A survey of Pliocene to Mid-Quaternary leaf cuticle from the North Island, New Zealand

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ABSTRACT

This study reports on a survey of dispersed leaf cuticle from a range of late Pliocene to mid-Pleistocene (mostly coastal) lignites in the North Island of New Zealand. With few exceptions, the samples contain plant taxa common in New Zealand today, for example, *Beilschmedia* (Lauraceae) and numerous conifers. However, some of them are outside of their current latitudinal-altitudinal range. The common presence of *Halocarpus* (Podocarpaceae) is noteworthy, as almost all records from the North Island today are from upland regions. There is no clear evidence of *Agathis* (Araucariaceae) having existed further south than its present limit. However, a single find of an araucarian in the Rangitawa Fossil Beds in the southern North Island is notable. Its identity is not clear, but it is possibly *Araucaria* (Araucariaceae), making it the youngest occurrence of this genus in New Zealand. If it is *Agathis*, then it is the most southern Pleistocene macro-record. *Cryptocarya* (Lauraceae) is recorded in an undated peat at Huriwai. It represents an apparent Pleistocene survival of a genus known from the Miocene in New Zealand, but now locally extinct. *Carmichaelia* (Fabaceae) cuticle was common, but unexpected, as it is very rare in the palynological record.

Keywords: epidermis; Lauraceae; Fabaceae; Araucariaceae; Podocarpaceae


INTRODUCTION

Each global glacial-interglacial cycle is effectively a naturally repeating experiment. Cold or dry conditions, or even the direct physical effects of glacial activity, substantially alter the vegetation and soil conditions. In otherwise forested regions, the trees return after each interglacial warming, but due to glacial period extinction or simple chance – the make-up of the forests may differ between interglacials. Vegetation from past interglacial periods is therefore of interest as analogues for the

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Holocene and possible future. However, as atmospheric carbon dioxide levels are now higher than their peaks throughout the Pleistocene, their efficacy as analogues becomes more doubtful. In that respect, earlier analogues must be sought, before 2.1 Ma, or even before 2.7 Ma—the Pliocene (Hönisch et al., 2009).

The sequence of vegetation change which occurred in New Zealand from the Last Glacial Maximum through the Holocene to the present is now understood in reasonable amount of detail. This has mostly been due to a concerted effort by palynologists (e.g., McGlone, 1983, 1988; McGlone et al., 2003) and stratigraphers. Now there is an increasing focus to extend this understanding back through previous Glacial-Interglacial cycles (e.g., Bussell, 1986, 1992; Bussell and Pillans, 1992; Moar and Suggate, 1996; Newnham et al., 1999; Moar and McKellar, 2001). Over this time scale, not only have there been major rearrangements of vegetation communities with respect to relative dominance, but there have been both extinctions, and a few natural introductions to New Zealand as well (e.g., Bussell and Mildenhall, 1990; Kershaw et al., 2000). A detailed knowledge of the timing of these events becomes more important in this time range, as other methods of dating become harder.

There have been a few Quaternary macrofossil studies in New Zealand, which complement the palynological work, for instance McQueen (1954), Couper and McQueen (1954), Newnham and Lusk (1987), and Ogden et al. (1993). However, none of these macrofossil studies have involved studying leaf cuticle. Cuticle is useful because small fragments (i.e., 1mm sq.) have the potential for greater taxonomic precision than using pollen or spores. They may have a typical resolution to the genus or species level, whereas pollen and spores are more typically to the genus or family level. Cuticle may therefore potentially identify extinct species. Several of the Pleistocene sites in the North Island have had pollen of *Acacia* reported (Mildenhall, 1975a, b; Bussell, 1992), a genus that is not now indigenous to New Zealand. If leaf cuticle of *Acacia* was present, it would allow both confirmation of the palynological record and an indication of the extant taxonomic affinities of the fossil within the genus. Cuticle may also identify taxa, which are usually invisible in the palynological record, the ‘classic’ case being Lauraceae, where cuticle may not only identify the family, but the genus as well. Cuticle is also more likely to record local vegetation rather than the regional, wind-dispersed pollen that is thought to dominate many palynological samples. Dispersed cuticle may also be present when there are no obvious signs of macrofossils. A relatively small sample (c. 100–300 g) may be sufficient to include many species of plant. The drawback of cuticle is that it is preserved much more infrequently than pollen, and for practical reasons, sample size (at least for a reconnaissance-level project such as this) is larger and temporal resolution lower than for palynology.

Lignite (or somewhat indurated peat) deposits of Late Pliocene-Early Pleistocene age are preserved in various localities in the North Island of New Zealand, mostly around the coast. In the north west of the North Island many of these isolated occurrences have been collectively called the Kaihu Group (Chappell, 1970) and the palynology has been documented in one of the largest palynological studies in New Zealand (Mildenhall et al., 1992) which incorporated several earlier works (e.g (Pocknall and Waterhouse, 1984; Mildenhall, 1985). Further to the south, the sediments of the Wanganui Basin have achieved a global importance as a detailed record of global sea level change over 2.6 Ma (Fleming, 1953; Pillans, 1994; Naish and Kamp, 1995; Carter and Naish, 1998a; Kamp et al., 2004; Pillans, 2017). The sediments are largely marine but include several lignite/peat horizons (Fleming, 1953). Previous macrofossil investigation of these peats includes that of McQueen (1954), Couper and McQueen (1954), and Newnham and Lusk (1987). The former paper described an extinct species of Proteaceae fruit, *Proteoides ohukaensis* from the mouth of Ohuka Creek. The third paper compared macrofossils with the palynological spectra.

The aim of this work is to investigate some of these lignites and associated muddy sediments for plant macrofossils, which have cuticle preservation. The initial goal, which stimulated this project, was to try and locate cuticle belonging to the plant that produced the *Acacia* pollen reported by Mildenhall (1975a, b) and Bussell (1992). The paper offers a photographic record of the cuticle of a range of extant New Zealand taxa that will be a useful reference for further work.

**MATERIAL AND METHODS**

The study focussed on sites known to be Pliocene or earlier Pleistocene in age (Figure 1). Dating in most cases is broad and based on a mixture of palynology, inferred sea-level, and direct dating of volcanigenic sediment or carbonaceous material.
Division of Pliocene-Pleistocene time in New Zealand (Figure 2) follows Raine et al. (2015), which largely follows Beu (2001). This does not recognise several terms that occur frequently in the earlier palynological literature and therefore to help the reader, some of the system of Carter and Naish (1998b) is also included.

The localities were visited over multiple trips from the mid-1990s. Sample size was about 400 g, and where there was no obvious change in lithology, they were taken across the full thickness of a carbonaceous bed (up to around 1 m in some cases). Grid references are given in New Zealand Transverse Mercator (NZTM). Appropriate New Zealand Geological Society Fossil Record Numbers given in Mildenhall et al. (1992) are also cited (e.g., 'P7 f9583').

Disaggregation of samples was usually achieved by immersing the sample in hot water and adding 45 % hydrogen peroxide. If this worked the remaining sludge was then sieved through an approximately 1 mm mesh and the fines discarded. Adhering silica was removed using 70 % hydrofluoric acid. Lignite that did not disaggregate in peroxide was placed directly into c. 10 % chromium trioxide or concentrated hydrogen peroxide and warmed gently. This process reduced any leafy plant material to translucent cuticle, which could then be stained with safranin. Cuticle fragments were then mounted on either microscope slides using thymol glycerine jelly or on scanning electron
microscope stubs with double-sided tape and coated with platinum. Microscope slides are catalogued with the prefix ‘SL’ or ‘LX’ and electron microscope stubs with ‘S-’. Slides were viewed under Transmitted Light Microscopy (TLM) or with Scanning Electron Microscopy (SEM). Comparative herbarium material is prefixed with OTA (Otago University Botany Department, CHR (the Allan Herbarium), and OPH (material collected by the author).

Results are discussed by location stratigraphically below, and taxonomically in the Systematic Palaeontology section.

Material is lodged in the Queensland Herbarium, Brisbane.

**RESULTS**

A. **Kaihu Group**

i. **Formation undifferentiated**

Scottys Camp Road (P7/f9579)

**Location.** Roadside cutting 0.75 km SE of intersection of State Highway 12 and Scottys Camp Road (grid reference: NZTM 1674159, 6024386).

**Age.** Hautawan (late Pliocene) or possibly earliest Marahauan (earliest Pleistocene) (Mildenhall, 1985).

**Palynology.** The pollen assemblage is dominated by Myrtaceae (incl. *Eucalyptus*) and *Acacia* pollen is present (Mildenhall, 1985).

**Macroflora (present study).** Includes *Dacrycarpus, Dacrydium, Beilschmiedia, Carmichaelia*.

Dargaville (P9/f9582)

**Location.** Roadside cutting 1.5 km NW of Dargaville (grid reference: NZTM 1676256, 6025190).

**Age.** Hautawan (late Pliocene) or possibly earliest Marahauan (earliest Pleistocene) (Mildenhall, 1985).

**Palynology.** The pollen assemblage is dominated by Myrtaceae (incl. *Eucalyptus*) and *Nothofagus s.g. Fuscospora. Acacia* pollen is abundant (Mildenhall, 1985; Mildenhall et al., 1992).

**Macroflora (previous work).** *Beilschmiedia taraire* leaves were reported by J. Braggins (personal commun., to Mildenhall, 1985).

**Macroflora (present study).** The dominant plant to be an unidentified monocot. Remains of other taxa are uncommon, but include *Dacrycarpus dacrydioides, Beilschmiedia, and Carmichaelia*.

Baylys Beach (P7 f9583)

**Location.** Coastal cliffs at the end of Baylys Coast road (grid reference: NZTM 1667272,6020671).

**Age.** Possibly early Pleistocene (Richardson, 1985).

**Palynology.** A palynological sample 6.4 km to the NW (P7/f9581) is dominated by *Nothofagus “fusca”* group and *Dacrydium cupressinum*. No extinct pollen taxa were found, and it was regarded as Castlecliffian or younger (Mildenhall, 1985).

**Macroflora (previous work).** C.J. Burrows (personal commun., to Richardson, 1985, p. 118) reported that P7/f9583 “Contains seeds of an apparently extinct plant.” C.J. Burrows and J. Braggins (personal commun., to Mildenhall, 1985) reported that P7/f9581 contained leaves of *Dacry-
Phyllocladus trichocarpus, Dacrydium cupressinum, Podocarpus totara, Agathis, Cyathodes, and Nothofagus truncata, along with wood of Dacrydium colensoi, Agathis and Myrtaceae, and seeds of Elaeocarpus dentatus, Eleocharis sphacelata, and D. cupressinum.

**Macroflora (present study).** Overwhelmingly dominated by Agathis australis leaves and cone scales, with Dacrydium cupressinum also present.

### i. Ohuka Formation

**Pourewa (R15/f73, 74)**

**Location.** Exposed at the edge of a beach at the high tide level, Pourewa Point, Aotea (grid reference: NZTM 1760484, 5791809).

**Age.** Hautawan-Marahauan (Late Pliocene – Early Pleistocene). Chappell (1970) referred to this as Pourewa Lignite, a term that Mildenhall et al. (1992) recommended be given member status of the Ohuka Formation.

**Palynology.** Pocknall and Waterhouse (1984) noted a dominance of Nothofagus s.g. fuscospora pollen and that there were no extinct species and correlated the Pourewa Lignite with the Te Kowiwi Lignite. Mildenhall et al. (1992) recorded similar palynological assemblages but also noted the extinct spore Polypodiisporites radiatus Pocknall and Mildenhall (1984).

**Macroflora (present study).** No stomatiferous leaf cuticle was found.

**Te Kowiwi (R15/f83, 84)**

**Location.** Shore exposure, west side of Kowiwi Bay, directly below power lines (grid reference: NZTM 1760185, 5790908).

**Age.** Mildenhall et al. (1992) regard this assemblage as Hautawan (late Pliocene). Pocknall and Waterhouse (1984) referred to this as Te Kowiwi Lignite, a term that Mildenhall et al. (1992) recommended be given member status of the Ohuka Formation.

**Palynology.** Pocknall and Waterhouse (1984) recorded the dominant pollen taxa as Nothofagus s.g. Fuscospora, Dacrydium, Phyllocladus, Prumnopitys, and Libocedrus. The extinct taxon Polypodiisporites radiatus Pocknall and Mildenhall was also recorded. This has a known upper limit of basalt Nukumaruan (The base of the Nukumaruan is about 2.5 Ma, Pliocene; Beu, 2001). Mildenhall et al. (1992) recorded similar assemblages with one including a large fraction of Dacrycarpus.

**Macroflora (present study).** This study found the macroflora to be overwhelmingly dominated by small and imbricate-leaved conifer shoots, predominantly of Halocarpus, and also Dacrycarpus and Libocedrus. Only two angiosperm taxa are present, Phormium and Pseudowintera. The macroflora suggests the original vegetation was a conifer-dominated swamp. However, the conifer leaves and shoots were mostly very twisted and distorted, perhaps indicating some disturbance before final burial.

**Waiuku (R12/f45, 46)**

**Location.** An exposure at water level in an inlet at the southern end of Manukau Harbour at the end of Racecourse Road, Waiuku (grid reference: NZTM 1753646, 5877698).

**Age.** Marahauan-lower Haweran; Late Pliocene-late Pleistocene (Mildenhall et al., 1992; See also Byrami et al., 2005).

**Palynology.** Mildenhall et al. (1992) placed this exposure in the Ohuka Formation, noting a dominance of Podocarpaceae pollen, but including no extinct species.

**Macroflora (present study).** The cuticle assemblage consists entirely of Beilschmiedia.

**Ihumatao**

**Location.** Peaty material with abundant wood and stumps is exposed by wave action on a coastal platform, (grid reference: NZTM 1755940, 5903759).

**Age.** A range of dates have been obtained by several researchers, including from radiocarbon from c. 29,000 to c. 43,600 years BP (Ferguson and Rafter, 1959; Grant-Taylor and Rafter, 1963; Polach et al., 1969), K-Ar dates on basaltic material of c. 74,000 (McDougall et al., 1969) and 38,000 years BP using thermoluminescence (Wood, 1991). Following this uncertainty, a focussed study including Optically Stimulated Luminescence dating by Marra et al. (2006) regarded the forest as much older—being deposited “during late MIS 7 (probably substage 7.1)” (c. 190,000-210,000 BP). Ricketts (1977) described this as the “Ihumatao Fossil Forest” and concluded it had been buried by volcanic tuff (See also Hayward, 1977). Hayward and Hayward (1995) recognised two forests, one growing within peat and a further one on the peat. They mapped the location of in situ stumps and larger fallen logs. The forest was subsequently investigated for insects and palynology (Marra et al., 2006).

**Macroflora (previous work).** Ricketts (1977) noted Agathis and cf. Dacrydium cupressinum. Hayward and Hayward (1995) listed Dacrydium cupressinum, Elaeocarpus dentatus, Prumnopitys ferrugineus, Agathis australis, Phyllocladus tricho-

**Macroflora (present study).** Found the peat around the lower forest to be dominated by Agathis australis, but also with Halocarpus, Beilschmiedia, Pseudowintera, and Toronia.

**Hamiltons Gap**

**Location.** Beds of highly compressed leaves and other carbonaceous matter occur in high cliffs to the south of Hamiltons Gap, (grid reference: from NZTM 1739935, 5888430 to 1739958, 5888394).

**Age.** The Hamiltons Gap material appears to be similar to that reported by Newnham and Lusk (1987) from about 15 km to the south and 1.5 km north of Karioitahi Beach, (grid reference> R12 564 357).

**Macroflora (previous work).** Newnham and Lusk's (1987) macrofossil material included Laurelia novae-zelandiae, Alectryon excelsus, Mysine salicina, Beilschmiedia taraire, B. tawa, Macropiper excelsum, Elaeocarpus sp., Cyathea sp., Griselinia sp., Podocarpus sp. (identified in the field only), and indeterminate monocot leaves. They noted the abundance of Podocarpaceae and Myrtaceae pollen, although corresponding macrofossils were, consecutively, rare and absent.

**Macroflora (present study).** The material is distinctly lithified and does not disaggregate well with peroxide treatment. However, chromium trioxide treatment showed fragmented cuticle was abundant and well-preserved, and includes Griselinia, Olearia, and Astelia.

**Huriwai**

**Location.** Peat in coastal cliff (grid reference: NZTM 1751200, 5854900).

**Age.** Unknown.

**Macroflora (present study).** The cuticle assemblage consists entirely of Pseudowintera, Beilschmiedia, and Cryptocarya.

**Hauroto Bay**

**Location.** Coastal exposure (grid reference NZTM 1771840, 5814030).

**Age.** Hautawan (Mildenhall, 1992).

**Macroflora (present study).** The cuticle assemblage consists mostly of Beilschmiedia.

**Pahurehure**

**Location.** Peaty material is exposed in the low cliff bordering Pahurehure Inlet a few meters west of the Auckland Motorway. (grid reference NZTM 1770980, 5896670).

**Age.** Unknown.

**Macroflora (present study).** Contains Dacrycarpus, Dacrydium, Libocedrus, Beilschmiedia, and the two extant New Zealand Proteaceae; Toronia toru and Knightia excelsa.

**Ohawe**

**Location.** Bussell (1990) recognised two broad localities: ‘Ohawe Waterfall’ samples came from within 24 below the Stratford Lahar and fell into his OH4 zone, and ‘Ohawe East’ samples came from just above the Stratford Lahar and fell into his OH5 zone. The present study found one productive sample from the ‘Ohawe Waterfall’ locality (my sample Ohawe-4, from lignite directly below the Stratford Lahar, grid reference: NZTM 1704360, 5616740) and two samples from ‘Ohawe East’ (my sample Ohawe-5, from the lower main lignite, and, separated by a c. 30 cm overlying tuff, sample Ohawe-6, from the upper main lignite, grid reference NZTM 1704700, 5616710).

**Age.** Bussell (1990) dated the waterfall section as Oxygen Isotope Stage 6 (c. 130,000-190,000 years BP, full glacial) and Ohawe east as OI5e (c. 110,000-130,000 years BP, full interglacial), reported by Bussell and Mildenhall (1990).

**Palynology.** Bussell (1990) interpreted the Water Fall samples as representing a rush-fern shrubland with sedges and common Phormium, but with patches of shrubland including Halocarpus. In contrast, he considered that Ohawe East samples probably represented forest with prominent Dacrycarpus and Syzygium with wetland taxa in patches.

**Macroflora (present study).** Ohawe-4 contains Phormium and Carmichaelia, Ohawe-5 contains Cyperaceae and Ohawe-6 contains Halocarpus.

**Rapanui**

**Location.** Exposure in coastal cliff (grid reference: NZTM 1764460, 5581563).

**Age.** Bussel (1992), with reference to work of Pillans (1983, 1988, 1990) estimated the age of the Rapanui Lignite as c. 110,000 years BP.

**Palynology.** Bussel (1992) interpreted as “low forest/shrubland with prominent Myrtaceae, and including a now-extinct Acacia-type.”
Macroflora (present study). Dominated by Halo-carpus and Dacrydium with some Dacrycarpus and Prumnopitys.

Omokoroa

Location. (grid reference NZTM 1869491, 5831249).

Age. Briggs et al. (2005). Note the Te Puna Ignimbrite (0.93 Ma) overlies lignites at Omokoroa.

Macroflora (present study). Contains Halo-carpus and Phyllocladus.

B. Wanganui Basin

i. Maxwell Group

The basic stratigraphy of the Wanganui Basin was documented by Fleming (1953), who described many marine units, but amongst these were three formations, which he described as containing 'lignite' - the Upper Maxwell, Middle Maxwell, and Lower Maxwell Formations. These crop out in the coastal cliffs east and west of the mouth of Ototaka Ck, south of Maxwell. The palynology of these three units was documented by W.F. Harris and R.A. Couper (in Fleming, 1953). Some major quantitative differences were noted between samples, but no extinct taxa were recorded.

Age. The Maxwell Formations were deposited from just before the Plio-Pleistocene boundary (1.81 Ma) and approximately 1.6 Ma (Marahauan, early Pleistocene; Carter and Naish, 1998b).

Palynology. Harris and Couper's (in Fleming, 1953) palynological analysis of two 'lignite' samples within the Formation suggested Nothofagus-Weinmannia forest, and podocarp-dicotyledonous forest. No extinct taxa were noted.

Macroflora (present study). Contains sparse Dacrydium, Dacrycarpus and Cyperaceae.

Upper Maxwell Formation

Location. From 2100 m southeast of the Ototoka Stream mouth (sample Maxwell-13, grid reference: NZTM 1758686, 5584805).

Palynology. Harris and Couper's (in Fleming, 1953) palynological analysis of two 'lignite' samples within the Formation suggested Nothofagus-Weinmannia forest, and podocarp-dicotyledonous forest. No extinct taxa were noted.

Macroflora (present study). No good cuticle.

Middle Maxwell Formation


Palynology. Harris and Couper's (in Fleming, 1953) palynological analysis of a 'lignite' bed suggested a Weinmannia-Nothofagus forest. No extinct taxa were noted.

Macroflora (present study). No good cuticle.

Lower Maxwell Formation

Location. From 830 m west of the Ototoka Stream mouth (sample Maxwell-17, grid reference: NZTM 1755840, 5585629).

Palynology. Harris and Couper's palynological analysis of three 'lignite' beds included within the Lower Maxwell Formation noted three distinct assemblages which they identified with Nothofagus-Weinmannia forest, Podocarp-dicotyledonous forest, and Metrosideros-podocarp forest. No extinct taxa were noted.

Macroflora (present study). Contains sparse Dacrydium, Dacrycarpus and Cyperaceae.

Rangitawa Fossil Beds

Location. The Rangitawa Fossil Beds (RFB) was first described by Te Punga (1952, 1962) and more detailed work followed by McQueen (1953), Bussell (1986), and Kohn et al. (1992). For the present study unambiguous RFB was sampled (samples Rangitawa-7, and Rangitawa-11) from outcrop close to normal flow level of the Rangitawa Stream (thus low in the local section which has a very gentle-dip). The peaty material also contains molluscs, as per original descriptions of the RFB, but may be 100 m downstream of where Bussell (1986) indicated samples on his map, (grid reference: NZTM 1809596, 5554665).

Age. Kohn et al. (1992) suggest the Rangitawa Fossil Beds (RFB) date to a little after the start of OIS 11, an interglacial, making them around 420,000 yrs BP. They date the overlying Rangitawa Pumice to about 350,000 yrs BP, which correlates with OIS 10, a glacial period.

Palynology. Bussell (1986) described the palynology of the RFB, and interpreted a “warm-temperate paleoclimate” under “lowland podocarp-hardwood-beech forest.”

Macroflora (previous work). Leaf and seed fossils were described from the Rangitawa Stream by McQueen (1953, 1954) and indicate forested conditions. Two extinct seed species were described; Podocarpites toleyi (Podocarpaceae) and Rumex pachyperianthus (Polygonaceae).

Macroflora (present study). Contains Carmichaelia and Beilschmiedia.

Rangitawa Pumice

Location. Roots, logs and plant macrofossils were reported by Bussell (1986) in mudstone associated with the Rangitawa Pumice about 17 m above the RFB and about 70 m downstream of his RFB sam-
ples. In the present study, what is regarded as the same horizon was located, exposed in a cliff, about 80 m upstream of Bussell’s RFB. The horizon is inaccessible, and therefore fossil material was collected from fallen blocks directly below the cliff (samples Rangitawa-8, -9, -10, Grid reference: NZTM 1809761, 5554759).

**Age.** Kohn et al. (1992) date the Rangitawa Pumice to about 350,000 yrs BP, which correlates with OIS 10, a glacial period.

**Palynology.** The palynology of the sediments associated with the Rangitawa Pumice was documented by Bussell (1986) and later by Kohn et al. (1992, as their ‘Pole Site 2’). These results indicated open, non-forested conditions of grassland-shrubland, similar to a “Present day subalpine environment” (Bussell, 1986).

**Macroflora (present study).** Contains *Beilschmiedia* and *Carmichaelia*.

**Rangitawa ‘Upstream’**

**Location.** In situ peaty material was collected near water level in the left bank of the Rangitawa River (sample Rangitawa-6, grid reference NZTM 1809990, 5554740). This is 300–400 m upstream and along strike of known RFB outcrop and is a likely correlate.

**Macroflora (present study).** Contains the conifers *Halocarpus*, *Libocedrus*, *Dacrydium*, *Phyllocladus*, *Prumnopitys*, and the angiosperms *Carmichaelia* and *Pseudowintera*.

**Rangitawa ‘Float’**

**Location.** Two blocks of ‘float’ fossiliferous mudstone were collected from the Rangitawa stream bed downstream of Rangitawa Pumice outcrop and of probable RFB outcrop. It is uncertain what horizon they derive from (samples Rangitawa-4 and Rangitawa-5, grid reference NZTM 1809750, 5554700).

**Macroflora (present study).** Contains the conifers *Araucaria* or *Agathis*, *Libocedrus*, *Phyllocladus*, and angiosperms *Carmichaelia* *Beilschmiedia*, and *Pseudowintera*.

**C. Uncertain unit**

**Mokau**

**Location.** In 1996 a collection was made from a c, 2 m thick carbonaceous mud horizon exposed in a fresh road-cutting of State Highway c. 3 km south of the bridge over the Mokau River. A return visit to clarify the stratigraphy in 2005 could not relocate it, as it appears to have been covered by a small slump. (grid reference: NZTM 1740577, 5712664).

**Stratigraphy and age.** Henderson and Ongley (1923, p.54) mentioned Early Pleistocene “soft sandy silt containing plant-remains” … “on the coast half a mile south of the Mokau River”. It is possible that this refers to similar material, although the locality collected here is inland.

**Macroflora (present study).** Contains *Dacrycarpus*, *Halocarpus* and an indet. conifer, and the angiosperms *Carmichaelia*, *Griselinia*, *Litsea*, *Pseudowintera*, and *Phormium*.

**DISCUSSION**

This survey demonstrates the existence of well-preserved leaf cuticle in some of the Plio-Pleistocene peat and associated sediments in the North Island (Figure 3). Both angiosperm and conifer cuticle is often abundant. Due to the sometimes higher taxonomic resolution of cuticle compared with pollen, this type of evidence makes it a useful contribution to the field.

The presence of a now-extinct taxon of Lauraceae, *Cryptocarya*, and perhaps *Araucaria*, in the Quaternary, is notable. However, these extinct taxa are rare finds. Overall the taxa recovered as cuticle indicate interglacial vegetation that was little different from the swamp-forest (with common conifers) and associated lower vegetation (*Phormium*-dominated) that grew around the coast of the North Island in Holocene times. The cuticle evidence allows the identification of *Beilschmiedia*, and sometimes *Litsea*, both locally common components of North Island forest today (with *Beilschmiedia* extending to the north-east of the South Island; (Wardle, 1991), but essentially absent from the palynological record.

The repeated occurrence of the conifers *Halocarpus* and *Libocedrus* confirms palynological evidence for these taxa. The occurrences are significant as they are outside of their current geographical range. Both of these taxa have restricted, upland distributions in the North Island today, although they occur at sea level much further to the south, in the South Island. *Libocedrus* and *Halocarpus* are, for example, important conifers today on Mount Hauhungatahi above 1000-1150 m (Atkinson, 1981; Druitt et al., 1990), but *Halocarpus* especially, is almost never below this level in the central North Island. The occurrence of both taxa at present sea level in the North Island suggests cooler conditions.

With the single exception of a possible fragment from Rangitawa, *Agathis* cannot be confirmed in any of the Plio-Pleistocene samples south of this range (i.e., southern North Island samples).
If the Rangitawa Araucariaceae is *Araucaria*, then it is the youngest known occurrence in New Zealand. If it is *Agathis*, then it is far south of its current southern limit of 38.11° S (Ecroyd, 1982) or its earlier Holocene range (Ogden et al., 1992). The general absence of *Agathis* macrofossils south of its current range suggests temperatures may not have been warmer than the present day for those periods. *Agathis* has been reported further south in some interglacials (e.g., Mildenhall, 1986) but based on pollen, which is difficult or impossible to distinguish from *Araucaria*. In this context the possible record of *Araucaria* cuticle from Rangitawa is significant.

The coexistence of both of New Zealand’s extant Proteaceae species, *Knightia excelsa* and *Toronia toru*, at Pahurehure is unusual, but this combination does occur in various regions of the North Island today. *K. excelsa* is a very widespread forest tree of the North Island (and extends a little into the South Island) while *T. toru* is more restricted. Its southern limit is the Bay of Plenty, but
it occurs from lowland to mountain forest (Dawson and Lucas, 2000). It is common in Agathis australis forest at Waipoua, Northland (McGregor, 1948, p.22), where it is described as a “distinctively northern tree”. Grove and Brandon (2006) reported both Proteaceae species in the foothills of the Coromandel Range, near Tairua. They specifically reported T. toru only from the upper slopes and ridge top of the generally secondary vegetation.

The common presence of Libocedrus and Winteraceae amongst the plant remains reinforces the findings of Pole (2007a) in Pliocene localities in the South Island. However, other taxa are noticeable by their general absence or rarity in this survey. Among the conifers for instance, Phyllocladus might be expected, given its importance today in relatively extreme environmental situations, but it was found only in two localities (Rangitawa and Rapanui). There are few records of Phyllocladus growing near sea level in the Auckland region today, although it is common in the uplands. However, all the extant sea-level records appear to be Phyllocladus trichomanoides, while it is unclear to which species the fossils belong to.

The widespread presence of Carmichaelia in these samples was surprising. McGlone et al. (1984) recorded a trace of Carmichaelia pollen in the late Pleistocene of the Inaha Terrace (near the Ohawe site), but in general the genus appears to be severely under-represented in the palynological record of these peats. However, the genus is known from swamps, even tidal swamps (Heads, 2017), and one species, Carmichaelia arborea, is known as ‘Swamp broom’. The cuticle comes from cladodes (it was initially suspected as the source plant of the Acacia pollen, another cladode-bearing plant), but despite not being shed like normal leaves, the remains clearly find their way into the fossil record with some ease. In this respect, the absence of Carmichaelia from any of the diverse early Miocene cuticle assemblages is notable (personal observ.).

The macrofossils found by Ogden et al. (1993) make an interesting comparison to those reported here. For example, the Lauraceae that are virtually ubiquitous in the present samples, were absent from theirs. Conversely, their material included Phyllocladus, although it is uncommon in samples studied here.

There is a clear discrepancy between Bussell’s (1986) palynological results of the Rangitawa Pumice, which indicated grassland-shrubland and the “logs, and plant macrofossils” indicated on his figure 2. The Beilschmiedia cuticle recovered in this work suggests forest, consistent with the logs. This is difficult to reconcile, but might, for example, result from long-distance river transport of forest debris from an isolated patch of forest in sheltered conditions upstream. However, for most of the cuticle in this study, its presence in peat, or at least very carbonaceous muds, suggests little fluvial transport, which would dilute the organic fraction with mud or sand. On that basis, the cuticle assemblages are inferred to represent local vegetation (Ferguson, 1985; Greenwood, 1991; Spicer, 1991; Gastaldo and Staub, 1999).

The broad composition of these samples – composed mainly of extant taxa plus a few, rare taxa destined to become extinct sometime within the Pleistocene, supports the conclusions of palynology (Bussell and Mildenhall, 1990).

As fossil cuticle samples the present material is of relatively low diversity. The most diverse sample was Rangitawa-6, with 11 taxa. Typically there are fewer taxa. In contrast, early Miocene dispersed cuticle assemblages from New Zealand typically contain more, and sometimes over 20 species (e.g., Pole, 2008a). Furthermore, Plio-Pleistocene samples typically have 1-2 species of Lauraceae, comparing well with present-day forest in the northern half of New Zealand, but significantly less than early Miocene where samples, which often have more than 10 Lauraceae species (Pole, 2007b). Both these observations confirm a relatively low diversity of the Plio-Pleistocene source vegetation compared with early Miocene and is consistent with some aspect of temperature being generally cooler than then.

**SYSTEMATIC PALAEONTOLOGY**

Illustrations of selected specimens of all taxa found in this survey are provided here. For most of these taxa a range of illustrations have been published previously and therefore in these cases only basic TLM are given here.

**Conifers**

*Araucariaceae*

*Agathis australis*  
Figure 4.1-2

**Specimens.** SL4401, Baylys Beach; SL5882, Ihamatao.

**Identification.** Agathis litter (as leaves and cone scales) is common at Baylys Beach and Ihamatao. Under TLM the stomatal complexes show the typical 'spoked' structure of the genus (Pole, 2008b).

**Araucaria**  
Figure 4.3-4
Specimens. SL4641, Rangitawa-5.
Identification. A single cuticle fragment from Rangitawa shows araucarian stomatal structure with oblique orientation. However, the 'spoked' structure of the stomatal complexes under TLM is not apparent, and it is possible that this is an *Araucaria* rather than an *Agathis*.

Cupressaceae
*Libocedrus* sp.
Figure 5

Specimens. SL4395, Te Kowiwi; SL4503, Pahurehure; SL4648, Rangitawa-6.
Identification. Monocyclic stomates and scattered papillae and scale-like leaves indicate *Libocedrus*. However, the two extant New Zealand species, *L. plumosa* and *L. bidwilli* are difficult to distinguish without whole shoot data. Florin and Boutelje (1954) noted that *L. plumosa* had papillae on normal epidermal cells, away from the stomatal bands at the base of the leaves. The present fossils display a range of papillae prominence, some appear to lack the fused papillate ring around the stomatal aperture, and one specimen is at least partially dicyclic.

Podocarpaceae
*Dacrycarpus dacrydioides*
Figure 6.1-2

Specimens. SL0202, Dargaville; LX2735, Maxwell-13; LX2837, Maxwell-17; SL2345, Mokau-1; SL4501, Pahurehure; SL4667, Rapanui; SL0225, Scottys Camp; SL4392, Te Kowiwi.
Identification. The very smooth and relatively thin-walled, unbuttressed epidermal cells of this genus are distinctive.

*Dacrydium cupressinum*
Figure 6.3-4
FIGURE 5. Cuticle of Libocedrus sp. (Cupressaceae); 1. SEM view of inner surface, note lateral subsidiary cells from adjacent stomatal complexes touching, and encompassing a shared polar subsidiary cell (S-1651, Rangitawa-6, scale bar equals 10 µm); 2. SEM view of outer surface (S-1651, Rangitawa-6, scale bar equals 20 µm); 3. TLM view (SL4504, Pahurehure, scale bar equals 50 µm); 4. TLM view, note prominent rim of fused papillae and numerous individual papillae (SL5448, Rangitawa-6, scale bar equals 50 µm); 5. TLM view (SL4503, Pahurehure, scale bar equals 50 µm); 6. TLM view, note some dicyclic stomatal complexes (SL4508, Pahurehure, scale bar equals 50 µm).

Specimens. SL5972, Baylys; SL4509, Pahurehure; SL4669, Rapanui; SL5879, Ihamatao; SL2370, Max-07; SL0221, Scottys Camp; LX2836, Maxwell-17.

Identification. The cuticle morphology of Dacrydium is broadly like that of Dacrycarpus, except that the epidermal cells can be buttressed (Florin, 1931). However, the buttressing of the New Zealand Dacrydium cupressinum can be subtle and much less than other species in this genus (see Wells and Hill, 1989). Within New Zealand, D. cupressinum can be consistently distinguished from
**Dacrycarpus dacrydioides** by its thicker and rougher epidermal cell walls and generally less elongate stomatal complexes.

*Halocarpus bidwilli*

**Specimens.** SL5974, Baylys; SL5878, Ihamatao; SL2360, Mokau; SL5965, Omokoroa; LX2293, Ohawe-6; SL4636, Rangitawa-4; SL4388, Te Kow-iwi.

**Identification.** New Zealand currently has three genera of podocarps that may display small, spirally oriented, scale-like leaves. These include *Halocarpus*, *Manoao*, and *Lepidothamnus*. The stomatal construction of all specimens in this study is consistent with *Halocarpus*, almost certainly *H. bidwilli*. These features include a prominent ring of subsidiary cells, which usually have distinctly thicker periclinal walls than normal epidermal cells, and this thickening is consistent from the outer margin to almost the edge of the guard cells, which are surrounded by a ring of thin cuticle (see figures in Florin, 1931; Wells and Hill, 1989; Jordan et al., 2011).

*Phyllocladus* sp.

**Specimens.** SL5968, Omokoroa; SL4980, Rangitawa-6.

**Identification.** Cuticle sheets with multiple rows of well-spaced, paratetraacytic stomatal complexes with lappets and well-developed Florin ring (Cookson and Pike, 1954; Hill, 1989). There are three extant species of *Phyllocladus* in New Zealand: *P. alpinus*, *P. toatoa*, and *P. trichomanoides*, but based on cuticle alone, these species are probably not distinct.

*Prumnopitys taxifolia*

**Specimens.** SL4511, Pahurehure.
Specimens. SL4981, Rangitawa-6; SL4675, Rapanui.

Identification. The specimens have the typical Prumnopitys stomatal distribution whereby stomates are widely-spaced along closely-spaced rows, making the nearest neighbour of a stomate usually in an adjacent row. They also have the typical stomates of this genus which have an "inflated"
FIGURE 9. Cuticle of Prumnopitys taxifolia (Podocarpaceae, all Rangitawa-6) and indet, conifer (all Mokau): 1. *P. taxifolia*, TLM view (SL4981, Rangitawa-6, scale bar equals 50 µm); 2. *P. taxifolia*, TLM view, detail of stomatal complex (SL4981, Rangitawa-6, scale bar equals 20 µm); 3. *P. taxifolia*, SEM view of inner surface, note marked ‘bulging’ of stomatal complex outline (S-1656, Rangitawa-6, scale bar equals 10 µm); 4. *P. taxifolia*, SEM view of outer surface (S-1656, Rangitawa-6, scale bar equals 10 µm); 5. Indet. conifer, TLM view (SL2355, Mokau, scale bar equals 50 µm); 6. Indet. conifer, TLM view (SL2355, Mokau, scale bar equals 50 µm); 7. Indet. conifer, SEM view of inner surface, note clear buttressing (S-1657, Mokau, scale bar equals 20 µm); 8. Indet. conifer, SEM view of outer surface (S-1657, Mokau, scale bar equals 20 µm).
appearance. Of the two extant New Zealand species, they match *P. taxifolia* (compare figures in Greenwood et al., 2013)

Conifer gen indet.
Figure 9.5-8

**Specimen.** SL2355, Mokau.

**Identification.** A single specimen has epidermal cells with sinuous anticlinal cell walls. Stomatal complexes longitudinally-aligned, widely-spaced (no clear row system), with a basically paratetracyclic construction but inconsistent. There are similarities to *Prumnopitys*, but also to what was described as 'Cycad gen. et sp. indet' in Pole (2007c).

Angiosperms
Monocots
CUT-Mo-FGD
*Phormium tenax*
Figure 10

**Specimens.** SL205, Dargaville; LX2731, Maxwell-13; SL2353, Mokau; LX2253, Ohawe-4; SL4649, Rangitawa-6; SL2325, Te Kowiwi.

**Identification.** The epidermis has alternating stomatal and non-stomatal zones. The stomatal complexes are aligned and tend to be arranged in an alternating fashion, the complexes are tightly-packed, and sometimes networked, so that some polar subsidiary cells are the lateral subsidiary cell of an adjacent complex. Subsidiary cells have several papillae. Some specimens are markedly less papillate, much less so than the *Phormium tenax* figured in Pole (2007d). It is not clear whether this is natural variation within the species, or is a different taxon (see also Conran et al., 2015, figure 6).

*Astelia*
Figure 11.1-4

**Specimens.** SL5978, Dargaville; SL5913 Hauroto; SL5436, Hamiltons Gap.

**Identification.** *Astelia* has widely-spaced stomata, hidden under pairs of adjacent trichome bases (see Pole, 2007d; Maciunas et al., 2011; Conran et al., 2015).

*Cyperaceae*
Gen et sp. Indet.
Figure 11.5-6

**Specimen.** SL2376, Maxwell-07; LX2838, Maxwell-17; LX2274, Ohawe-5.

**Identification.** A single fragment of monocot cuticle from Maxwell shows stomatal complexes with pairs of distinct elongate and rounded subsidiary cells, and sinuous epidermal cells. This is regarded as *Cyperaceae*, genus unknown (cf. Conran et al., 2015, figure 8).

*Areucaceae*
Gen et sp. Indet.
Figure 11.7-8

**Specimens.** SL4671, Rapanui.

**Identification.** A single specimen shows broad palm morphology. However, it appears distinct from *Rhopalostylus sapida*, the single extant New Zealand palm (see Pole, 2007d).

*Dicots*

*Winteraceae*

*Pseudowintera colorata*
Figure 12

**Specimens.** SL5457, Huriwai; SL2382, Maxwell-07; SL2361, Mokau; SL5452, Rangitawa-6.

**Identification.** The paracytic stomates and very granular cuticle clearly place this cuticle in the Winteraceae where it can be compared directly with extant New Zealand *Pseudowintera colorata* (see Pole 2007a, figure 11).

*Lauraceae*

*Beilschmiedia* sp.
Figures 13, 14, 15.1-4

**Specimens.** SL0207, Dargaville; SL5959, Hauroto; SL4977, Huriwai; SL5883, Ihumatao; SL0211, Scottys Camp; SL4638, Rangitawa-4; SL2590, Waiuku; SL4495, Pahurehure.

**Identification.** Lauraceae cuticle is relatively easy to identify as such based on the presence of paracytic stomates overarched by subsidiary cells, and separated from them by cuticular “scales”, and, unless glabrous, by simple deciduous trichomes (Hill, 1986). Most of the Lauraceae cuticle found in this survey can mostly be compared directly with extant New Zealand *Beilschmiedia*. Depending on the authority followed, there are two or three *Beilschmiedia* species (Figure 13) in New Zealand (Wright, 1984) and also *Litsea calicaris* (Figure 14). *Beilschmiedia tawa* and *B. tawaroa* are probably not distinguishable on cuticular details. Both have somewhat angular stomatal complexes with truncated poles, thinner cuticle than normal epidermal cells, and narrow cuticular scales. On the external surface the outlines of the epidermal cells and subsidiary cells are clearly marked, and the stomatal pore is broad. *B. tararoides* is similar, except the external surface is much more subdued. The outlines of the epidermal cells are not obvious, and only the area of the subsidiary cells around the narrow stomatal pore is raised. In all *Beilschmiedia* the epidermal cell flanges are very clear in TLM. Thus,
FIGURE 10. Cuticle of *Phormium tenax* (Asphodelaceae): 1. TLM view (SL4983 Rangitawa-6, scale bar equals 200 µm); 2. TLM view (SL4983 Rangitawa-6, scale bar equals 100 µm); 3. TLM view (SL4649, Rangitawa-6, scale bar equals 50 µm); 4. TLM view, detail of stomatal complexes (SL4649, Rangitawa-6, scale bar equals 20 µm); 5. SEM view of inner surface (S-1655, Rangitawa-6, scale bar equals 10 µm); 6. SEM view of outer surface, note marked alternation of stomatal complexes (S-1655, Rangitawa-6, scale bar equals 10 µm); 7. TLM view (SL2353, Mokau, scale bar equals 50 µm); 8. TLM view, detail of stomatal complexes (SL2353, Mokau, scale bar equals 20 µm).
FIGURE 11. Cuticle of *Astelia* (Asteliaceae), Cyperaceae and Arecaceae; 1. *Astelia* sp., TLM view (SL5436, Hamiltons Gap, scale bar equals 50 µm); 2. *Astelia* sp., TLM view (SL5436, Hamiltons Gap, scale bar equals 20 µm); 3. *Astelia* sp., SEM view of inner surface (S-1661, Hamiltons Gap, scale bar equals 50 µm); 4. *Astelia* sp., SEM view of outer surface (S-1661, Hamiltons Gap, scale bar equals 50 µm); 5. Cyperaceae gen. et sp. indet., TLM view (SL2376, Maxwell-07, scale bar equals 50 µm); 6. Cyperaceae gen. et sp. Indet., TLM view, detail of stomatal complexes (SL2376, Maxwell-07, scale bar equals 20 µm); 7. Arecaceae gen. et sp. indet., TLM view (SL4671, Rapanui, scale bar equals 100 µm); 8. Extant *Rhopalostylus sapida*, TLM view (OPH2638, scale bar equals 50 µm).
B. tawa/tawaroa and B. tararoides might be distinguishable using an SEM, though the difference is subtle.

Most of the fossil Lauraceae cuticle can be referred to Beilschmiedia, and most of that to B. tawa/tawaroa, particularly on the clearly defined cells on the outer surface. Material from Rangitawa is illustrated.

Litsea sp.
Figure 15.5-8

Specimens. SL2344, Mokau; SL4982, Rangitawa-6.

Identification. Litsea calicaris is distinct from New Zealand Beilschmiedia, having a much smoother and more defined stomatal outline, raised ledges around the outer stomatal pore, very subdued cuticular scales and a granularity to the inner and outer surface. The epidermal cell flanges can appear almost diffuse in TLM. Rangitawa has material comparable with Litsea calicaris, although at this stage it is prudent to refer to it as Litsea sp.

Cryptocarya sp.
Figure 16

Specimens. SL4977, Huriwai.

Identification. At Huriwai there is lauraceous cuticle that differs from extant New Zealand Beilschmiedia and Litsea in having well-developed, "butterfly-like", cuticular scales, and a granular cover over the subsidiary cells and stomatal pore. The most likely identity of this is Cryptocarya (Christophel and Rowett, 1996), a genus now extinct in New Zealand.

Proteaceae
Toronia toru
Figure 17.1-2

Specimens. SL5880, Ihamatao; SL4498, Pahurehure.
FIGURE 13. Cuticle of extant *Beilschmiedia* (Lauraceae): 1. *Beilschmiedia tawaroa*, TLM view (CHR367126, scale bar equals 50 µm); 2. *B. tawaroa*, TLM view, detail of stomatal complex (left) and trichome attachment site (right) (CHR367126, scale bar equals 20 µm); 3. *B. tawaroa*, SEM view of outer surface (S-1784, CHR367126, scale bar equals 10 µm); 4. *B. tawaroa*, SEM of inner surface SEM view (S-1784, CHR367126, scale bar equals 10 µm); 5. *B. tawa*, TLM view (OTA01439, scale bar equals 50 µm); 6. *B. tawa*, TLM view, detail of stomatal complexes (OTA01439, scale bar equals 20 µm); 7. *B. tawa*, SEM view of outer surface with clearly defined stomatal complexes and slit-like apertures (S-1719, OPH2627, scale bar equals 10 µm); 8. *B. rawa*, SEM view of inner surface with two stomatal complexes (S-1719, OPH2627, scale bar equals 10 µm).
FIGURE 14. Cuticle of extant *Beilschmiedia* and *Litsea* (Lauraceae); 1. *Beilschmiedia tarairi*, TLM view (OTA46493, scale bar equals 50 µm); 2. *B. tarairi*, TLM view, detail of two stomatal complexes (OTA46493, scale bar equals 20 µm); 3. *B. tarairi*, SEM view of outer surface (S-1718 OPH2628, scale bar equals 10 µm); 4. *B. tarairi*, SEM view of inner surface (S-1718 OPH2628, scale bar equals 10 µm); 5. *Litsea calicaris*, TLM view (OTA04767, scale bar equals 50 µm); 6. *L. calicaris*, TLM view, detail of stomatal complex, note very rounded outline (OTA04767, scale bar equals 20 µm); 7. *L. calicaris*, SEM view of outer surface (S-1153, OTA04767, scale bar equals 10 µm); 8. *L. calicaris*, SEM view of inner surface (S-1153, OTA04767, scale bar equals 10 µm).
FIGURE 15. Cuticle of fossil *Litsea* and *Beilschmiedia* (Lauraceae, all Rangitawa-6); 1. *Beilschmiedia* sp., TLM view (SL4979, scale bar equals 50 µm); 2. *Beilschmiedia* sp., TLM view (SL4979, scale bar equals 20 µm); 3. *Beilschmiedia* sp., SEM view of outer surface (S-1654, scale bar equals 10 µm); 4. *Beilschmiedia* sp., SEM view of inner surface (S-1654, scale bar equals 10 µm); 5. *Litsea* sp., TLM view (SL4982, scale bar equals 50 µm); 6. *Litsea* sp., TLM view, detail of stomatal complexes (SL4982, scale bar equals 20 µm); 7. *Litsea* sp., SEM view of outer surface with two stomatal complexes. Note distinctive raised aspect of the subsidiary cells (SEM3957, scale bar equals 10 µm); 8. *Litsea* sp., SEM view of inner surface (SEM3959, scale bar equals 10 µm).
Identification. Relatively large and aligned stomates and trichome bases associated with one or two modified epidermal cells identify it as Toronia toru (see figures in Pole, 1998; Carpenter et al., 2010).

Knightia excelsa
Figure 17.3-4

Specimens. SL4514, Pahurehure.

Identification. Randomly-oriented stomates and trichome attachment sites with over two to four modified epidermal cells and very sinuous epidermal cell flanges can be directly related to the extant Knightia excelsa (see figures in Pole, 1998).

Fabaceae
Carmichaelia sp.
Figure 18

Specimens. SL0204, Dargaville; SL2379, Maxwell-07; SL2351, Mokau; SL2365, Mokau; SL4515, Pahurehure; LX2251, Ohawe-4; LX2304, Rangitawa-4; SL4642, Rangitawa-5; SL5946, Rangitawa-6; SL0228, Scottys Camp.

Identification. This cuticle morphology is common and widespread. It is distinctive in having transversely-oriented stomates, organised in loose rows or narrow zones, and in elongate, flattened structures. Transversely-oriented stomates are such a restricted character that an unusual organ – i.e., a cladode, was immediately suggested. There was an initial suspicion that this may be the macro fossil corresponding to the *Acaciapollenites* pollen, which has been reported as from several Plio-Pleistocene localities in the North Island, including those covered here (Mildenhall, 1975a, b). However, *Acacia* was subsequently ruled out, and an extant New Zealand taxon, the cladodes of *Carmichaelia*, was recognised as the match. The species is not yet clear. Heenan (1998) recognised 23 species of *Carmichaelia* indigenous to New Zealand (and noted that they have transverse stomates).
The cuticle fragments suggest one a species with larger cladodes, such as *C. australis*. This species would be consistent with the peaty facies (many other extant *Carmichaelia* are species of drylands), but the fossils have trichome-bases, which have not been observed on *C. australis*.

**Grisseliniaceae**  
*Grisellinia lucida*  
Figure 19

**Identification.** The relatively large, round stomatal complexes, with a prominent Outer Stomatal Ledge and relatively isodiametric epidermal cell outlines identify this as *Grisellinia lucida* (see Pole 2008a, figure 35).

**Asteraceae**  
*Olearia* sp.  
Figure 20

**Specimen.** SL5480, Hamiltons Gap.  
**Identification.** The key characters are an ovoid stomatal complex surrounded by a sharply-defined, narrow rim, and a mixture of abundant smooth, discrete papillae on the epidermal cells and trichome bases. This combination of characters is found in some of the extant New Zealand *Olearia*, for example *O. arborescens*.

**Myrsinaceae**  
*Myrsine* sp.  
Figure 21

**Specimen.** SL4655, Rangitawa-6.  
**Identification.** Cuticle with anisocytic stomatal complexes with either a pervasive ornamentation of ridges or restricted to the subsidiary cells is probably *Myrsine*. The peltate trichomes that would confirm the identification (see Pole, 2008a) have not been noted.
FIGURE 18. Cuticle of Carmichaelia sp. (Fabaceae): 1. TLM view showing stomatal complexes. The plant axis in the image is horizontal and the stomatal orientation is transverse (SL2365, Mokau, scale bar equals 100 µm); 2. TLM view, detail of stomatal complex (SL2365, Mokau, scale bar equals 20 µm); 3. SEM view of inner surface of a stomata complex (S-1648, Rangitawa-6, scale bar equals 10 µm); 4. SEM view of outer surface of a stomatal complex (S-1648, Rangitawa-6, scale bar equals 10 µm); 5. TLM view of stomatal complexes and trichome bases near upper and lower margins (SL2379, Maxwell-07, scale bar equals 50 µm); 6. TLM view of stomatal complexes and a trichome base at lower left (SL2379, Maxwell-07, scale bar equals 50 µm).
CONCLUSIONS

Leaf cuticle has been found to be a common component in the Plio-Pleistocene 'lignites' in the North Island. A variety of taxa is represented, and is almost entirely of extant New Zealand species of forest and wetland angiosperms and conifers. However, some taxa are outside of their present-day range - the araucarian specimen at Rangitawa is notably far south of the range of Agathis today, while the presence of Halocarpus at sea level in the North Island is probably lower than it would occur today at that latitude. The existence of Cryptocarya, a possible Araucaria, and an indeterminate conifer from Mokau, support palynological evidence, of relatively recent plant extinctions in New Zealand.

There is clear scope for further work on this kind of material. Some sites still need to be dated, for example, Huriwai, and refining some identifications to species level would be useful. The New Zealand record has potential for providing stomatal indice variation of extant taxa, and thus further refining atmospheric carbon dioxide levels over several glacial-interglacial cycles. The macrofossils belonging to the Acacia pollen that stimulated this work still await discovery.

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FIGURE 20. Cuticle of fossil and extant Olearia sp. (Asteraceae); 1. TLM view of stomatal complexes, with papillae and trichomes between (SL5480, Hamiltons Gap, scale bar equals 50 µm); 2. TLM view, detail of two stomatal complexes, note the prominent peristomal rims (SL5480, Hamiltons Gap, scale bar equals 20 µm); 3. SEM view of inner surface, note hollow papillae (S-1659, Hamiltons Gap, scale bar equals 10 µm); 4. SEM view of outer surface, note common papillae and trichomes (S-1659, Hamiltons Gap, scale bar equals 20 µm); 5. Extant Olearia arborescens, TLM view of stomatal complexes, papillae and trichomes (OTA43559, scale bar equals 100 µm); 6. Extant Olearia rani, TLM view of stomatal complexes, papillae and trichomes (OTA23174, scale bar equals 50 µm).
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