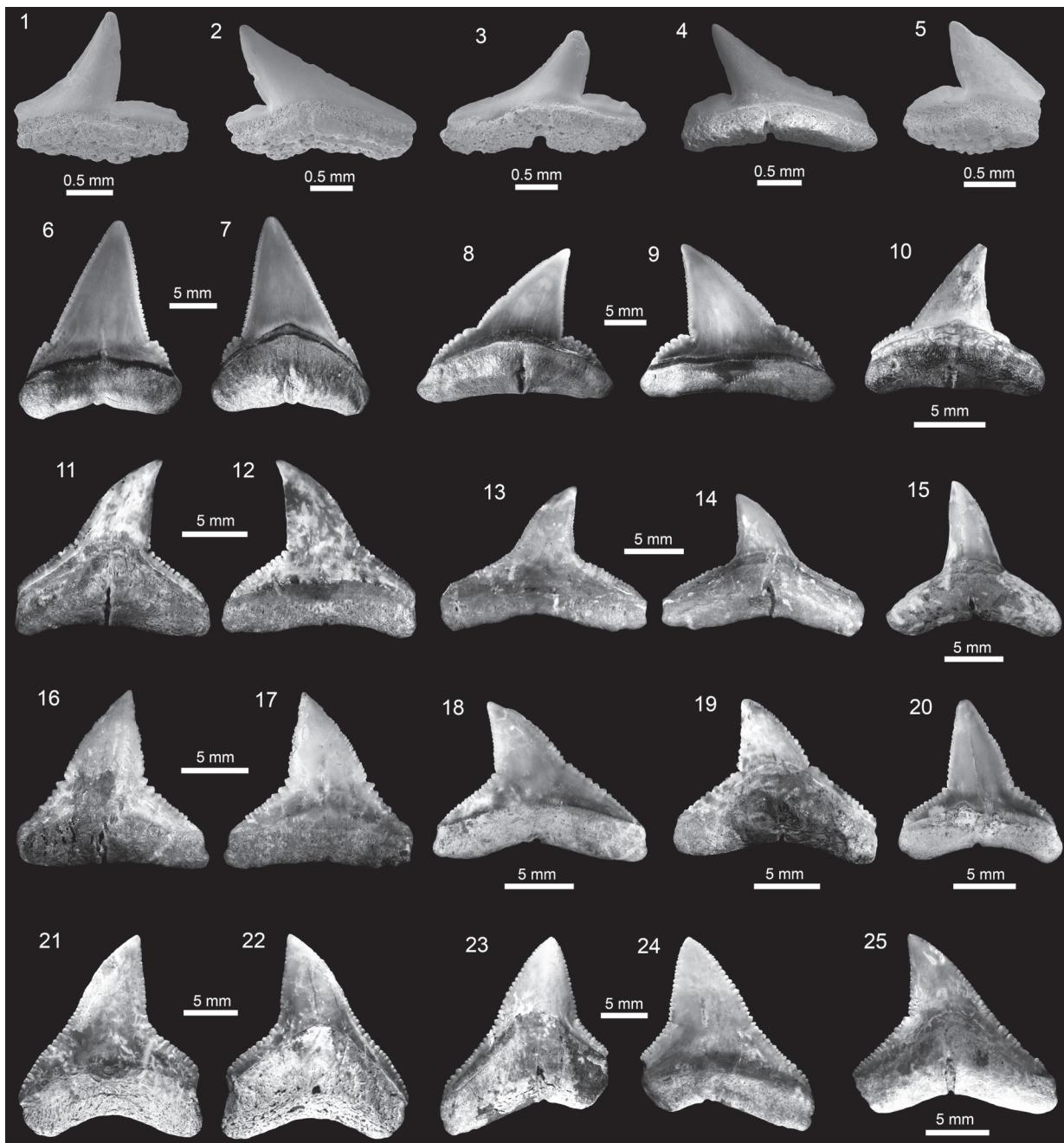


FIGURE 5. Carcharhiniformes of the Cantaire Formation. 1-5. *Rhizoprionodon* sp. (1-3, 5: AMU-CURS-975 and 4: AMU-CURS-960). 6-10. †*Carcharhinus ackermannii* (6-7: AMU-CURS-713 and 8-10: AMU-CURS-714). 11-15. *Carcharhinus brachyurus* (AMU-CURS-990). 16-20. *Carcharhinus falciformis* (AMU-CURS-991). 21-25. *Carcharhinus* cf. *C. galapagensis* (AMU-CURS-973). Jaw position: upper (2, 4-14, 16-25) and lower (1, 3, 15). View: labial (6, 9, 12-13, 17-18, 20-21, 24) and lingual (1-5, 7-8, 10-11, 14-16, 19, 22-23, 25).

representing 32% of the total specimens collected. Specimens referred to *Mustelus* sp. (Triakidae) and *Rhizoprionodon* sp. (Carcharhinidae), are few and fragmentary, lacking diagnostic characters for accurate species assignment.

The genus *Carcharhinus* Blainville, 1816, with nine species, represents the most diverse genus

from the Cantaire assemblage (Figures 2, 5.6-5.25, 6.1-6.24; Table 1; Appendix 1, 4). With the exception of the extinct †*Carcharhinus ackermannii* Santos and Travassos, 1960 (Figure 5.6-5.10) and †*Carcharhinus gibbesii* (Woodward, 1889) (Figure 6.1-6.4), the species *Carcharhinus brachyurus* (Günther, 1870) (Figure 5.11-5.15),

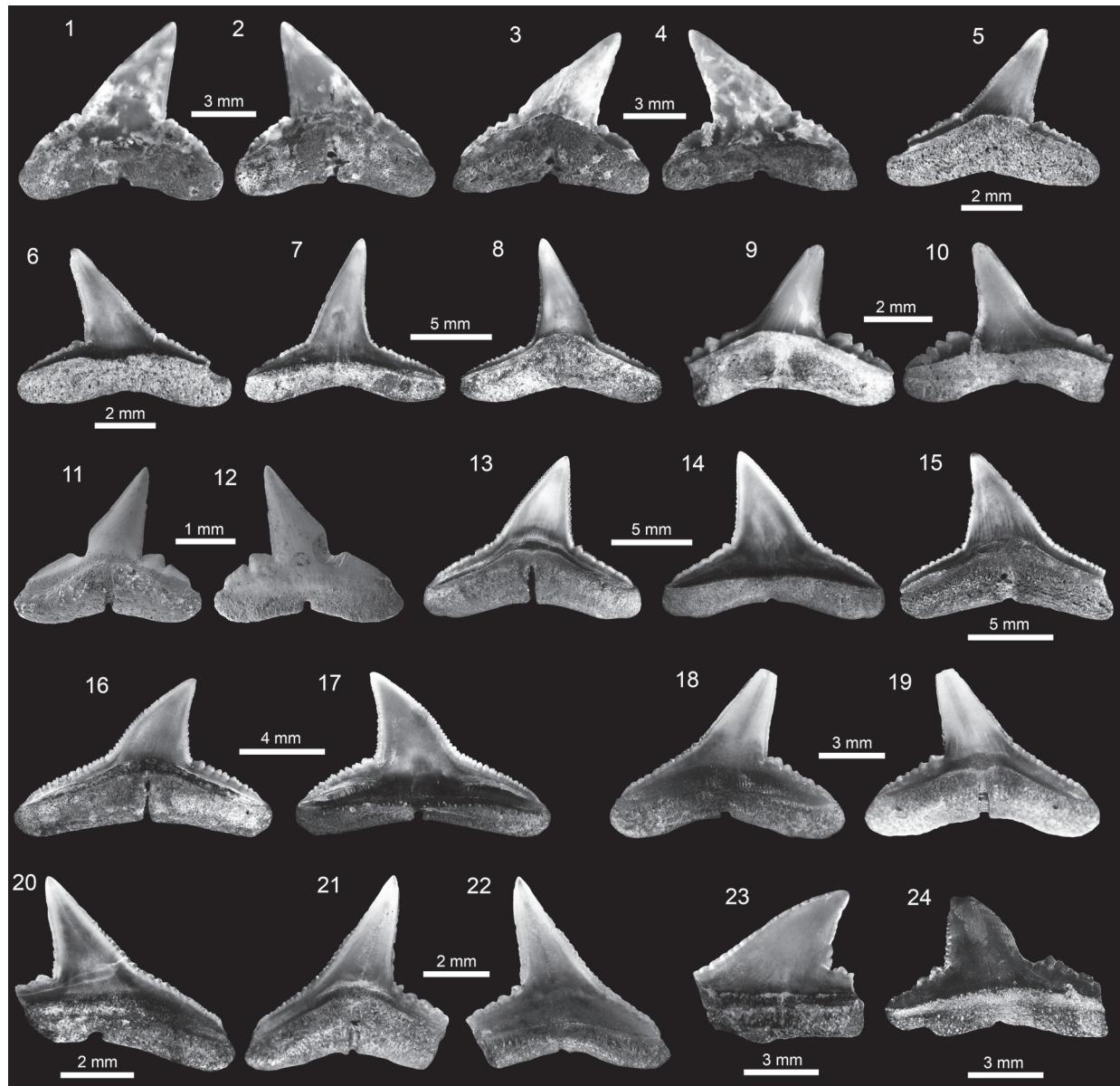


FIGURE 6. Carcharhiniformes of the Cantaure Formation. 1-4. †*Carcharhinus gibbesii* (AMU-CURS-958). 5-8. *Carcharhinus* cf. *C. limbatus* (AMU-CURS-944). 9-12. *Carcharhinus* cf. *C. maclooti* (AMU-CURS-957). 13-17. *Carcharhinus perezi* (AMU-CURS-956). 18-22. *Carcharhinus* sp.1 (AMU-CURS-943). 23-24. *Carcharhinus* spp. (23: AMU-CURS-941 and 24: AMU-CURS-940). Jaw position: upper (1-24). View: labial (1, 4, 6-7, 10, 12, 14, 17-18, 20, 22-24) and lingual (2-3, 5, 8-9, 11, 13, 15-16, 19, 21).

Carcharhinus falciformis (Müller and Henle, 1839) (Figure 5.16-5.20), *Carcharhinus* cf. *C. galapagensis* (Snodgrass and Heller, 1905) (Figure 5.21-5.25), *Carcharhinus* cf. *C. limbatus* (Müller and Henle, 1839) (Figure 6.5-6.8) and *Carcharhinus perezi* (Poey, 1876) (Figure 6.13-6.17), have living counterparts and most of them inhabit the Atlantic and Pacific coast of America (Compagno et al., 2005). *Carcharhinus maclooti* (Müller and Henle, 1839) (Figure 6.9-6.12) is today an exclusive

inhabitant of the Indo-west Pacific Ocean (Compagno et al., 2005). The specimens referred to *Carcharhinus* sp.1 (Figure 6.18-6.22), have a dental morphology different, especially in the form of the crown, which differentiate these teeth from the rest of the *Carcharhinus* species found in the Cantaure assemblage. Due to the scarcity of specimens (Appendix 1) and the lack of preserved diagnostic characters in *Carcharhinus* teeth we have used for comparison (fossil and recent speci-

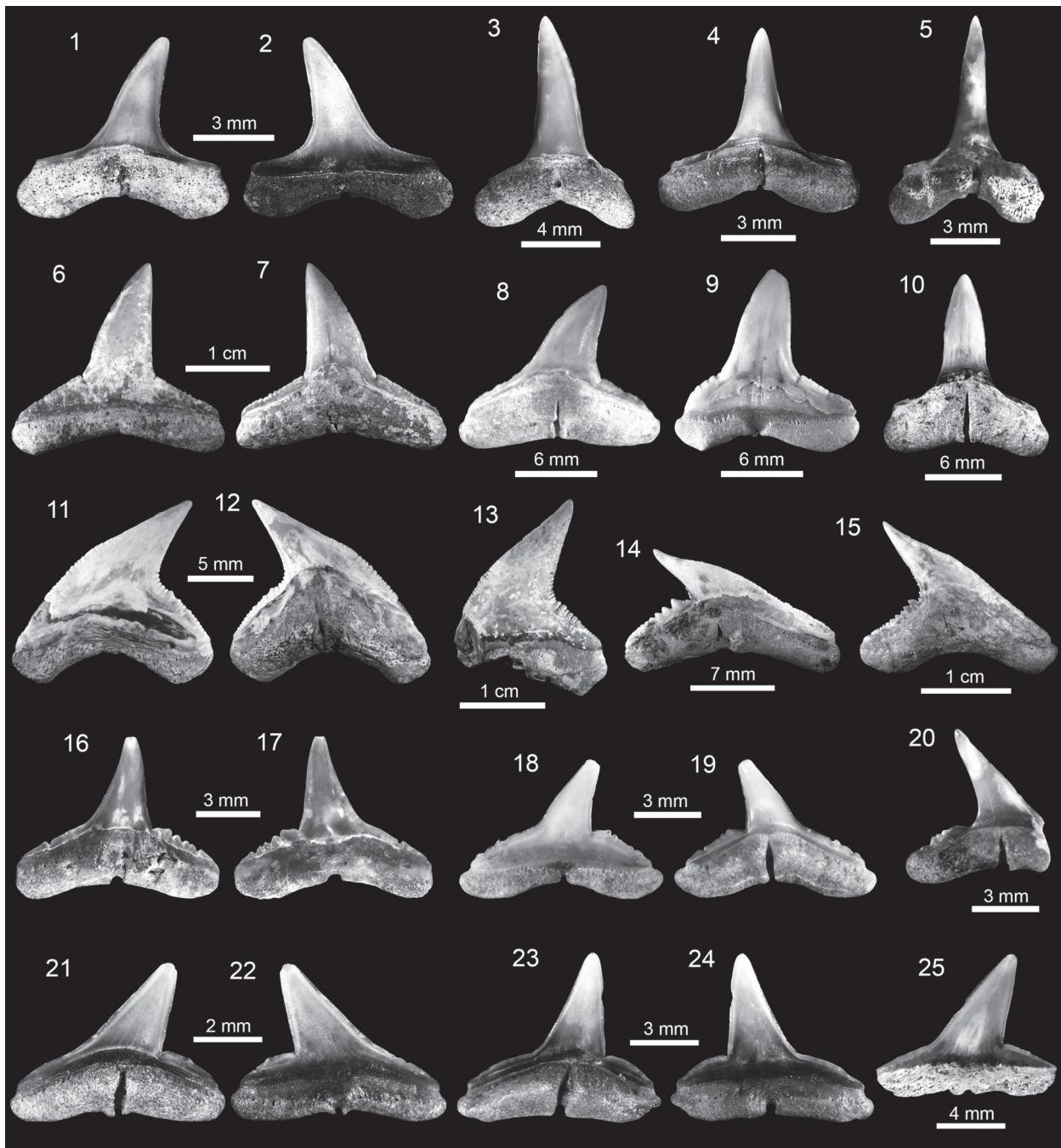


FIGURE 7. Carcharhiniformes of the Cantaure Formation. 1-5. †*Isogomphodon acuaricus* (1-2: AMU-CURS-950 and 3-5: AMU-CURS-951). 6-10. †*Negaprion eurybathrodon* (6-9: AMU-CURS-989 and 10: AMU-CURS-987). 11-15. †*Physogaleus contortus* (11-12: AMU-CURS-719 and 13-15: AMU-CURS-648). 16-20. Carcharhinidae indet. (AMU-CURS-939). 21-25. †*Sphyrna arambourgi* (21-24: AMU-CURS-961 and 25: AMU-CURS-962). Jaw position: upper (1-3, 6-9, 16-19? 21-22, 25), lower (4-5, 10, 20? 23-24) and indet. (11-15). View: labial (2, 6, 9, 11, 13, 17-18, 22, 24-25) and lingual (1, 3-5, 7-8, 10, 12, 14-16, 19-21, 23).

mens), we prefer to keep these specimens to a generic level. We have referred all specimens that are broken and eroded, without any diagnostic features for specific identification (Figure 6.23-6.24), to *Carcharhinus* spp. Some of the specimens clas-

sified as Carcharhinidae indet. (Figure 7.16-7.20; Table 1; Appendix 1, 4), exhibit a very prominent lingual protuberance with a sigmoid cusp, which could be the basis of a more detailed taxonomical identification in future studies.

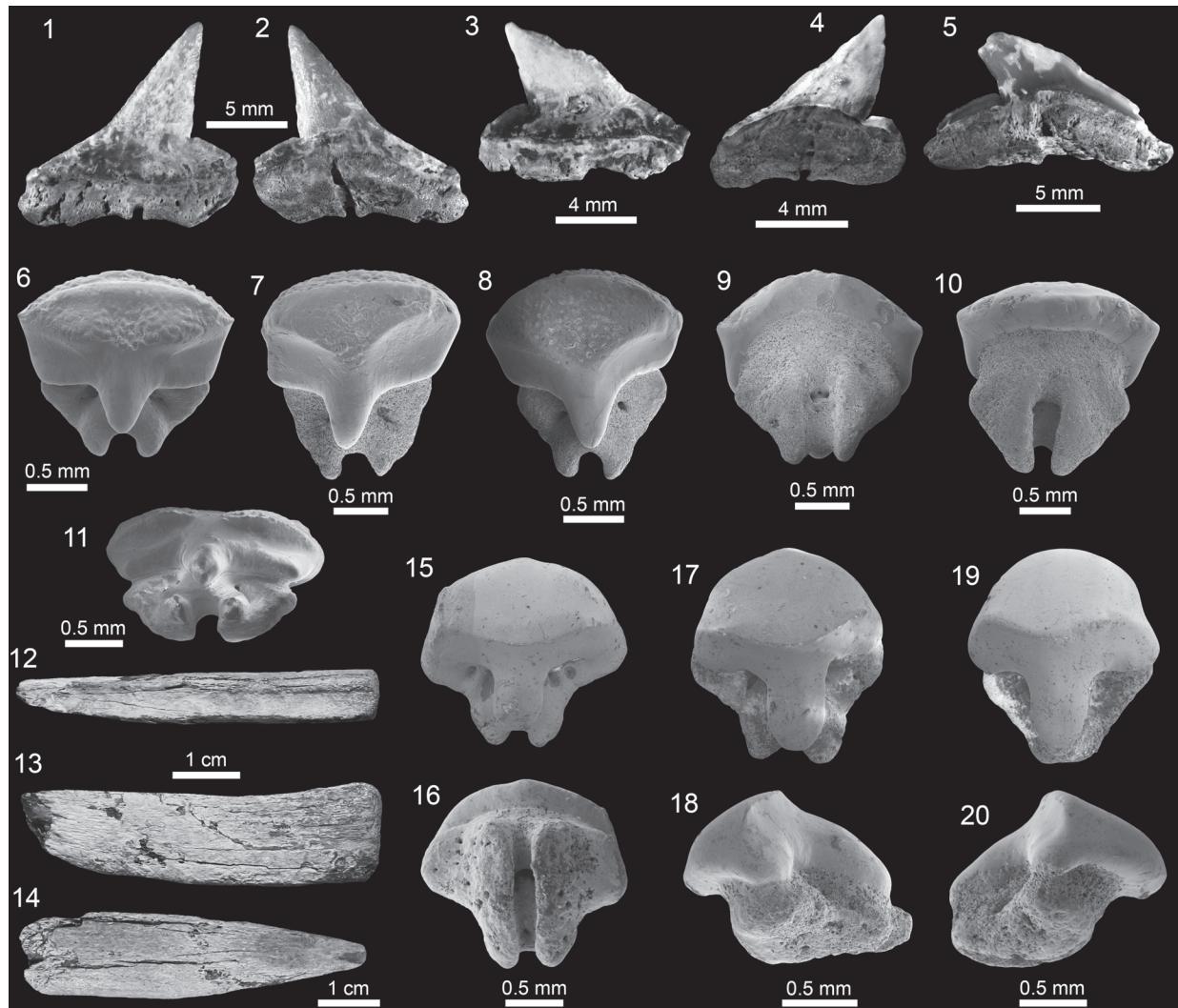


FIGURE 8. Carcharhiniformes and Rajiformes of the Cantaure Formation. 1-5. *Sphyrna* cf. †*S. laevissima* (1-3, 5: AMU-CURS-964 and 4: AMU-CURS-963). 6-11. *Rhynchobatus* sp. (6, 9: AMU-CURS-968; 7-8, 10: AMU-CURS-969 and 11: AMU-CURS-970). 12-20. *Pristis* sp. (rostral denticle 12-14: AMU-CURS-242; buccal teeth 15-20: AMU-CURS-967). Jaw position: upper (1-2) and indet. (3-11, 15-20). View: labial (1, 3), lingual (2, 4-5, 11), occlusal (6-8, 15, 17, 19), profile (18, 20), basal (9-10, 16), posterior (12) and dorsal (13-14).

Rajiformes Berg, 1937. This group of batoids is represented by two families, two genera and two species (Figure 2). Concerning Rhynchobatidae Garman, 1913, we report indeterminate teeth of *Rhynchobatus* Müller and Henle 1837 (Figure 8.6-8.11; Appendix 1, 4), resembling previously figured isolated teeth of *Rhynchobatus* from the late Miocene of Costa Rica, Panama and Venezuela (Laurito, 1999; Pimiento et al., 2013b; Carrillo-Briceño et al., 2015b). We refrain from taxonomic identification at the species level of the Cantaure *Rhynchobatus*, because the range of dental variation in extant species is unknown, and little is known about fossil species from the Americas. The *Pristi-*

dae Bonaparte, 1838 is represented by a few specimens of oral teeth and rostral denticles of *Pristis* Linck, 1790 (Figure 8.12-8.20, Appendix 1, 4). As noted by Carrillo-Briceño et al. (2015b), *Pristis* rostral denticles are not diagnostic at species level, while isolated oral teeth are somewhat difficult to identify, because those of extant species are poorly known and are morphologically diverse.

Myliobatiformes Compagno, 1973. This group is represented by four families, eight genera and 11 species, being thus the second most abundant and diverse order from the Cantaure assemblage (Figure 2; Table 1; Appendix 1, 4). The Dasyatidae Jordan, 1888, is represented by three genera and four

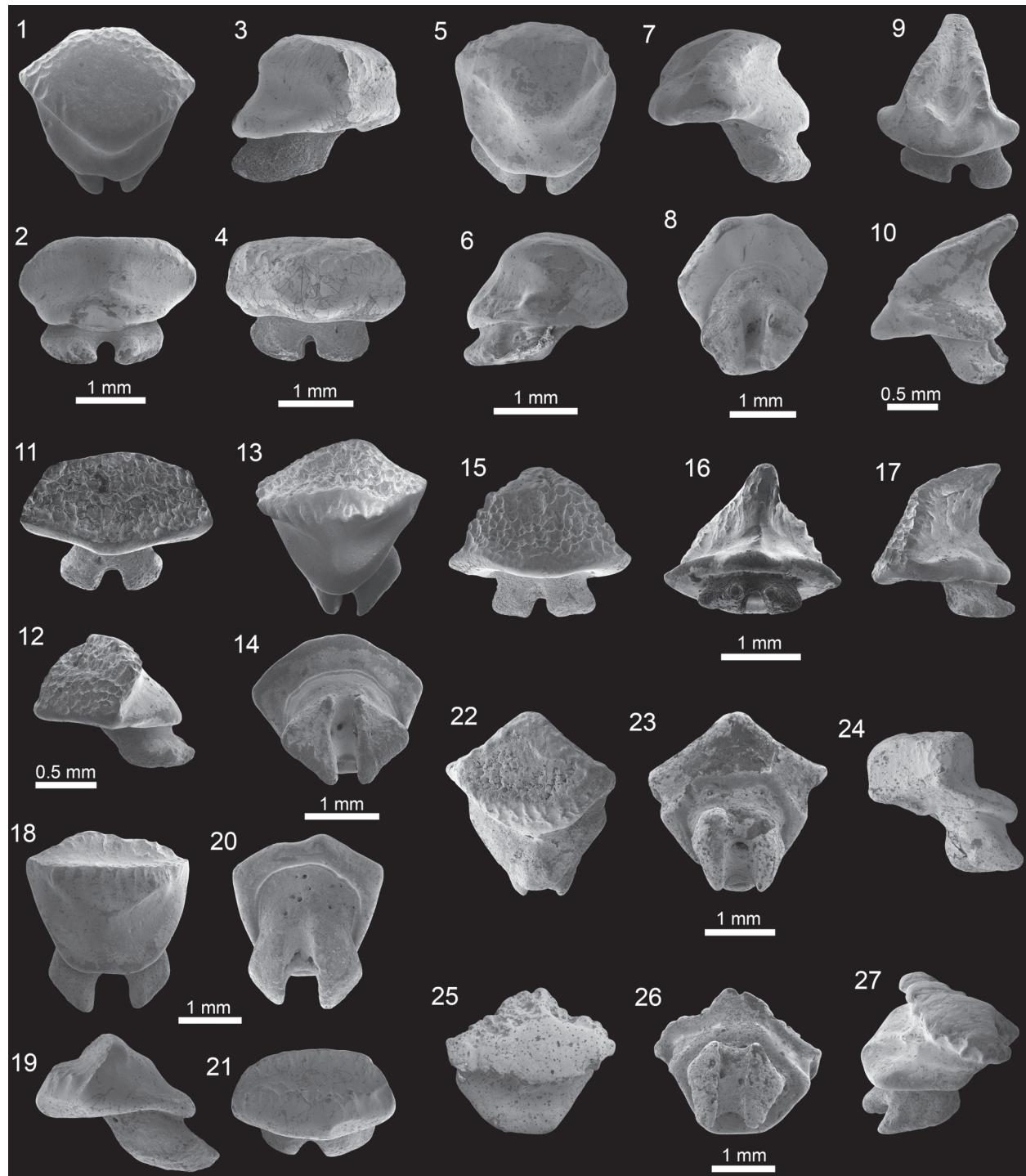


FIGURE 9. Myliobatiformes of the Cantaire Formation. 1-10. *Dasyatis* sp. (1-2, 7-8: AMU-CURS-930; 3-4: AMU-CURS-932; 5-6: AMU-CURS-933 and 9-10: AMU-CURS-931). 11-17. cf. *Pteroplatytrygon* sp. (AMU-CURS-926). 18-21. cf. *Taeniurus* sp. (AMU-CURS-925). 22-27. Dasyatidae indet. (22-24: AMU-CURS-936 and 25-27: AMU-CURS-937). Jaw position: indet. (1-27). View: labial (4, 9, 11, 15, 21), lingual (2, 16), occlusal (1, 5, 13, 18, 22, 25), profile (3, 6-7, 10, 12, 17, 19, 24, 27) and basal (8, 14, 20, 23, 26).

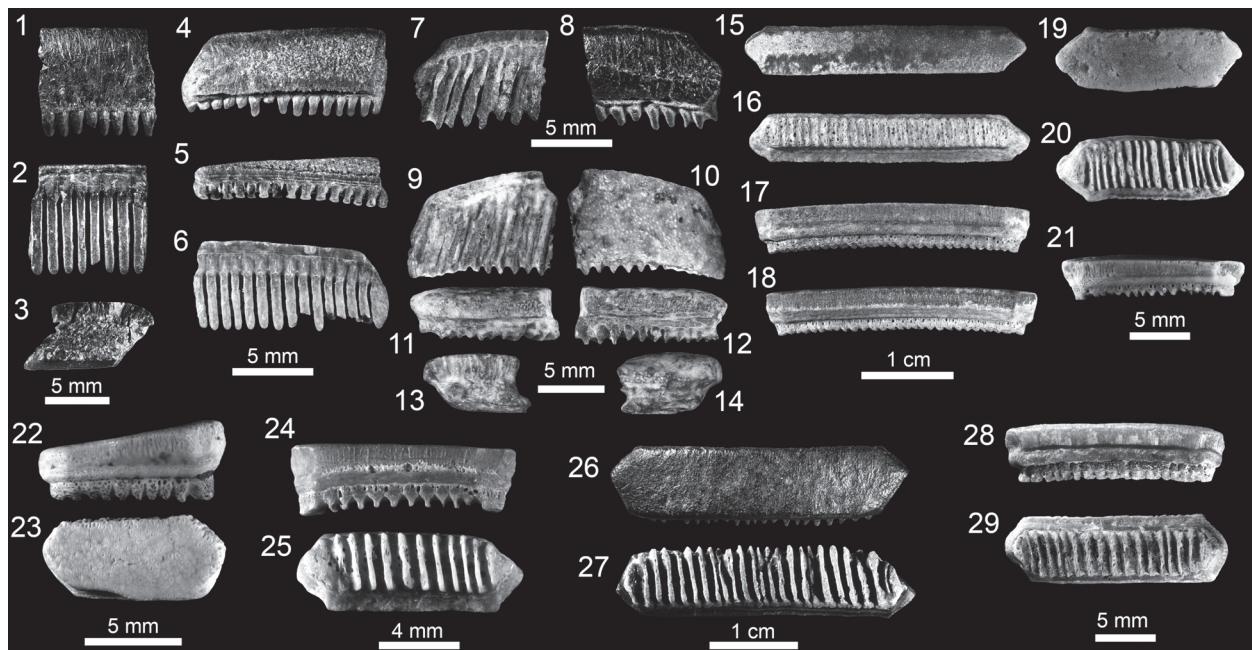


FIGURE 10. Myliobatiformes of the Cantaire Formation. 1-3. *Aetobatus* sp. (1-3: AMU-CURS-716). 4-14. *Aetomylaeus* sp. (AMU-CURS-750). 15-29. *Rhinoptera* sp. (15-18: AMU-CURS-979; 19-21, 22-23, 24-25, 28-29: AMU-CURS-977 and 26-27: AMU-CURS-981). Jaw position: indet. (1-29). View: labial (11, 18), lingual (5, 12, 17, 21-22, 24, 28), occlusal (1, 4, 8, 10, 15, 19, 23, 26), profile (3, 13-14) and basal (2, 6, 7, 9, 16, 20, 25, 27, 29).

species, and the Myliobatidae Bonaparte, 1838 and Rhinopteridae Jordan and Evermann, 1896 by one genus and one species each (Figure 2; Table 1). Due to the small number of specimens and the lack of diagnostic characters, teeth referred to *Dasyatis* Rafinesque, 1810 (Figure 9.1-9.10), cf. *Pteroplatytrygon* Fowler, 1910 (Figure 9.11-9.17), cf. *Taeniurus* Garman, 1913 (Figure 9.18-9.21), and *Dasyatidae* indet. (Figure 9.22-9.27), cannot be assigned to species. As noted by Carrillo-Briceño et al. (2015b), there is very little known available about the morphological diversity of dental patterns in extant and fossil Dasyatidae from the Americas.

Teeth referred to *Aetobatus* Blainville, 1816 (Figure 10.1-10.3), and *Aetomylaeus* Garman, 1913 (Figure 10.4-10.14), are fragmented and incomplete; but preserve diagnostic characters that made their generic identification possible. Teeth of *Rhinoptera* Cuvier, 1829 (Figure 10.15-10.29) are common in all localities (especially at locality 2), being the third most abundant species of the assemblage (Appendix 1, 4). Previous work on the Cantaire Formation has reported the presence of *Myliobatis* Cuvier, 1816 teeth (Aguilera and Rodrigues de Aguilera, 2004; Aguilera and Lundberg, 2010; Aguilera, 2010). However, our revision of fossils led to the identification of isolated medial

and lateral teeth, which exhibit the typical imbrication pattern and morphology of *Rhinoptera*. Mobulidae Gill, 1893 is represented by two genera and four species (Figure 2; Table 1; Appendix 1, 4). The mobulids †*Mobula fragilis* Cappetta, 1970 (Figure 11.1-11.11), *Mobula* cf. †*M. loupianensis* Cappetta, 1970 (Figure 11.12-11.18) and *Mobula* sp. (Figure 11.19-11.23) are scarce (Appendix 1, 4), and their teeth have been recovered from localities 1 and 2 only. One tooth of †*Plinthicus stenodon* Cope, 1869 was found at locality 3.

All eroded, broken and non-diagnostic caudal spines (Figure 12.1-12.6) are referred to Myliobatiformes indet. In addition, specimens referred to Chondrichthyes indet. correspond to fragmented and non-diagnostic vertebrae (Figure 12.7-12.8).

Elasmobranch Paleobathymetric Analysis

The Cantaire elasmobranch fauna is characterized by a predominance of benthopelagic sharks and rays, the extant relatives of which inhabit diverse environments and have diverse bathymetric preferences, especially in neritic/epipelagic environments (Appendix 3). Our paleobathymetric analysis indicates that the most probable deposition depth for the Cantaire assemblage is the 0 to 50 m interval, followed by the 50 to 100 m interval as the second most probable (Figure 13). This sug-

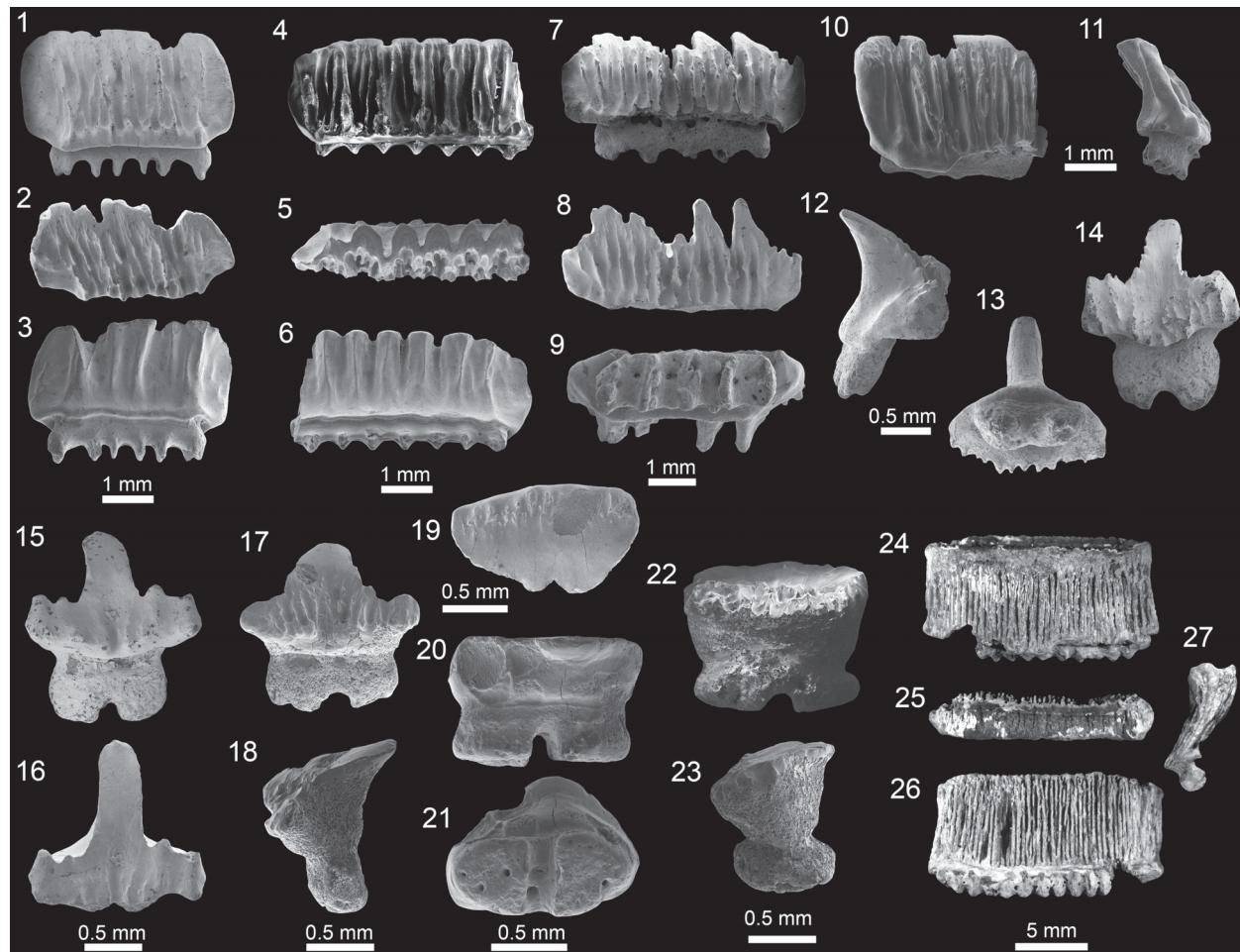


FIGURE 11. Mobulids of the Cantaura Formation. 1-11. †*Mobula fragilis* (1-3, 10-11: AMU-CURS-947; 4-6: AMU-CURS-948 and 7-9: AMU-CURS-949). 12-18. *Mobula* cf. †*M. loupiensis* (AMU-CURS-946). 19-23. *Mobula* sp. (AMU-CURS-927). 24-27. †*Plinthicus stenodon* (AMU-CURS-715). Jaw position: indet. (1-27). View: labial (1, 4, 7, 10, 14-15, 17, 22, 24), lingual (3, 6, 20, 26), occlusal (2, 5, 8, 16, 19, 25), profile (11-12, 18, 23, 27) and basal (9, 13, 21).

gests that the Cantaura Formation was deposited in an inner-middle shelf environment. The presence of *Nebrius* sp., *Carcharhinus* cf. *C. limbatus*, *C. perezi*, *Isogomphodon acuarius*, *Negaprion eurybathrodon*, *Rhynchobatus* sp. and *Rhinoptera* sp. support this hypothesis, due to the fact that their extant representatives inhabit waters shallower than those of middle shelf environments (Appendix 3). Other taxa recognized in the first and second most probable depth intervals (Figure 13) include Alopiidae, Carcharhinidae, Sphyrnidae, Rinchobatidae, Pristidae, Dasyatidae, Myliobatidae and Mobulidae, whose extant representatives are associated with coastal environments. However, it is important to keep in mind that some of those species also have a wide vertical distributional range (Appendix 3), or are even able to move

along significant distances over oceanic basins (Compagno, 1984b; Compagno et al., 2005; Voigt and Weber, 2011; Froese and Pauly, 2015).

Dietary Preferences of the Cantaura Elasmobranchs

The dietary composition of the elasmobranch assemblage from the Cantaura Formation can be represented by five food categories, from filter feeders to benthic-pelagic predators (Figure 14; Appendix 3). In this assemblage the most abundant feeder group is that of the piscivorous, representing 49% of the total species count (Figure 14; Appendix 3). We have included in this group the pelagic lamnid *Paratodus benedenii* (Cappetta, 1987, Kent and Powell, 1999) and the carcharhinid *Physogaleus contortus*, extinct species with a cut-

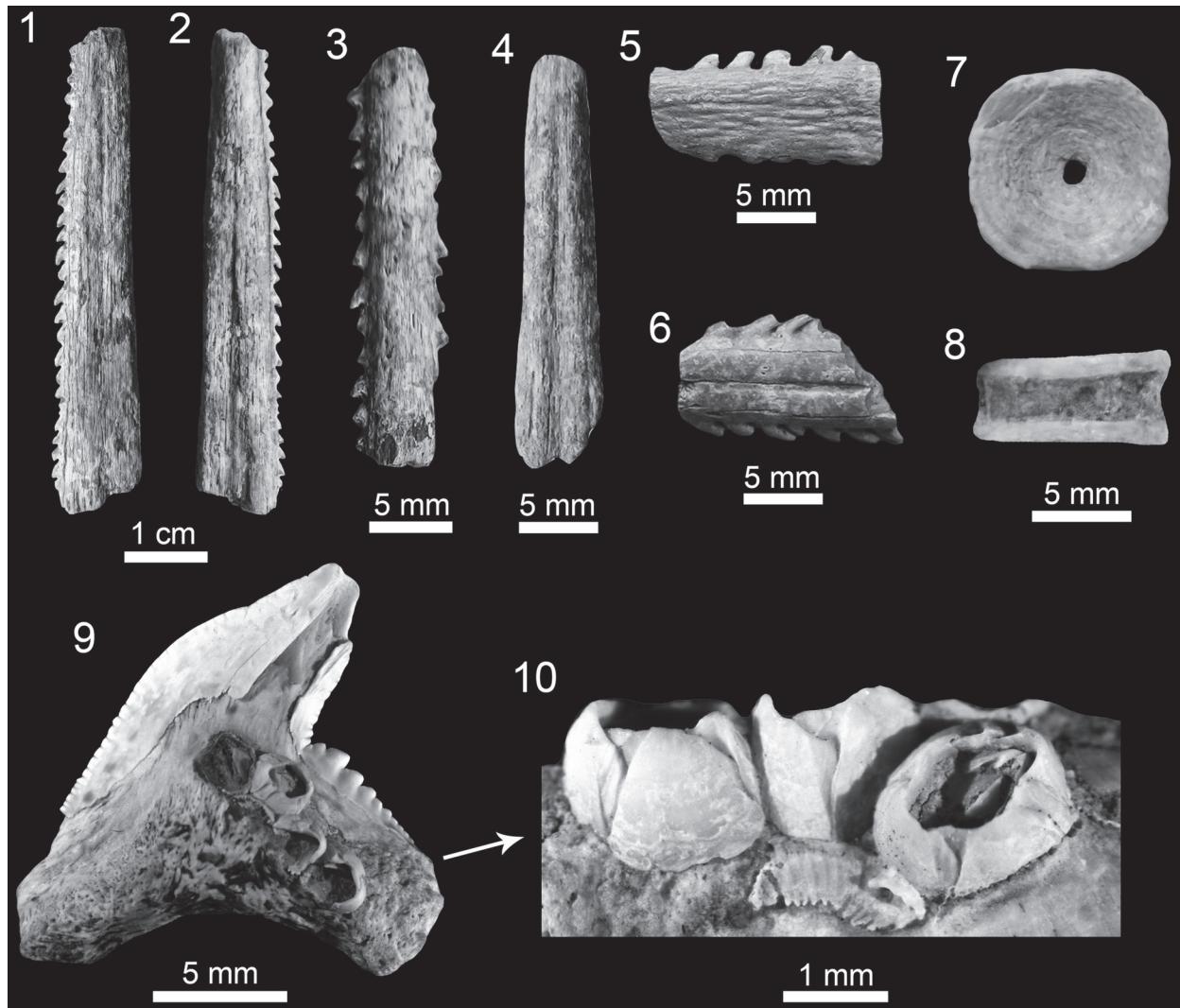


FIGURE 12. Myliobatiformes and other chondrichthyans of the Cantaire Formation. 1-6. Myliobatiformes indet. (AMU-CURS-507). 7-8. vertebra of chondrichthyan Indet. (AMU-CURS-928). 9-10. tooth of *†Galeocerdo aduncus* in association with barnacles of the family Balanidae (AMU-CURS-720). Jaw position: indet. (9). View: lingual (9), dorsal (1, 4, 5-6), ventral (2-3) and indet. (7-8).

ting-clutching dental type that could suggest a fish-based diet. The durophagous/cancritrophic (mollusk, crustacean, coral feeders) group is the second most abundant, corresponding to 30% of the species in the Cantaire assemblage, with a predominance of Myliobatiformes and Rajiformes (Figure 14; Appendix 3). In this group the most abundant taxon is *Rhinoptera* sp., living species of which are opportunistic/generalists consuming mainly mollusks. The filter feeder group (Figure 14) is represented in the Cantaire assemblage only by mobulids (Table 1; Appendix 1, 4), whose living counterparts are characterized by a diet based mainly on planktonic microorganisms (Klimley,

2013). One of its representatives is the extinct mobulid *Plinthicus stenodon*, whose fragile teeth also suggest a possible filter diet of soft prey (Cappetta, 2012). In the eurytrophic/sarcophagous group (diverse food source: fishes, reptiles, birds, mammals, etc.) (Figure 14), we have included large sharks with a broad habitat preference. Those species include *Carcharocles megalodon*, *Galeocerdo aduncus* and *G. mayumbensis*, which could have had the role of apex predators. In reference to the teuthitrophic group, it is represented only by *Alopias* cf. *A. vulpinus*, whose living counterpart has a higher preference for cephalopods (Cortés, 1999; Appendix 3).

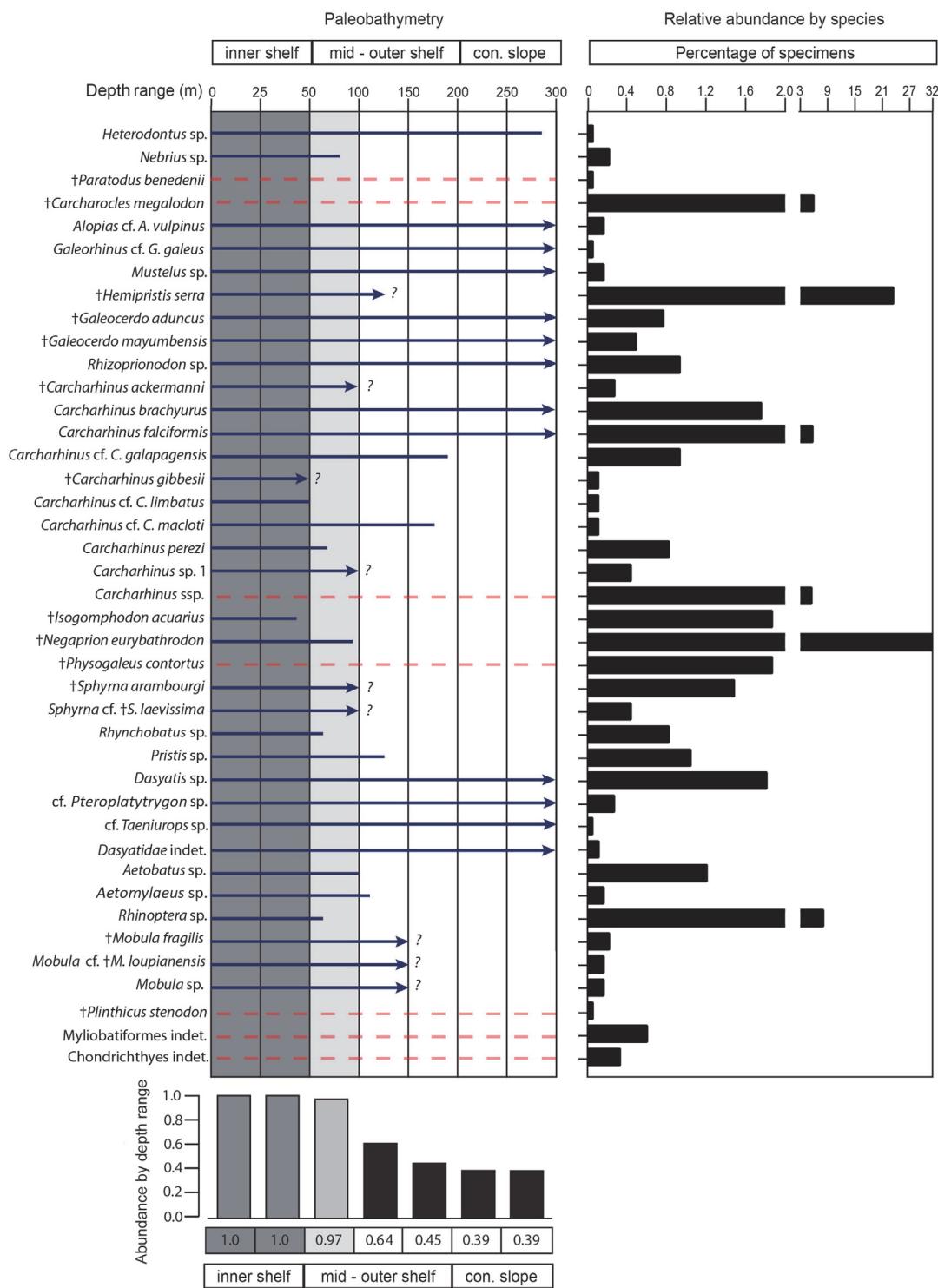


FIGURE 13. Bathymetric ranges estimated for Cantare Formation taxa and their relative abundance. Arrows indicate that the bathymetric range is greater; question mark symbol indicates uncertainty about the extent of the range. A dashed red line indicates that there is no information for the taxa. The gray shadow indicates the more probably depth range for this assemblage. In the relative abundance graphic the y-axis is shortened between 2 and 3 to improve the values visualization; note the greater abundance of shallow water taxa. The extinct species including *P. benedenii*, *C. megalodon*, *P. contortus*, and *P. stenodon*, and taxa without clear identification, including *Carcharhinus* spp., Carcharhinidae indet., Myliobatiformes indet. and Chondrichthyes indet., were removed from the analysis.

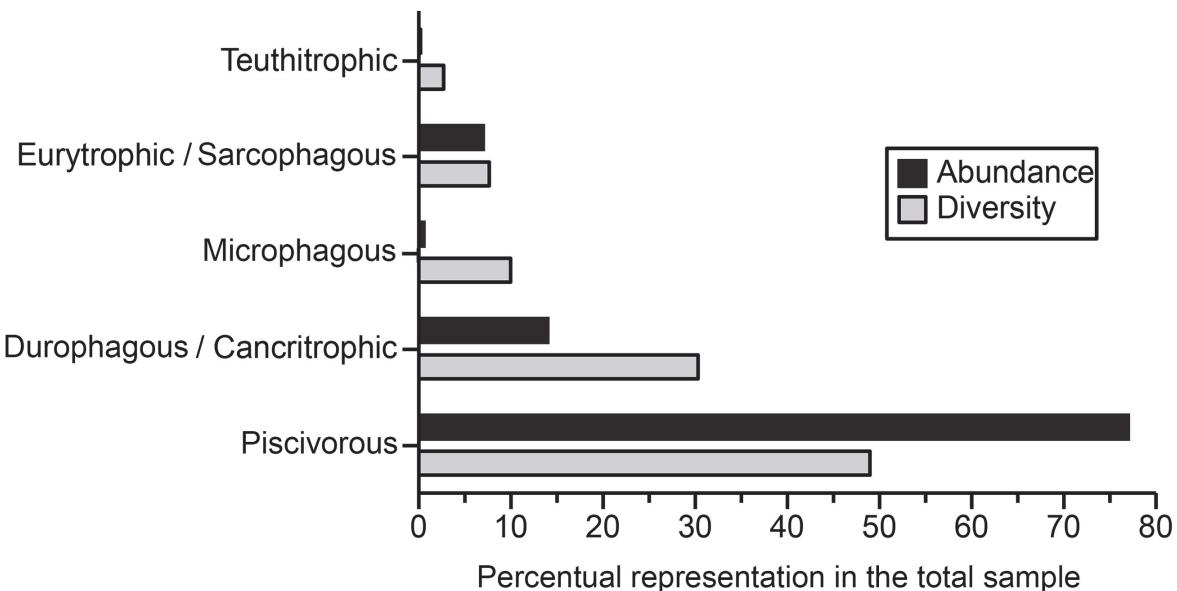


FIGURE 14. Dietary preference composition of the elasmobranchs from the Cantaura Formation. Percentages do not include *Carcharhinus* spp., Myliobatiformes indet. and Chondrichthyes indet.

DISCUSSION

Paleodiversity Significance

Forty-four percent of the Cantaura elasmobranch assemblage consists of genera or species with living representatives in Tropical America (Compagno, 1984a, 1984b; Compagno et al., 2005; Voigt and Weber, 2011). Five genera and 15 species present are extinct (Table 1). *Nebrius* sp., *Carcharhinus* cf. *C. maclovi*, and *Rhynchobatus* sp., are present in our fossil sample and are currently extinct in the Caribbean, but have living counterparts in the eastern Atlantic and Indo-West Pacific (Compagno et al., 2005; Froese and Pauly, 2015). Their fossil record confirms that these taxa became extinct in the Caribbean, Western Atlantic and Eastern Pacific, possibly as a consequence of ecological competition with other species or environmental changes during the final stage of the CAS (e.g., Coates and Stallard, 2013).

The galeomorphs *Paratodus benedenii*, *Galeocerdo mayumbensis*, *C. brachyurus*, *Carcharhinus* cf. *C. galapagensis*, *C. gibbesii*, *Carcharhinus* cf. *C. maclovi*, *Isogomphodon acutus*, *Physogaleus contortus*, *Sphyrna arambourgi*, *Sphyrna* cf. *S. laevissima*, and the batoids cf. *Pteroplatytrygon* sp., cf. *Taeniurops* sp., *Mobula fragilis* and *Mobula* cf. *M. loupi* are reported for the first time from Neogene Venezuelan deposits. *Paratodus benedenii*, *Carcharhinus* cf. *C. gal-*

pagensis, *Carcharhinus* cf. *C. maclovi*, cf. *Pteroplatytrygon* sp., *Taeniurops* sp., *Mobula fragilis* and *Mobula* cf. *M. loupi* are new records for the Neogene of Tropical America. With the exception of the records mentioned above from Tropical America, the fossil record of most elasmobranch taxa from the Cantaura Formation have been found in other Neogene marine deposits of the Americas and are summarized in Kruckow and Thies (1990), Laurito (1999), Purdy et al. (2001), Aguilera et al. (2011), González-Rodríguez et al. (2013) and Carrillo-Briceño et al. (2014, 2015a, 2016).

Early Neogene marine elasmobranch assemblages known from the Americas are summarized in Figure 15 and include North America (Purdy et al., 2001; Visaggi and Godfrey, 2010), Southern South America (Suarez et al., 2006; Cione et al., 2011) and Tropical America (Leriche, 1938; Casier, 1966; MacPhee et al., 2003; Alván et al., 2006; Portell et al., 2008; Costa et al., 2009; Aguilera and Lundberg, 2010; Pimiento et al., 2013b; Laurito et al., 2014; Carrillo-Briceño et al., 2016). From these assemblages the Uitpa Formation (Colombia) has been referred to the Aquitanian (Carrillo-Briceño et al., 2016), while the Calvert and Pungo River formations in North America (Purdy et al., 2001; Visaggi and Godfrey, 2010) and the Cantaura (in this work) and Castillo (Rincón et al., 2014) formations in Venezuela, have been referred to the Burdiga-

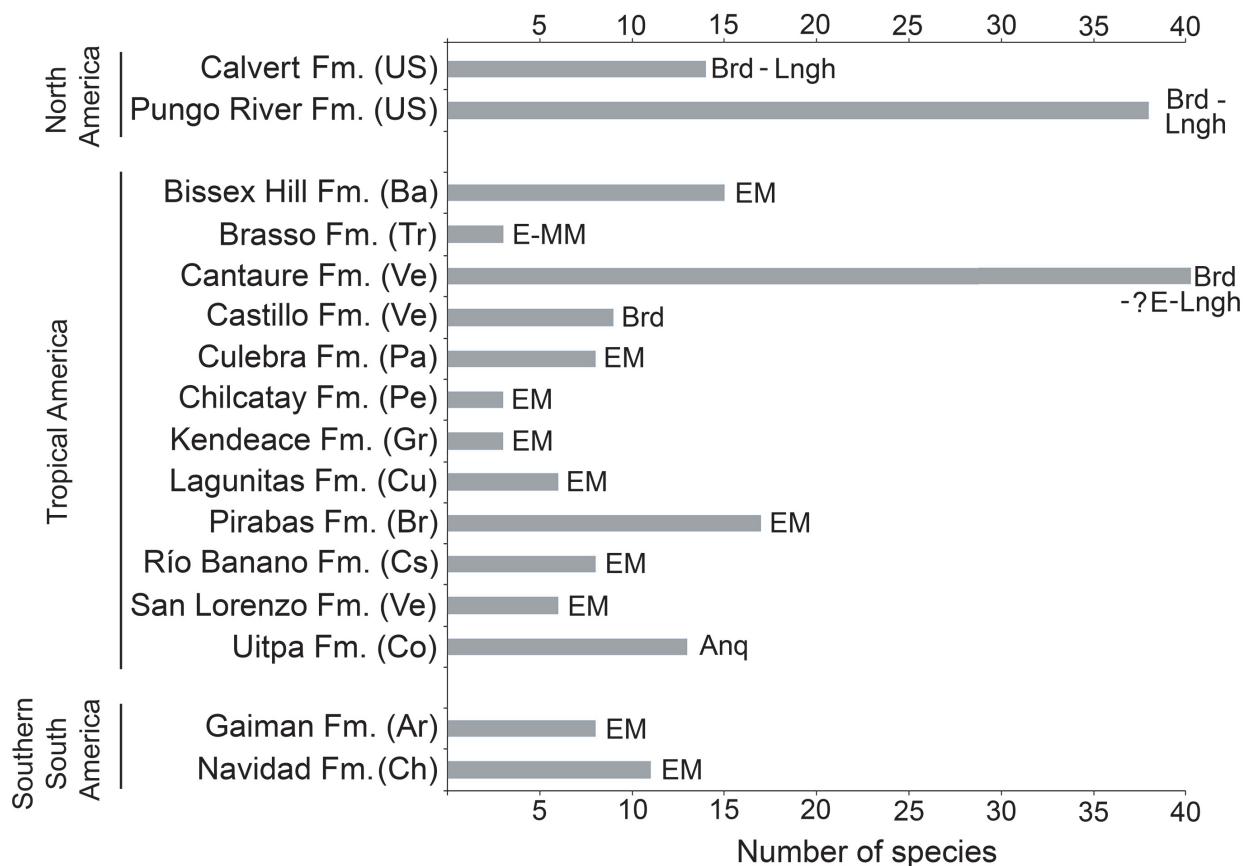


FIGURE 15. Early Neogene chondrichthyan diversity in the Americas. Tropical America: Barbados (Ba), Brazil (Br), Colombia (Co), Costa Rica (Cs), Cuba (Cu), The Grenadines (Gr), Panamá (Pa), Perú (Pu), Trinidad (Tr) and Venezuela (Ve). North America: United States of America (US). Southern South America: Argentina (Ar) and Chile (Ch). Aquitanian (Anq), Burdigalian (Brd), Langhian (Lngh); early Miocene without differentiation (EM), early-Middle Miocene (E-MM). Formation(Fm.), early (E). Note: to references see “Paleodiversity significance” (see Discussion).

lian. The remaining formations (Figure 15) have been assigned to an unresolved early Miocene age. The Pungo River (~38 spp.) and Cantaure (39 spp.) assemblages preserve the most diverse elasmobranch faunas known from early Miocene of the Americas.

The late Burdigalian age (16.5 ± 0.4 Ma.) of the Cantaure Formation based on Sr isotopes from corals from locality 2 of Griffiths et al. (2013) (Figure 1.2), suggests that all specimens found in localities 1 and 2 could be coetaneous or older than 16.5 ± 0.4 Ma. The poorly known stratigraphic position of localities 3 and 4, in reference to localities 1 and 2 (different vertical strata along the Cantaure section or horizontal variation of localities 1 and 2), does not permit an accurate age estimation for them, but their stratigraphic position (underlying the limestones interpreted as post-Cantaure sedimentation; Jung, 1965; Hunter and Bartok, 1974),

suggests that these localities (3 and 4) are close to the top section of the Cantaure Formation. A Burdigalian age (Griffiths et al., 2013), or even that of a Burdigalian-Langhian boundary for the upper section of the Cantaure Formation (assuming that the biozones N7-8 and NN4-5 represent the lower to middle Miocene transition), would extend the stratigraphic range of two galeomorphs and of a batoid found in this assemblage. Fossils of *Carcharhinus* cf. *C. galapagensis*, from the Cantaure Formation, represent the oldest record for this species. It was previously known from the late Miocene of Mexico (González-Rodríguez et al., 2013). The same applies to *Carcharhinus gibbesii*, the youngest record was reported by Carrillo-Briceño et al. (2016) as the Aquitanian of Colombia. In reference to the batoid cf. *Pteroplatytrygon* sp., its record from the Cantaure Formation could be the oldest known to this stingray species.

The presence of *Carcharocles megalodon* in Burdigalian sediments of the Cantaura Formation (in at least localities 1 and 2) confirms the presence of this species during late early Miocene. It has been generally accepted that the record of *Carcharocles chubutensis* spans from the early to the middle Miocene, whereas *Carcharocles megalodon* spans from the middle Miocene to the late Pliocene (Pimiento and Clements, 2014; Pimiento and Balk, 2015). However, other authors have suggested that *C. megalodon* also occurs in the Burdigalian of Europe and North America (Leriche, 1938; Purdy et al., 2001; Visaggi and Godfrey, 2010; Reinecke et al., 2011). According to Leriche (1938), the two *C. megalodon* specimens he collected (Leriche, 1938, Planche III, Figures 5-6) in southern Mexico (Tabasco and Isthmus of Tehuantepec) came from early Miocene strata and are similar to teeth he previously collected in the Burdigalian of Europe. Another early Miocene unpublished occurrence of *C. megalodon* from Tropical America is from the late Burdigalian section (~16.2 Ma.) of the Castilletes Formation at North East Colombia (under study by JDCB). The specimens of *C. megalodon* from the Cantaura sequence have been found in all four localities studied here (Figure 1, Appendix 1), but with a greater abundance in localities 2 and 3 (Appendix 1). The presence of *C. megalodon* tooth (UNEFM-PF-322) at least 6-10 m deep in the artisan well (Figure 1.2) could suggest an older age than the 16.5 ± 0.4 Ma for the species. More detailed information about the American and worldwide Neogene stratigraphic range of the remaining elasmobranch taxa recovered from the Cantaura Formation (Table 1) can be found in Kruckow and Thies (1990), Cicimurri and Knight (2009), Reinecke et al. (2011), Bor et al. (2012), Adnet et al. (2012), Cappetta, (2012), Pimiento et al. (2013b), and Carrillo-Briceño et al. (2014, 2015a, 2016).

Furthermore, it is important to note that the elasmobranch fauna from the Cantaura Formation shows a clear differentiation in paleodiversity amongst the four localities studied here (Appendix 1, 4). As is shown in Appendix 4.2, localities 1 and 2 (Figure 1) have a higher diversity than those of localities 3 and 4, a fact that could be attributable to: 1) less intensive sieving and sampling of the localities 3 and 4 (see Material and Methods section) or 2) different lithologic, taphonomic and preservational conditions. Future, more accurate dating of the deposits could offer new insights into the age of the Cantaura deposits, and especially of the sections located south of Barbasco Creek.

Paleoenvironment and Paleoecology

On the basis of planktonic foraminifera, calcareous nannofossils and mollusks, it has been proposed that the Cantaura Formation was deposited in a shallow, coastal tropical marine environment, with clear waters and normal salinity (Jung, 1965; Díaz de Gamero, 1974; Rey, 1996; Landau et al., in press). A conspicuous absence of typical herbivore gastropods suggests that sea-grass habitats were probably not represented in this environment (Landau et al., in press). The teleostean assemblage reported by Nolf and Aguilera (1998) reflects a shallow environment with a water depth of less than 50 m. The crustacean assemblage (Aguilera et al., 2010) is also suggestive of inner-shelf environments. The barnacle epifauna attached to a shark tooth AMU-CURS-720 (Figure 12.9-12.10) further indicates that the Cantaura Formation was deposited in a shallow marine environment, as barnacles live either in intertidal (26%) or sublittoral areas of ≤ 100 m (73%) (Doyle et al., 1996). Our bathymetric analysis of the Cantaura Formation, suggesting most likely an estimated deep interval depth between 0 to 50 m (Figure 13), supports the assumption of an inner-middle shelf environment (Figure 16). Additional support to this estimation is derived from the presence of the benthopelagic sharks *Nebrius* sp., *Carcharhinus* cf. *C. limbatus*, *C. perezi* and *Isogomphodon acuaris*, whose extant representatives usually prefer waters shallower than 40 m deep (Appendix 3). As shown in Figure 13, many of the benthopelagic and pelagic elasmobranch taxa of the Cantaura fauna have extant representatives associated with coastal environments, but can also occur in adjacent deep waters (Appendix 3), and even move significant distances over oceanic basins, suggesting a wide environmental range (Compagno et al., 2005; Voigt and Weber, 2011).

The Cantaura sequence rests on metamorphic and igneous rocks that were above water level during early Miocene (Gonzalez de Juana et al., 1980; Macellari, 1995). During the Cantaura sedimentation, these positive reliefs played the role of islands or archipelagos, as illustrated in the paleogeographic reconstruction of the Falcon Basin by Gonzalez de Juana et al. (1980) and Macellari (1995). According to Landau et al. (in press), 47.8% of the gastropod species (188 spp.) in the Cantaura Formation are exclusive, representing a very high level of endemism. This species endemism could support the hypothesis of an isolated, insular mollusk-community in the northern part of the Falcon Basin during early Miocene. However,

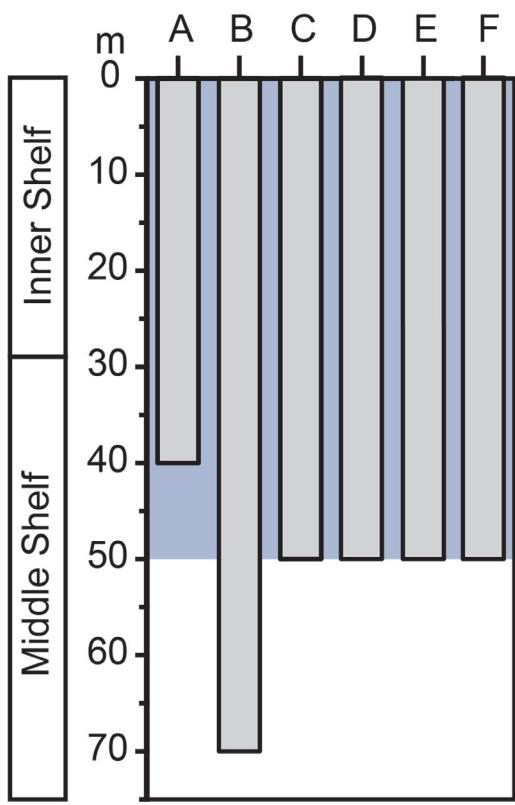


FIGURE 16. Paleobathymetric interpretation of the Cantaure Formation using several proxies. A, Foraminifera (Díaz de Gamero, 1974), B, corals (Griffiths et al., 2013), C, mollusks (Jung, 1965; Landau et al., in press), D, crustaceans (Aguilera et al., 2010), E, sharks (this study), F, teleosteans (Nolf and Aguilera, 1998; Aguilera and Rodrigues de Aguilera, 2001).

expanded knowledge of their localities will certainly lower that estimate. The Cantaure gastropod assemblage also has a small freshwater component, corresponding to typical species of late Oligocene early Middle Miocene NW South-American inland freshwater faunas (Landau et al. in press). This is likely suggestive of insular areas allowing the existence of terrestrial faunas. In addition, the diverse teleostean fauna from the Cantaure Formation, with more than 65 species (Nolf and Aguilera, 1998; Aguilera, 2010; Aguilera and Lundberg, 2010), suggests a high-productivity environment associated with a coastal upwelling (Aguilera and Rodrigues de Aguilera, 2004).

Although it is difficult to approach the precise ecological role that the Cantaure elasmobranch fauna played in this ancient marine environment, we can speculate about trophic interactions using the dietary (nutrition mode) categorization of taxa in the assemblage (Figure 14). The Cantaure

assemblage is dominated by piscivorous shark species (Figure 14, Appendix, 3). This could be related to the diverse teleostean fauna of the Cantaure Formation (Nolf and Aguilera, 1998; Aguilera, 2010; Aguilera and Lundberg, 2010), whose remains (mainly otoliths) are extremely abundant, with approximately more than 60 specimens per 10 kg of sifted matrix (OAA and JDCB pers. obs.). The durophagous/cancritrophic feeders represent the second most abundant elasmobranch group in the assemblage (Figure 14, Appendix 3). Their potential benthonic prey includes some 445 mollusks species (Jackson et al., 1999; Landau et al., in press) and more than 20 crustacean species (Aguilera et al., 2010) and must have offered wide feeding opportunities. Durophagous rays with capacity to triturate hard shells were abundant in the Cantaure fauna (e.g., *Rhinoptera* sp. being the third most abundant species of the assemblage, Appendix 1, 3). The presence of mobulids (Table 1, Appendix 1, 4), whose living counterparts are characterized by a diet based mainly on planktonic microorganisms (Klimley, 2013), support the assumption of a high-productive environment associated to a coastal upwelling (Aguilera and Rodrigues de Aguilera, 2004).

A few large sharks from the Cantaure assemblage such as *Carcharocles megalodon*, *Galeocerdo aduncus* and *G. mayumbensis*, could have alternatively occupied different environments and fed on various resources, such as teleosteans, chondrichthyans, reptiles, birds and mammals. In extant marine communities many sharks are perceived as apex predators, occupying the same trophic level with marine mammals (Klimley, 2013). In the Cantaure fauna such large sharks (*C. megalodon*, *G. aduncus* and *G. mayumbensis*), together with a medium-sized odontocete (?Squalodelphinidae) (Cozzuol and Aguilera, 2008; likely occupied the apex predator role. Moreover, cetacean (squalodontids) and sirenid (dugongids) skeletal remains found in the Cantaure Formation exhibit shark bite marks, suggesting that these mammals also were prey (Cozzuol and Aguilera, 2008). A cetacean lumbar vertebra with an embedded tooth of a *Carcharocles megalodon* was reported from the early Pliocene Paraguáná Formation, in the “El Yacural” locality (Aguilera et al., 2008). Recent field work in the “El Yacural” produced new specimens that include cetacean remains (vertebra and a skull fragment; AMU-CURS-709 and 710, respectively) and a crocodilian vertebra (AMU-CURS-708).

Those specimens lack any evidence for predation. The geographic coordinates of locality 4

(Figure 1.1) are very close to the “El Yacural” locality studied by Aguilera et al. (2008), with a horizontal distance of less than 200 m. Both localities are overlain stratigraphically by limestone layers on the top of the Cerro San Carlos (OAA, JDCB, GAF, pers. obs.), which have been interpreted by Jung (1965) and Hunter and Bartok (1974) as the top of the Cantaure Formation, representing the transition between the latter unit and post-Cantaure and pre-Pliocene sedimentation. We tentatively suggest that the vertebra with an embedded *Carcharocles megalodon* tooth could belong to the Cantaure Formation and not to the early Pliocene Paraguaná Formation (Aguilera et al., 2008). As previously mentioned, the Cantaure Formation is a complex sequence that needs a new detailed stratigraphic redescription, which could offer new geological information. The sections located at the south of the Barbasco Creek are of particular relevance for future research. In addition, the crocodilian vertebra (AMU-CURS-708) and the small eroded turtle-shell fragments, although not offering useful information for taxonomical classification, verify the presence of at least two reptiles in this ancient high-productive marine environment of the Cantaure Formation.

CONCLUSIONS

The elasmobranch fauna from the Cantaure Formation, containing 39 species, is characterized by a predominance of benthopelagic sharks with piscivorous feeding preferences, followed by durophagous/cancritrophic feeders. The elasmobranch habitat preferences and paleobathymetric analyses support the hypothesis that the Cantaure Formation was deposited in a high productivity, inner-middle shelf environment, around positive reliefs (island/archipelago), in the northern part of Falcon Basin during early-?middle Miocene. This shark and ray fauna, is one of most diverse assemblages from early Neogene of the Americas, and contains many forms now extinct in the Caribbean, but present in the eastern Atlantic and western Pacific. Crocodile and turtle remains are also preserved in the Cantaure Formation.

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APPENDIX 1.

Elasmobranchii of the Cantaure Formation and their record per localities, jaw position and tooth measurements. Abbreviations: number of specimens (N°), total number of specimens (N° T) and indeterminate (indet.).

Taxon	Localities								Jaw position			Teeth measurements (mm)			
	1	N°	2	N°	3	N°	4	N°	N° T	Upper	Lower	indet.	Height	Width	Length
<i>Heterodontus</i> sp.	—	—	X	1	—	—	—	—	1	—	—	1	?	?	?
<i>Nebrius</i> sp.	X	3	X	1	—	—	—	—	4	—	—	4	2 to 11	3 to 7	—
† <i>Paratodus benedenii</i>	—	—	X	1	—	—	—	—	1	1?	—	—	32	?	—
† <i>Carcharocles megalodon</i>	X	1	X	18	X	80	X	7	106	50	31	25	23 to 127	23 to ~ 66	—
<i>Alopias</i> cf. <i>Alopias vulpinus</i>	X	1	X	1	X	1	—	—	3	2	—	1	9 to 11	10 to 11	—
<i>Galeorhinus</i> cf. <i>G. galeus</i>	X	1	—	—	—	—	—	—	1	—	—	1	2.5	5	—
<i>Mustelus</i> sp.	—	—	X	3	—	—	—	—	3	—	—	3	<3	<3	<3
† <i>Hemipristis serra</i>	X	2	X	182	X	195	X	46	425	230	70	125	5 to 32	4.5 to 35	—
† <i>Galeocerdo aduncus</i>	X	1	X	13	—	—	—	—	14	—	—	14	11 to 18	16 to 24	—
† <i>Galeocerdo mayumbensis</i>	—	—	X	8	X	1	—	—	9	—	—	9	12 to 22	20 to 25.5	—
<i>Rhizoprionodon</i> sp.	X	3	X	14	—	—	—	—	17	4	6	7	2 to 3.8	2.2 to 4	—
† <i>Carcharhinus ackermannii</i>	X	5	—	—	—	—	—	—	5	3	2	—	11 to 18.2	12.7 to 20	—
<i>Carcharhinus brachyurus</i>	—	—	X	32	—	—	—	—	32	29	3	—	8 to 14	8 to 16.1	—
<i>Carcharhinus falciformis</i>	—	—	X	85	—	—	X	17	102	96	6	—	3 to 17	3 to 15.2	—
<i>Carcharhinus</i> cf. <i>C. galapagensis</i>	—	—	—	—	X	17	—	—	17	—	—	—	12 to 17.5	13 to 16	—
† <i>Carcharhinus gibbesii</i>	—	—	X	2	—	—	—	—	2	2	—	—	9 to 9.9	7.8 to 8	—
<i>Carcharhinus</i> cf. <i>C. limbatus</i>	X	2	—	—	—	—	—	—	2	2	—	—	5 to 9	6.5 to 11	—
<i>Carcharhinus</i> cf. <i>C. maclovi</i>	—	—	X	2	—	—	—	—	2	2	—	—	3.4 to 6.3	4 to 6.7	—
<i>Carcharhinus perezi</i>	X	11	X	3	X	1	—	—	15	15	—	—	4 to 8	5 to 13	—
<i>Carcharhinus</i> sp. 1	X	8	—	—	—	—	—	—	8	8	—	—	2 to 4.5	3.2 to 6	—
<i>Carcharhinus</i> spp.	X	21	X	50	X	17	X	10	98	40	52	6	5 to 15	4 to 14	—
† <i>Isogomphodon acuaricus</i>	X	8	X	26	—	—	—	—	34	18	16	—	3 to 10.1	3.1 to 9.5	—
† <i>Negaprion eurybathrodon</i>	X	13	X	101	X	437	X	35	586	280	306	—	6 to 19	6 to 22	—
† <i>Physogaleus contortus</i>	—	—	X	30	X	3	X	1	34	—	—	34	10 to 19	19 to 19	—
<i>Carcharhinidae</i> indet.	X	2	X	8	—	—	—	—	10	2	8	—	3 to 6.2	4.5 to 8	—
† <i>Sphyraena arambourgi</i>	X	14	X	13	—	—	—	—	27	14	12	1	3 to 7	4 to 7.2	—
<i>Sphyraena</i> cf. † <i>S. laevissima</i>	X	2	X	1	X	5	—	—	8	—	3	5	3.5 to 11.9	5.5 to 14.9	—
<i>Rhynchosciurus</i> sp.	X	8	X	7	—	—	—	—	15	—	—	15	<4	<4	<4
<i>Pristis</i> sp.	X	5	X	13	X	1	—	—	19	—	—	19	<3	<3	<3
<i>Dasyatis</i> sp.	X	21	X	12	—	—	—	—	33	—	—	33	<3	<3	<3
cf. <i>Pteroplatytrygon</i> sp.	X	5	—	—	—	—	—	—	5	—	—	5	<3	<3	<3
cf. <i>Taeniurops</i> sp.	X	1	—	—	—	—	—	—	1	—	—	1	<3	<3	<3
<i>Dasyatidae</i> indet.	X	2	—	—	—	—	—	—	2	—	—	2	<3	<3	<3
<i>Aetobatus</i> sp.	X	6	X	9	X	4	—	—	19	—	—	19	4.1 to 6	?	~4.5 to 10
<i>Aetomyialeus</i> sp.	X	2	X	1	—	—	—	—	3	—	—	3	?	?	?
<i>Rhinoptera</i> sp.	X	15	X	129	X	1	—	—	145	—	—	145	2 to 6	5 to 29	3 to 8
† <i>Mobula fragilis</i>	X	4	—	—	—	—	—	—	4	—	—	4	2.5 to 3	3.5 to > 6	1 to 1.6
<i>Mobula</i> cf. † <i>M. loupianensis</i>	X	3	—	—	—	—	—	—	3	—	—	3	<2	<2	<2
<i>Mobula</i> sp.	—	—	X	3	—	—	—	—	3	—	—	3	<2	<2	<2
† <i>Plinthiscus stenodon</i>	—	—	—	X	1	—	—	—	1	—	—	1	8	15.5	4
Myliobatiformes indet.	—	—	X	11	—	—	—	—	11	—	—	—	?	?	?
Chondrichthyes indet.	—	—	X	6	—	—	—	—	6	—	—	—	?	?	?
Total:	170	786	764	116	1836										

APPENDIX 2.

Referred fossil specimens and collection numbers.

Referred fossil specimens

- *Nebrius* sp. [1 tooth (AMU-CURS-982), 1 tooth (AMU-CURS-986)].
- †*Paratodus benedenii* [1 tooth (AMU-CURS-645)].
- †*Carcharocles megalodon* [4 teeth (AMU-CURS-512), 3 teeth (AMU-CURS-513), 2 teeth (AMU-CURS-514), 2 teeth (AMU-CURS-515), 3 teeth (AMU-CURS-516), 4 teeth (AMU-CURS-517), 3 teeth (AMU-CURS-518), 1 tooth (AMU-CURS-519), 4 teeth (AMU-CURS-650), 1 tooth (AMU-CURS-996), 1 tooth (AMU-CURS-997), 1 tooth (AMU-CURS-998), 7 teeth (AMU-CURS-1034), 1 tooth (UNEFM-PF-321), 1 tooth (UNEFM-PF-322), 13 teeth (UNEFM-PF-s/n), 17 teeth (UNEFM-PF-s/n), 21 teeth (UNEFM-PF-s/n), 17 teeth (MCNC-s/n)].
- *Alopias* cf. *A. vulpinus* [1 tooth (AMU-CURS-983), 1 tooth (AMU-CURS-984), 1 tooth (AMU-CURS-985)].
- *Galeorhinus* cf. *G. galeus* [1 tooth (AMU-CURS-974)].
- *Mustelus* sp. [3 teeth (AMU-CURS-975)].
- †*Hemipristis serra* [66 teeth (AMU-CURS-502), 1 tooth (AMU-CURS-506), 6 teeth (AMU-CURS-520), 24 teeth (AMU-CURS-644), 3 teeth (AMU-CURS-712), 1 tooth (AMU-CURS-731), 1 tooth (AMU-CURS-973), 1 tooth (AMU-CURS-993), 1 tooth (AMU-CURS-994), 1 tooth (AMU-CURS-1014), 46 1 tooth (AMU-CURS-1031), 156 (UNEFM-PF-s/n), 118 teeth (UNEFM-PF-s/n)].
- †*Galeocerdo aduncus* [13 teeth (AMU-CURS-647), 1 tooth (AMU-CURS-730)].
- †*Galeocerdo mayumbensis* [8 teeth (AMU-CURS-646), 1 tooth (AMU-CURS-995)].
- *Rhizoprionodon* sp. [14 teeth (AMU-CURS-959), 3 teeth (AMU-CURS-960)].
- †*Carcharhinus ackermannii* [3 teeth (AMU-CURS-713), 2 teeth (AMU-CURS-714)].
- *Carcharhinus brachyurus* [28 teeth (AMU-CURS-668), 4 teeth (AMU-CURS-990)].
- *Carcharhinus falciformis* [72 teeth (AMU-CURS-669), 6 teeth (AMU-CURS-991), 11 teeth (AMU-CURS-953), 17 teeth (AMU-CURS-1035)].
- *Carcharhinus* cf. *C. galapagensis* [17 teeth (AMU-CURS-973)].
- †*Carcharhinus gibbesii* [2 teeth (AMU-CURS-958)].
- *Carcharhinus* cf. *C. limbatus* [2 teeth (AMU-CURS-944)].
- *Carcharhinus* cf. *C. macloti* [2 teeth (AMU-CURS-957)].
- *Carcharhinus perezi* [1 tooth (AMU-CURS-954), 3 teeth (AMU-CURS-955), 11 teeth (AMU-CURS-956)].
- *Carcharhinus* sp. 1 [8 teeth (AMU-CURS-943)].
- *Carcharhinus* spp. [22 teeth (AMU-CURS-651), 26 teeth (AMU-CURS-652), 14 teeth (AMU-CURS-940), 7 teeth (AMU-CURS-941), 10 teeth (AMU-CURS-942), 7 teeth (AMU-CURS-972), 1 tooth (AMU-CURS-1015), 1 tooth (AMU-CURS-1017), 10 teeth (AMU-CURS-1085)].
- †*Isogomphodon acuarius* [8 teeth (AMU-CURS-950), 25 teeth (AMU-CURS-951), 1 tooth (AMU-CURS-952)].
- †*Negaprion eurybathrodon* [165 teeth (AMU-CURS-503), 65 teeth (AMU-CURS-504), 56 teeth (AMU-CURS-505), 1 tooth (AMU-CURS-509), 47 teeth (AMU-CURS-649), 13 teeth (AMU-CURS-987), 1 tooth (AMU-CURS-998), 52 teeth (AMU-CURS-989), 1 tooth (AMU-CURS-1013), 35 (AMU-CURS-1032), 150 teeth (UNEFM-PF-s/n)].
- †*Physogaleus contortus* [29 teeth (AMU-CURS-648), 2 teeth (AMU-CURS-992), 1 tooth (AMU-CURS-1033), 1 tooth (AMU-CURS-719), 1 tooth (AMU-CURS-720)].

APPENDIX 2 (CONTINUED).

- Carcharhinidae indet. [2 teeth (AMU-CURS-938), 8 teeth (AMU-CURS-939)].
- †*Sphyraena arambourgi* [14 teeth (AMU-CURS-961), 13 teeth (AMU-CURS-962)].
- Sphyraena* cf. †*S. laevissima* [1 tooth (AMU-CURS-963), 5 teeth (AMU-CURS-964), 1 tooth (AMU-CURS-965)].
- Rhynchobatus* sp. [4 teeth (AMU-CURS-653), 2 teeth (AMU-CURS-654), 3 teeth (AMU-CURS-968), 4 teeth (AMU-CURS-969), 1 tooth (AMU-CURS-970), 1 tooth (AMU-CURS-971)].
- Pristis* sp. [7 rostral denticles and 1 rostrum fragment (AMU-CURS-242), 1 rostral denticle (AMU-CURS-717), 3 teeth (AMU-CURS-655), 1 tooth (AMU-CURS-1019), 1 tooth (AMU-CURS-966), 5 teeth (AMU-CURS-717)].
- Dasyatis* sp. [1 tooth (AMU-CURS-657), 1 tooth (AMU-CURS-658), 12 teeth (AMU-CURS-929), 5 teeth (AMU-CURS-930), 10 teeth (AMU-CURS-931), 1 tooth (AMU-CURS-932), 1 tooth (AMU-CURS-933), 1 tooth (AMU-CURS-934), 1 tooth (AMU-CURS-935)].
- cf. *Pteroplatytrygon* sp. [5 teeth (AMU-CURS-926)].
- cf. *Taeniurops* sp. [1 tooth (AMU-CURS-925)].
- Dasyatidae indet. [1 tooth (AMU-CURS-936), 1 tooth (AMU-CURS-937)].
- Aetobatus* sp. [1 tooth (AMU-CURS-508), 1 tooth (AMU-CURS-659), 4 teeth (AMU-CURS-663), 4 teeth (AMU-CURS-665), 6 teeth (AMU-CURS-716), 2 teeth (AMU-CURS-718), 1 tooth (AMU-CURS-729)].
- Aetomylaeus* sp. [1 tooth (AMU-CURS-750), 2 teeth (AMU-CURS-766)].
- Rhinoptera* sp. [80 teeth (AMU-CURS-660), 1 tooth (AMU-CURS-664), 4 teeth (AMU-CURS-977), 6 teeth (AMU-CURS-978), 28 teeth (AMU-CURS-979), 11 teeth (AMU-CURS-980), 15 teeth (AMU-CURS-981)].
- †*Mobula fragilis* [2 teeth (AMU-CURS-947), 1 tooth (AMU-CURS-948), 1 tooth (AMU-CURS-949)].
- Mobula* cf. †*M. loupiensis* [3 teeth (AMU-CURS-946)].
- Mobula* sp. [3 teeth (AMU-CURS-945)].
- †*Plinthicus stenodon* [1 tooth (AMU-CURS-715)].
- Myliobatiformes indet. [10 caudal spines (AMU-CURS-507), 1 tooth (AMU-CURS-510), 1 tooth (AMU-CURS-1018)].
- Chondrichthyes indet. [4 vertebrae (AMU-CURS-511), 2 denticles (AMU-CURS-656)].

APPENDIX 3.

Bathymetric and feeding preferences of the Cantare Formation elasmobranch taxa, based on the biology of their extant relatives. The gray shadow indicates the feeding preferences. Abbreviations: minimum (Mn), maximum (Mx) and meters (m).

Cantare Formation taxa	Extant species	Habitat of living representatives								Feeding ecology					
										Behavior		Diet			
		Benthic	Benthopelagic	Pelagic	Neritic / Epipelagic	Bathyal / Meso-	Bathypelagic	Comments	Filter feeders	Benthic / pelagic predators	Piscivorous	Durophagous / Cancritrophic	Teuthitrophic	Eurytrophic / Sacrophagous	Microphagous
						Mn	Mx								
<i>Heterodontus</i> sp.	<i>Heterodontus</i> spp.	X		X	X	0	280			X	X	X			
<i>Nebrius</i> sp.	<i>Nebrius ferrugineus</i>	X		X		1	70	usually less than 30 m		X	X	X	X		
† <i>Paratodus benedenii</i>	—			X	X	?	?	?		X	X		?		
† <i>Carcharocles megalodon</i>	—	?	X	X	?	0	?				X		X		
<i>Alopias</i> cf. <i>A. vulpinus</i>	<i>Alopias vulpinus</i>		X	X	X	0	550	usually less than 200 m			X		X		
<i>Galeorhinus</i> cf. <i>G. galeus</i>	<i>Galeorhinus galeus</i>	X	X	X	0	1100	usually 2-400 m			X	X	X	X		
<i>Mustelus</i> sp.	<i>Mustelus</i> spp.	X	X	X	0	900				X	X	X			
† <i>Hemipristis</i> serra	<i>Hemipristis elongata</i>	X	X		0	~ 130				X	X		?	?	
† <i>Galeocerdo aduncus</i>	<i>Galeocerdo cuvier</i>	X	X	?	0	~ 370	usually less than 150 m			X	X	X	X	X	
† <i>Galeocerdo mayumbensis</i>	<i>Galeocerdo cuvier</i>	X	X	?	~ 370	usually less than 150 m				X	X	X	X	X	
<i>Rhizoprionodon</i> sp.	<i>Rhizoprionodon</i> spp.	X	X		0	500	usually less than 100 m			X	X	X	X		
† <i>Carcharhinus ackermannii</i>	<i>Carcharhinus</i> spp.	?	X	X	?	0	>100?			X	X				
<i>Carcharhinus brachyurus</i>	<i>Carcharhinus brachyurus</i>	X	X	X	0	~ 360	usually less than 100 m			X	X		X		
<i>Carcharhinus falciformis</i>	<i>Carcharhinus falciformis</i>		X	X	X	0	4000	most in waters > 200 m		X	X		X		
<i>Carcharhinus</i> cf. <i>C. galapagensis</i>	<i>Carcharhinus galapagensis</i>	X	X		0	180				X	X				
† <i>Carcharhinus gibbesii</i>	<i>Carcharhinus</i> spp.	?	X	X	?	0	>100?			X	X		X		
<i>Carcharhinus</i> cf. <i>C. limbatus</i>	<i>Carcharhinus limbatus</i>	X	X		0	100	usually less than 30 m			X	X		X		
<i>Carcharhinus</i> cf. <i>C. maclovi</i>	<i>Carcharhinus maclovi</i>		X	X		0	170			X	X		X		
<i>Carcharhinus perezi</i>	<i>Carcharhinus perezi</i>	X	X		0	65	usually less than 30 m			X	X				
<i>Carcharhinus</i> sp. 1	<i>Carcharhinus</i> spp.	?	?	X	?	0	>100?						X		
<i>Carcharhinus</i> spp.	<i>Carcharhinus</i> spp.	X	X	X	?	0	>100?						X	?	?
† <i>Isogomphodon acarius</i>	<i>Isogomphodon oxyrinchus</i>	X	X		0	40				X	X				
† <i>Negaprion eurybathrodon</i>	<i>Negaprion brevirostris</i>	X	X		0	92				X	X				
† <i>Physogaleus contortus</i>	—	?	X	X	?	?	?			X	X				
<i>Carcharhinidae</i> indet.	<i>Carcharhinidae</i> spp.	?	?	X	?	0	>100?			X	X				
† <i>Sphyraena arambourgi</i>	<i>Sphyraena</i> spp.	X	X	?	0	>100?				X	X				
<i>Sphyraena</i> cf. † <i>S. laevissima</i>	<i>Sphyraena zygaena</i>	X	X		0	>100?	usually less than 20 m			X	X				
<i>Rhynchobatus</i> sp.	<i>Rhynchobatus</i> spp.	X	X		0	~ 60				X	X		X		
<i>Pristis</i> sp.	<i>Pristis</i> spp.	X	X		0	~ 120	most in waters < 20 m			X	X	X			
<i>Dasyatis</i> sp.	<i>Dasyatis</i> spp.	X		X	0	~ 470	most in waters < 50 m			X	X				
cf. <i>Pteroplatytrygon</i> sp.	<i>Pteroplatytrygon violacea</i>	X		X	0	381	usually 1-100 m			X	X	X	X		
cf. <i>Taeniurops</i>	<i>Taeniurops</i> spp.	X		X	10	500	usually 20-60 m			X	X				
<i>Dasyatidae</i> indet.	<i>Dasyatidae</i> spp.	X		X	0	~ 470				X	X	X			
<i>Aetobatus</i> sp.	<i>Aetobatus</i> spp.	X		X	0	~ 100	usually less than 40 m			X	X	X	X		
<i>Aetomylaeus</i> sp.	<i>Aetomylaeus</i> spp.			X	0	110				X	X	X	X		
<i>Rhinoptera</i> sp.	<i>Rhinoptera</i> spp.	X		X	0	< 60	usually less than 30 m			X	X	X			
† <i>Mobula fragilis</i>	<i>Mobula</i> spp.		X	X	?	0	>200	usually less than 100 m	X					X	
<i>Mobula</i> cf. † <i>M. loupianensis</i>	<i>Mobula</i> spp.		X	X	?	0	>200	usually less than 100 m	X					X	
<i>Mobula</i> sp.	<i>Mobula</i> spp.		X	X	?	0	>200	usually less than 100 m	X					X	
† <i>Plinthicus stenodon</i>	—		X	X	?	?	?			X					X
Myliobatiformes indet.	—	?	?	?	?	?	?			?	?				
Chondrichthyes indet.	—	?	?	?	?	?	?								

APPENDIX 4.

Relative abundance of elasmobranchs in the Cantaire Formation. 1. Abundance percentage per locality per taxon. 2. Total abundance per locality. Ht: Heterodontiformes; Or: Orectolobiformes; La: Lamniformes; Ca: Carcharhiniformes; Ra: Rajiformes; My: Myliobatiformes; In: Indeterminate; Loc.: locality; †: extint taxon. *: Percentage does not include *Carcharhinus* spp., Myliobatiformes indet. and Chondrichthyes indet.

